

1 Data assimilation experiments inform monitoring needs for near-term ecological forecasts in a
2 eutrophic reservoir

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13 **Abstract:**

14 Ecosystems around the globe are experiencing increased variability due to land use and climate

15 change. In response, ecologists are increasingly using near-term, iterative ecological forecasts to

16 predict how ecosystems will change in the future. To date, many near-term, iterative forecasting

17 systems have been developed using high temporal frequency (minute to hourly resolution) data

18 streams for assimilation. However, this approach may be cost-prohibitive or impossible for

19 forecasting ecological variables that lack high-frequency sensors or have high data latency (i.e., a

20 delay before data are available for modeling after collection). To explore the effects of data

21 assimilation frequency on forecast skill, we developed water temperature forecasts for a

22 eutrophic drinking water reservoir and conducted data assimilation experiments by selectively

23 withholding observations to examine the effect of data availability on forecast accuracy. We used

24 in-situ sensors, manually collected data, and a calibrated water quality ecosystem model driven
25 by forecasted weather data to generate future water temperature forecasts using FLARE
26 (Forecasting Lake And Reservoir Ecosystems), an open-source water quality forecasting system.
27 We tested the effect of daily, weekly, fortnightly, and monthly data assimilation on the skill of 1
28 to 35-day-ahead water temperature forecasts. We found that forecast skill varied depending on
29 the season, forecast horizon, depth, and data assimilation frequency, but overall forecast
30 performance was high, with a mean 1-day-ahead forecast root mean square error (RMSE) of
31 0.94°C, mean 7-day RMSE of 1.33°C, and mean 35-day RMSE of 2.15°C. Aggregated across
32 the year, daily data assimilation yielded the most skillful forecasts at 1-7-day-ahead horizons, but
33 weekly data assimilation resulted in the most skillful forecasts at 8-35-day-ahead horizons.
34 Within a year, daily to fortnightly data assimilation substantially outperformed monthly data
35 assimilation in the stratified summer period, whereas all data assimilation frequencies resulted in
36 skillful forecasts across depths in the mixed spring/autumn periods for shorter forecast horizons.
37 Our results suggest that lower-frequency data (i.e., weekly) may be adequate for developing
38 accurate forecasts in some applications, further enabling the development of forecasts broadly
39 across ecosystems and ecological variables without high-frequency sensor data.

40

41 **Key Words:** data collection frequency; FLARE; high-frequency sensors; initial conditions;
42 observations; uncertainty; water temperature

43

44 **Introduction**

45 In the face of increasing ecological variability due to climate and land use change (e.g.,
46 Gilarranz et al., 2022, Malhi et al., 2020), ecological forecasting is increasingly being used for

47 understanding and predicting future ecological change (Carey et al., 2022d, Lewis et al., 2022).
48 Here, we define ecological forecasts as predictions of future environmental conditions with
49 quantified uncertainty (see Carey et al., 2022d, Lewis et al., 2022). Applications of ecological
50 forecasts can improve understanding of ecosystem processes (e.g., carbon cycling, Bett et al.,
51 2020), quantify predictability of environmental variables (e.g., rodent abundances, White et al.,
52 2019), and inform management of ecosystem services (e.g., fisheries management, Lindegren et
53 al., 2010). Because of their broad utility, forecasts are increasingly being developed by the
54 research community to predict population, community, and ecosystem dynamics (Lewis et al.,
55 2022). For example, an ongoing, community-based forecasting challenge organized by the
56 Ecological Forecasting Initiative's Research Coordination Network has received thousands of
57 ecological forecast submissions of National Ecological Observatory Network (NEON) data (e.g.,
58 lake water temperature, tick abundances, forest net ecosystem production, beetle communities)
59 before the data have been collected (Thomas et al., 2023a).

60 Many near-term (daily to decadal) ecological forecasts are produced using the iterative,
61 near-term forecasting cycle, in which models are updated as new observational data become
62 available to generate forecasts into the future with quantified uncertainty (Dietze et al., 2018).
63 The process of updating forecast models with newly available data, termed data assimilation
64 (DA), is a critical component of the iterative, near-term forecast cycle (Dietze et al., 2018, Luo et
65 al., 2011). DA allows for iterative updating of ecological hypotheses and models as forecasts are
66 continuously assessed and updated with the most recent ecosystem observations (Dietze et al.,
67 2018; White et al., 2019). DA can also improve forecast accuracy by updating forecast model
68 initial conditions (i.e., starting values given to the model), states, and/or parameters at the
69 timestep that the new observations become available (Cho et al., 2020, Gottwald and Reich,

70 2021, Luo et al., 2011, Niu et al., 2014). For example, McClure et al. (2021) developed 1 to 2-
71 week-ahead forecasts of reservoir methane emissions with and without weekly DA and found
72 that the accuracy of forecasts with DA was 44 - 128% higher than forecasts without DA over a
73 five-month forecasting period. Despite the usefulness of DA for improving forecasts, however,
74 the optimal frequency of observations for updating ecological models to produce skillful
75 forecasts is not well characterized.

76 While there are a number of best practices proposed for applying the near-term, iterative
77 forecast cycle in ecology (e.g., Clark et al., 2001, Harris et al., 2018, Lewis et al., 2022, White et
78 al., 2019), few recommendations exist for choosing the optimal frequency of DA to produce
79 accurate forecasts. Specifically, determining the appropriate frequency of observations for DA
80 across a range of ecological variables is needed to improve the scalability of ecological
81 forecasting, particularly if accurate forecasts can be developed using lower frequency
82 observations. For example, if weekly or fortnightly DA yielded similarly accurate lake dissolved
83 oxygen forecasts as daily DA, then water quality forecasting systems could be developed for
84 lakes that have weekly or fortnightly routine monitoring program data without needing expensive
85 high-frequency sensors, thereby enabling forecasts to be generated for many waterbodies
86 globally.

87 Currently, many automated ecological forecasting systems rely on high-frequency
88 sensors to assimilate data at each time step and generate accurate forecasts (e.g., Baracchini et
89 al., 2020b, Corbari et al., 2019, Marj and Meijerink, 2011, Page et al., 2018, Tanut et al., 2021),
90 but it is possible that high-frequency sensor data collection may not be needed for DA.
91 Moreover, deployment of high-frequency sensors is not always feasible for all ecological
92 variables (e.g., zooplankton abundance, biogeochemical concentrations, Marcé et al., 2016) and

93 some remote locations have additional logistical constraints for maintaining autonomous sensor
94 operation (Steere et al., 2000). Furthermore, some remotely sensed variables may only be
95 available as satellite orbits and weather conditions (e.g., cloud cover) allow (e.g., Herrick et al.,
96 2023). Thus, identifying how best to integrate observational data collected at different temporal
97 frequencies into forecast models has emerged as a critical need for ecological forecasters
98 (LaDeau et al., 2017).

99 Studies on the frequency of DA for environmental forecasts have generally shown that
100 more temporally frequent DA improves forecast accuracy, but not always, which may be related
101 to the sensitivity of forecasts to model initial conditions. For example, DA occurring every 24
102 hours using in-situ snow data (e.g., snow depth, density, snow water equivalent) resulted in
103 better predictions of these snow variables in an alpine snowpack model compared to DA
104 occurring every 3 hours (Piazzini et al., 2018). Conversely, DA ‘experiments’ performed for
105 NOAA (National Oceanic and Atmospheric Administration)’s Global Forecasting System using
106 meteorological observations collected at different frequencies showed that DA occurring every 2
107 hours resulted in more accurate air temperature and wind speed forecasts compared to DA
108 occurring every 6 hours (He et al., 2020). These differences are likely because uncertainty in
109 meteorological forecasts is primarily driven by the forecast model’s initial conditions. Thus,
110 more frequent DA, which constrains the model’s initial conditions, will almost always improve
111 the skill of meteorological forecasts (e.g., Clark et al., 2016, He et al., 2020, Simonin et al.,
112 2017). In contrast, for forecasts of environmental systems in which model process uncertainty
113 and model driver data uncertainty are more important sources of uncertainty (e.g., Dietze, 2017a,
114 Heilman et al., 2022, Lofton et al., 2022, Thomas et al., 2020), it is unknown whether more

115 frequent DA can improve forecast skill by generating initial conditions more consistent with
116 observations.

117 To the best of our knowledge, there have been only a few ecological DA experiments that
118 have tested the effects of different observation frequencies on forecast skill (e.g., Massoud et al.,
119 2018, Piazzzi et al., 2018, Weng and Luo, 2011, Ziliani et al., 2019), and none that have
120 considered how the frequency of data used for assimilation affects forecast skill across both
121 spatial and temporal scales. Weng & Luo (2011) assimilated eight different carbon datasets (e.g.,
122 root biomass, litter fall, soil respiration), each with different collection frequencies, to identify
123 the relative importance of these data sources in constraining long-term carbon dynamics, but did
124 not consider how different frequencies of the same dataset could affect forecast skill. Piazzzi et al.
125 (2018) assimilated multiple snow observations at two different frequencies (3 and 24 hours) for
126 predicting different snow-related variables (e.g., depth, density, and snow water equivalent), and
127 Ziliani et al. (2019) performed DA tests using 1-20 second assimilation of water depth data to
128 assess water level forecast skill, but neither considered the effect of less frequent assimilation
129 (e.g., >24 hours). Massoud et al. (2018) performed DA tests using a wider range of temporal
130 frequencies (e.g., ~3-34-day abundance data) to predict plankton community dynamics, but did
131 not consider the effects of DA across spatial scales (i.e., how DA affects forecast skill across
132 multiple sites or depths in an aquatic ecosystem). As a result, further work is needed to quantify
133 the utility of increased observation and DA frequency over both time and space to forecast
134 performance in ecological systems with varying sensitivities to initial conditions.

135 Among ecosystems, freshwater lakes and reservoirs are particularly important systems
136 for developing near-term forecasts because they provide essential ecosystem services, including
137 drinking water, food, irrigation, and recreation (Carpenter et al., 2011, Meyer et al., 1999,

138 Williamson et al., 2016). Because freshwaters are experiencing greater variability and adverse
139 water quality issues in response to land use and climate change (e.g., O'Reilly et al., 2015, Paerl
140 and Paul, 2012, Woolway et al., 2021), some water managers have used forecasts to
141 preemptively address poor water quality events (reviewed by Lofton et al., 2023). To date,
142 iterative, near-term freshwater forecasts have been developed for a number of water quality
143 variables, including water temperature (e.g., Carey et al., 2022d, Thomas et al., 2023b),
144 dissolved oxygen (e.g., Wang et al., 2016), and phytoplankton (e.g., Page et al., 2017, Woelmer
145 et al., 2022). These forecasts have been developed using DA with observations collected by
146 high-frequency sensors at intervals ranging from 4 minutes to 24 hours. However, most manual-
147 sampling water quality monitoring programs collect observations on weekly to fortnightly scales
148 (e.g., Francy et al., 2015, Kirchner and Neal, 2013, Romero et al., 2002), currently precluding the
149 scaling of existing forecasting systems broadly and underscoring the need to determine whether
150 less frequent observations can be used to produce accurate forecasts.

151 To quantify how DA at different frequencies affects forecast skill up to 35 days into the
152 future, we performed DA experiments in which we separately assimilated daily, weekly,
153 fortnightly, and monthly data into reservoir water temperature forecasts. Water temperature
154 forecasts are used to inform management decisions on water extraction depth and preemptive
155 water quality interventions (Georgakakos et al., 2005; Kehoe et al., 2015; Mi et al., 2020), and
156 thus our study has much utility for both informing how best to forecast complex ecosystem
157 dynamics, as well as manage drinking water supplies. Our research questions were: 1) Which
158 frequency of DA generates the most skillful water temperature forecasts? 2) How does forecast
159 skill vary across time (specifically focusing on the mixed vs. stratified seasons within a year) and
160 space (i.e., reservoir depth)? and 3) How does DA frequency influence total forecast uncertainty

161 and what is the relative contribution of initial condition uncertainty to total forecast uncertainty?
162 As previous work has suggested that reservoir water temperature forecasts can sometimes exhibit
163 sensitivity to initial conditions (Thomas et al. 2020), we expected that less frequent DA would
164 result in decreased forecast skill and increased total uncertainty. In addition, we expected that
165 forecast skill would be better at deeper depths, especially during thermally-stratified periods
166 (e.g., Mercado-Bettín et al. 2021; Thomas et al. 2020).

167

168 **Methods**

169 *Forecasting system overview*

170 We applied the Forecasting Lake And Reservoir Ecosystems (FLARE) forecasting
171 system (Thomas et al., 2020) to Beaverdam Reservoir, Virginia, USA to produce daily water
172 temperature forecasts for 1-35 days into the future (hereafter referred to as forecast horizon)
173 during 1 January 2021 - 31 December 2021. FLARE is an open-source forecasting system that
174 incorporates real-time water quality sensor data, DA, ensemble-based forecasts, and uncertainty
175 quantification to predict near-term water quality conditions (Thomas et al., 2020).

176 Forecast generation via FLARE can be summarized by four steps (Figure 1). First, 10-
177 min resolution water temperature data were collected by sensors deployed in the reservoir
178 (Figure 1 step 1). Second, these data were transferred to the cloud and stored in a GitHub
179 repository, where they were downloaded daily and made available for DA (Figure 1 step 2).
180 Simultaneously, 1 to 35-day-ahead NOAA meteorological forecasts were downloaded daily as
181 driver data for the reservoir hydrodynamic model to generate the water temperature forecasts.
182 Third, during the forecast generation step, DA was used to update initial conditions and
183 parameters with the most recent observations using an ensemble Kalman filter, a numerical

184 approach that allows for the updating of model states and parameters using data (Evensen, 2003)
185 (Figure 1 step 3a). Following DA, the reservoir hydrodynamic model was initialized with the
186 updated model states and parameters to produce 1-35-day-ahead forecasts for each 0.5 m depth
187 interval across the water column (Figure 1 step 3b). Finally, forecast skill was assessed by
188 comparing observed vs. predicted water temperatures for each daily forecast at each depth
189 (Figure 1 step 4). We repeated steps 3a-4 for daily, weekly, fortnightly, and monthly intervals of
190 DA throughout the year as part of the DA experiments to compare forecast skill over time.

191

192 *Study site and monitoring*

193 Beaverdam Reservoir (BVR) is a small (0.28 km²), shallow ($Z_{\max} = 11$ m), dimictic,
194 eutrophic reservoir in southwestern Virginia, USA (37.31° N, 79.82° W; Figure 2). BVR is
195 managed by the Western Virginia Water Authority as a secondary drinking water supply and is
196 located in a deciduous forest catchment (Doubek et al., 2019). During a typical year, BVR is
197 stratified from mid-March to late October and mixed from November to early March (Hounshell
198 et al., 2021). BVR experiences summer hypolimnetic anoxia and cyanobacterial blooms, both of
199 which are controlled by water temperature and thermal stratification (Doubek et al., 2019; Hamre
200 et al., 2018), making forecasts of water temperature important for water quality management.

201 Water quality monitoring of BVR includes both manual sampling and high-frequency
202 sensors. From 2014-present, manual water quality sampling occurred weekly to fortnightly
203 during the summer stratified period and fortnightly to monthly during the remainder of the year
204 (Carey et al., 2022c). Starting in June 2020, high-frequency sensors were deployed in the
205 reservoir, enabling a range of DA frequencies to be compared in this study. We deployed
206 NexSens T-Node FR Temperature Sensors (NexSens Technology, Fairborn, OH, USA) at 1 m

207 intervals from the surface to sediments and a YSI EXO2 sonde (YSI Incorporated, Yellow
208 Springs, OH, USA) that monitored temperature at 1.5 m at the deepest site in BVR (Figure 1; see
209 Carey et al., 2023 for sensor information). These sensors collected data every 10 minutes, which
210 was transmitted every 3 to 9 hours via secure sensor gateways to a Git repository in the cloud
211 (Carey et al., 2023, Daneshmand et al., 2021). We removed observations collected during
212 periods of sensor maintenance, as well as depth-adjusted the data using an offset calculated from
213 a CS451 Stainless-Steel Pressure Transducer (Campbell Scientific, Logan, UT, USA) to account
214 for water level changes (Wander et al., 2023b). Because of this range in latency, or the time that
215 it takes for data to become available for modeling after they are initially collected, we used the
216 daily mean in our forecasting application. Following quality checks, these data were integrated
217 into the FLARE forecasting system to produce depth-specific daily water temperature forecasts.

218

219 *Hydrodynamic model configuration*

220 For modeling reservoir hydrodynamics, we used the General Lake Model (GLM) v.3.3.0
221 (Hipsey et al., 2022) to forecast water temperature in BVR. GLM is an open source, 1-D process-
222 based hydrodynamic model commonly used within the freshwater research community to
223 simulate water quality in lakes and reservoirs (Hipsey et al., 2019). GLM uses a Lagrangian
224 approach for simulating different water layers and has been applied to a variety of lakes
225 worldwide for modeling (e.g., Bruce et al., 2018, Read et al., 2014) and forecasting
226 hydrodynamics (e.g., Thomas et al., 2020, 2023b).

227 We configured GLM for BVR using historical bathymetric data (Carey et al., 2022b) and
228 water temperature observations for initial conditions (Carey et al., 2023). We configured GLM
229 with two sediment zones to simulate epilimnetic (surface) and hypolimnetic (bottom) sediment

230 temperature dynamics following Carey et al. (2022a). GLM requires meteorological and
231 reservoir inflow observations as driver data to run the model. Because we were applying GLM
232 for forecasting, meteorological forecasts, not observed meteorology, were used as driver data in
233 the model, as described below. Additionally, we set the inflow to equal outflow in this study
234 given limited inflow data for validation and the relatively short forecast horizons (≤ 35 days). We
235 initiated the model using its default parameter set (Hipsey et al., 2019) and performed calibration
236 via a 35-day spin-up period with DA to tune parameters before the start of our focal forecasting
237 period (described below).

238

239 *FLARE configuration for DA and uncertainty*

240 We configured FLARE for BVR following its application to other lakes and reservoirs
241 (Thomas et al., 2020, 2023b). We set the number of forecast ensemble members to 256 to ensure
242 an adequate representation of uncertainty and prevent the ensemble Kalman filter from
243 developing erroneous correlations among ensemble members that can occur with low ensemble
244 sizes (Duc et al., 2021, Machete and Smith, 2016). While we used default values for most GLM
245 parameters, we used the ensemble Kalman filter in FLARE to tune three model parameters that
246 we identified as important for water temperature simulations using GLM in a similar, nearby
247 reservoir (Carey et al., 2022a, Thomas et al., 2020): 1) the longwave radiation scaling factor
248 (hereafter, *longwave*); 2) epilimnetic sediment temperature parameter (hereafter, *epi_sed_temp*);
249 and 3) hypolimnetic sediment temperature parameter (hereafter, *hypo_sed_temp*).

250 We used state augmentation to tune the three parameters in the ensemble Kalman filter
251 (Thomas et al. 2020). Specifically, correlations between the parameter values and the model
252 states with observations (i.e., water temperatures at the depths with sensor observations) were

253 used to adjust parameters to be consistent with the most recent data used in DA. The three tuned
254 parameters were initially calibrated during a spin-up period from 27 November to 31 December
255 2020 and were subsequently updated via DA throughout the forecasting period. To avoid the
256 common issue of artificially low parameter uncertainty in sequential DA (Dietze, 2017a), we
257 specified the standard deviation of a normal distribution for each parameter (1.0°C for the
258 sediment temperature parameters and 0.02 for the longwave radiation scaling factor). Initial
259 exploration of parameter fitting in this study indicated that the application of FLARE over the
260 full year resulted in low parameter uncertainty, necessitating us to specify the standard deviation
261 a priori rather than estimating it using DA. The distributions we chose were adapted from a prior
262 application of FLARE that estimated the standard deviation of parameter distributions across six
263 lakes (Thomas et al., 2023b).

264 FLARE uses a numerical ensemble-based approach to simulate and propagate forecast
265 uncertainty (Thomas et al., 2020). We represented the contribution of uncertainty from
266 meteorological driver data, initial conditions, model process, and model parameters using the
267 256-member ensemble, following Thomas et al. (2020). First, to represent the contribution of
268 meteorological driver data uncertainty, we assigned each of the 256 FLARE ensemble members
269 one of the 30 ensemble members from the 1-35-day-ahead meteorological forecasts (National
270 Oceanic and Atmospheric Administration’s Global Ensemble Forecasting System) to drive GLM
271 for forecasting. Second, we represented uncertainty in the initial conditions of the forecasts using
272 the spread in model states among the 256 ensemble members on the first day of each forecast.
273 This spread was determined by either using the prior day’s forecast as a starting point for the
274 next day’s forecast (when no data were available for DA) or the updated states following DA
275 (when data were available for DA). We set the observation uncertainty standard deviation to

276 0.1°C, determined from the standard deviation of temperature observations and following prior
277 applications of FLARE (Thomas et al. 2020). Third, we represented model process uncertainty
278 by adding random noise to the water temperature predictions from each of the 256 FLARE
279 ensemble members at each daily time-step in a 1-35-day-ahead forecast horizon. The random
280 noise for each modeled depth within an ensemble member was drawn from a normal distribution
281 with a standard deviation of 0.75°C, as used in a previous application of FLARE that reported
282 well-calibrated forecast uncertainty (Thomas et al. 2020). The random noise was spatially
283 correlated so that it was most similar for nearby depths and most different for further-apart
284 depths. The strength of the spatial correlation was determined by the exponential decay of the
285 correlation strength with distance (Thomas et al., 2023b). Fourth, we represented parameter
286 uncertainty using the standard deviations of the distributions for the three tuned GLM parameters
287 described above. A unique parameter value drawn from each of the three distributions was
288 assigned to each of the 256 FLARE ensemble members. The parameter value assigned to an
289 ensemble member was only updated when DA occurred. Parameters not tuned by the ensemble
290 Kalman filter were assumed to have fixed values and uncertainty in these parameters was not
291 calculated.

292 To determine whether there was a relationship between the magnitude of initial
293 conditions uncertainty and the sensitivity of forecast skill to more frequent DA (following Clark
294 et al., 2016, He et al., 2020, Simonin et al., 2017), we quantified the contribution of initial
295 conditions uncertainty to total forecast uncertainty in our DA forecasts for all DA frequencies.
296 For this analysis, we isolated the magnitude of initial conditions uncertainty by generating the
297 water temperature forecasts for all 365 days with and without initial conditions uncertainty and
298 compared the variance among all 256 ensemble members. We also calculated the proportion of

299 initial conditions uncertainty within total forecast uncertainty for all depths, horizons, and
300 stratified vs. mixed periods.

301

302 *Data assimilation experiments*

303 To quantify the effect of DA at different frequencies on forecast skill, we conducted DA
304 experiments in BVR from 1 January to 31 December 2021 (n = 365 days). As noted above, we
305 used a spin-up period from 27 November - 31 December 2020 (n = 35 days) during which DA
306 occurred, but no forecasts were generated. During the one-year forecast period in 2021, we
307 forecasted daily water temperature at 23 depths in the reservoir (spanning 0.1 to 11 m depth at
308 0.5 m intervals) and assessed forecast performance relative to observations across each forecast's
309 daily predictions for 1 to 35-day-ahead horizons and depth intervals. We focused on three focal
310 depths (1, 5, and 9 m) when reporting results, as these depths are representative of the surface,
311 middle, and bottom layers of the water column, respectively. We chose 9 m to represent the
312 bottom of the reservoir because deeper depths were not always observed due to variability in
313 water levels throughout the year (within ~1 m due to seasonality in flows).

314 We performed DA experiments using four different DA frequencies (daily, weekly,
315 fortnightly, and monthly) to represent different data collection latencies that are commonly used
316 by water quality monitoring programs (e.g., Engelhardt and Kirillin, 2014, Francy et al., 2015,
317 Kirchner and Neal, 2013, Liu et al., 2019, Romero et al., 2002). We assimilated water
318 temperature data across different temporal frequencies by downsampling from the high-
319 frequency observations collected by our sensors. This resulted in four different temporal
320 frequencies for which DA occurred, corresponding to either daily (representing standard FLARE
321 DA), weekly, fortnightly, or monthly DA (see Data Assimilation Experiments box in Figure 1

322 and Appendix S1: Figure S1 for visualization of DA frequencies). For example, for the weekly
323 DA frequency, observations were selected every seven days starting on 4 January 2021 and
324 ending on 31 December 2021. In this example, DA only occurred once per week; the forecasts
325 that were generated on the six other days in the same week did not include DA (i.e., no DA
326 occurred during 5 January - 10 December 2021 even though forecasts were still generated daily
327 during this interval; Figure 1). Fortnightly and monthly DA occurred every 14 days and 30 days,
328 respectively, throughout the year.

329 We generated 365 daily forecasts starting on 1 January 2021 for each of the four DA
330 frequencies. While we recognize that we are producing hindcasts for a historical period, because
331 the model was forced with only forecasted drivers and out-of-sample forecast evaluation
332 occurred, we refer to these retrospective forecasts or hindcasts as forecasts throughout for
333 consistency (following Jolliffe and Stephenson, 2012).

334

335 *Analysis*

336 *Question 1:* For all $n = 1460$ forecasts produced (365 forecasts generated daily over a
337 year for four different DA frequencies), we used root mean square error (RMSE) and continuous
338 ranked probability score (CRPS; Gneiting et al., 2005) to quantify forecast skill. We defined
339 skillful water temperature forecasts as those with an $RMSE < 2^{\circ}C$, a commonly-used threshold
340 for lake and reservoir hydrodynamic modeling following Bruce et al. (2018), Read et al. (2014),
341 and many others. Mean full water column RMSE was calculated for each of the 35 days across
342 all forecast horizons for each DA frequency regardless of whether data were assimilated the day
343 the forecast was generated. We aggregated RMSE across depths and dates to determine the
344 lowest temporal frequency of DA required to generate the most skillful water temperature

345 forecasts across the full water column and throughout the entire forecast period. We focus on
346 RMSE in the results and all CRPS values are reported in the SI.

347 *Question 2:* Using RMSE and CRPS, we compared forecast skill across depths and
348 seasons to identify how the frequency of DA affected forecast accuracy over space and time. To
349 quantify spatial forecast performance, we calculated RMSE and CRPS for each depth (1-11 m) at
350 each forecast horizon (1-35 days ahead) and DA frequency in BVR. To quantify temporal
351 forecast performance, we compared forecast skill at each horizon aggregated within thermally-
352 stratified vs. mixed periods in BVR. The stratified period began on the first day that the water
353 density difference between the reservoir surface (0.1 m) and the maximum depth observed for
354 the reservoir on each day (e.g., between 9-11 m) was $\geq 0.1 \text{ kg/m}^3$ for at least three consecutive
355 days (following Ladwig et al., 2021). Conversely, the mixed period began on the first day that
356 surface and bottom water density differences were $< 0.1 \text{ kg/m}^3$ for at least three consecutive
357 days. Altogether, we compared forecast skill between stratified vs. mixed periods; among depths
358 (1, 5, and 9 m), and among forecast horizons (focusing in on 1, 7, and 35-day-ahead forecasts)
359 for each of the four DA frequencies.

360 *Question 3:* We quantified total forecast uncertainty for each day in the 1-35-day forecast
361 horizon using the variance of the 256-member FLARE ensemble. The relative contribution of
362 initial condition uncertainty to total forecast uncertainty was calculated for each forecasted day
363 by comparing the variance in the 256-member FLARE ensemble between the set of forecasts
364 with initial condition uncertainty included and the set without initial condition uncertainty.

365 All statistical analyses were conducted in R v.4.2.0 (R Core Team, 2022). All R code and
366 data files used to run these analyses are archived and available in the Zenodo repository (Wander
367 et al., 2023a, 2023b).

368 **Results:**

369 *BVR water temperature dynamics*

370 BVR exhibited typical annual water temperature dynamics during the forecasting period
371 in 2021. Water temperature throughout the water column ranged from 1.4 to 29.9°C during the
372 year. The summer stratified period began on 12 March and ended on 7 November 2021, and the
373 reservoir was mixed from 1 January - 11 March and 8 November - 31 December (Figure 3).
374 Thermocline deepening occurred throughout the summer stratified period, starting at 1.5 m in
375 March with stratification onset and deepening to 9.5 m in November before fall turnover (Figure
376 3). During the winter, there were three brief periods of ice cover of one to three days in duration
377 in January and February when inverse stratification occurred (Figure 3; Carey and Breef-Pilz,
378 2022). We removed these few ice-cover days from the analysis and grouped mixed (n = 118
379 days) vs. summer stratified data (n = 241 days) for analysis.

380

381 *Data assimilation frequency altered forecast output and parameters over time*

382 We were able to successfully forecast water temperature throughout the water column
383 over the year using DA to update model states and parameters (Figures 4-5). Across all depths,
384 DA constrained uncertainty by updating initial conditions with the most recent water temperature
385 observations. Forecast uncertainty for the lower DA frequencies was strongly dependent on the
386 time since last assimilation (Figure 4). On average, forecast variance at the one-day horizon
387 across 2021 for forecasts with daily DA was 1.56°C while mean forecast variance at the one-day
388 horizon for forecasts with monthly DA was 3.25°C.

389 We observed that DA frequency altered parameter evolution of the forecasts (Figure 5).
390 The daily DA frequency resulted in more variable parameter estimates through time for all three

391 tuned parameters, reflecting the more frequent adjustment that occurred each time data were
392 assimilated. Importantly, parameter evolution for forecasts with daily DA yielded very different
393 estimates than the weekly, fortnightly, and monthly DA forecast frequencies (Figure 5). For
394 example, the evolution of the longwave radiation scaling parameter (longwave) over the 365-day
395 forecast period showed that forecasts with weekly, fortnightly, and monthly DA frequencies
396 converged at ~ 0.91 by December 2021, whereas the longwave parameter for forecasts with daily
397 DA was at ~ 0.85 by the end of the year (Figure 5a). Similarly, the parameter controlling the
398 surface layer sediment temperature (epi_sed_temp) in daily DA forecasts began to diverge from
399 the other DA frequencies in early April (Figure 5c). The non-daily DA frequencies (i.e., weekly,
400 fortnightly, monthly DA) surface sediment layer temperature parameter (epi_sed_temp) values
401 ranged from 13.51 to 15.62°C, whereas the daily DA frequency epi_sed_temp values ranged
402 from 13.29°C to 17.0°C during April-December. For the parameter controlling the bottom layer
403 sediment (hypo_sed_temp), daily DA forecasts exhibited much more variable values (ranging
404 from 10.24°C to 11.21°C) than forecasts for any other DA frequency (range 10.65°C to 10.72°C;
405 Fig 5b) from April to December.

406

407 *Question 1: Which frequency of data assimilation generates the most skillful water temperature*
408 *forecasts?*

409 Aggregated among depths and time periods, weekly DA resulted in the most skillful
410 water temperature forecasts of the four DA frequencies for the greatest number of 1-35-day-
411 ahead horizons (Figure 6). Among horizons, we observed that the frequency of DA needed to
412 produce skillful forecasts varied (Figure 6). At shorter horizons (1-7 days ahead), daily DA

413 resulted in the most skilled forecasts, but at longer horizons (8-35 days ahead), weekly DA
414 resulted in the most skilled forecasts (Figure 6).

415 The skill of all forecasts degraded as the forecast horizon increased, but the decrease in
416 performance was greatest for daily DA forecasts, such that forecasts generated using monthly,
417 fortnightly, and weekly DA all outperformed daily DA forecasts by the 19-day forecast horizon
418 (Figure 6), when aggregating across all depths and time periods. The daily DA forecasts
419 exceeded the 2°C RMSE metric of skill on the 28-day-ahead horizon, whereas the weekly,
420 fortnightly, and monthly forecasts never exceeded that metric for any of the 1-35-day-ahead
421 horizons. These results were consistent across forecast evaluation metrics, including the CRPS
422 metric that evaluates the full ensemble forecast (Appendix S1: Figure S2).

423

424 *Question 2: How does forecast skill vary across time and space?*

425 Aggregated across depths, horizons, and DA frequencies over the year, forecast skill
426 overall was high, with a mean water temperature forecast RMSE of $1.53 \pm 1.86^\circ\text{C}$ (1 S.D.).
427 Forecast skill was generally best at 9 m regardless of horizon or DA frequency. Aggregated 9 m
428 forecast skill was $1.29 \pm 1.80^\circ\text{C}$, followed by aggregated 5 m forecast skill ($1.63 \pm 1.85^\circ\text{C}$), and
429 aggregated 1 m forecast skill ($1.69 \pm 1.82^\circ\text{C}$). As expected, forecast skill generally decreased
430 with horizon, with a mean 1-day-ahead forecast RMSE of $0.80 \pm 1.20^\circ\text{C}$, mean 7-day RMSE of
431 $1.15 \pm 1.60^\circ\text{C}$, and mean 35-day RMSE of $1.99 \pm 2.17^\circ\text{C}$. However, we observed an exception to
432 this pattern for 1 m mixed forecasts, which is further described below.

433 On average, forecast skill was slightly better (as indicated by smaller RMSE) during the
434 stratified period than during the mixed period, aggregated among all depths and horizon
435 regardless of DA frequency (aggregated mixed RMSE = $1.59 \pm 1.57^\circ\text{C}$, stratified RMSE = 1.46

436 $\pm 2.13^{\circ}\text{C}$; Figure 7). Forecast skill was more variable among forecast horizons than depths in the
437 mixed period, whereas forecast skill was variable across both depths and horizons in the
438 stratified period (Figure 7). In the stratified period, forecast skill was best at 9 m, with relatively
439 similar skill over the forecast horizon (Figure 7f). In the mixed period, forecast skill varied very
440 little among depths aggregated across horizons (Figure 7a, c, e), with consistently greater
441 decreases in skill with increasing horizon than in the stratified period, except for at 1 m. Forecast
442 skill at 1 m decreased rapidly until ~the 19-day horizon, after which forecast skill remained
443 constant for the daily DA and increased for the weekly, fortnightly, and monthly DA frequencies
444 until the end of the forecast horizon (Figure 7a).

445 While daily DA always resulted in the best forecast skill for 1-day-ahead horizons, lower
446 frequency DA typically outperformed daily DA as the forecast horizon increased. An exception
447 was for 9 m stratified forecasts, when daily DA resulted in the lowest RMSE for all forecast
448 horizons and never exceeded 1.51°C for the duration of the 35-day forecast horizon (Figure 7f).
449 Additionally, 9 m stratified forecasts were the only forecasts where skillful ($\text{RMSE} < 2^{\circ}\text{C}$)
450 forecasts were produced for all DA frequencies and horizons (Figure 7f).

451

452 *Question 3. How does DA frequency influence total forecast uncertainty and what is the relative*
453 *contribution of initial condition uncertainty to total forecast uncertainty?*

454 Lower frequency DA forecasts consistently had more total uncertainty (Figure 8). We
455 found that the differences between uncertainty for daily and monthly DA were largest at 1-day-
456 ahead horizons and largely converged by the end of the 35-day horizon (Figure 8). At 1 m depth,
457 total uncertainty was similar between the mixed and stratified periods across the 35-day horizon,
458 but at 5 and 9 m, total uncertainty was on average higher in the stratified than mixed period. Both

459 RMSE and total variance were similar for forecasts run with and without initial conditions
460 uncertainty included (Appendix S1: Figures S3-S4).

461 Forecasts with less frequent DA had a greater contribution of initial condition uncertainty
462 to total forecast uncertainty during the first few days of the forecast horizon. However, overall,
463 initial conditions uncertainty contributed a minimal proportion of the total uncertainty for
464 forecasts generated with daily DA (Figure 9). At the 1-day-ahead forecast horizon, daily DA
465 initial conditions uncertainty was 0% of total uncertainty, whereas initial conditions uncertainty
466 contributed 55 - 71% of total forecast uncertainty in forecasts for all other DA frequencies
467 (Figure 9). The role of initial conditions uncertainty for all depths in the mixed period and
468 surface forecasts in the stratified period was minimal (< 1%) across all DA frequencies after the
469 10-day horizon (Figure 9a-c, e). Conversely, initial conditions uncertainty made up a larger
470 proportion of total forecast uncertainty for stratified 5 m and 9 m forecasts for forecast horizons
471 between 10 and 20 days (ranging from 5-10%; Figure 9d, f).

472

473 **Discussion:**

474 Across a year of water temperature forecasts in our focal reservoir, we found that weekly
475 DA generally resulted in the most skillful water temperature forecasts. However, skill varied
476 among depths, forecast horizons, and time of year, suggesting that DA frequency should be
477 chosen based on the specific forecast application. For example, if water temperature forecasts are
478 specifically needed to guide decision-making that involves the deeper reservoir layers (e.g., 5 m
479 or 9 m) at short time horizons (e.g., <5 days ahead), daily DA might be most advantageous
480 (Figure 7). Conversely, if water temperature forecasts are needed for the surface water at 20-35
481 day-ahead horizons, then weekly to monthly DA may be sufficient (Figure 7). Despite the

482 usefulness of DA for improving forecast skill, more frequent DA did not always lead to more
483 skillful water temperature forecasts, in part because initial conditions uncertainty only comprised
484 a significant proportion of total forecast uncertainty within the first few days of the forecast
485 horizon (Figure 9). Below, we interpret our results for each research question and make
486 recommendations for considering which DA frequency might be appropriate for different
487 ecological forecast applications.

488

489 *Q1: Which frequency of data assimilation generates the most skillful water temperature*
490 *forecasts?*

491 In this study we found that less frequent DA (e.g., weekly, fortnightly, and monthly DA)
492 sometimes led to more skillful water temperature forecasts than daily DA for all depths during
493 the mixed period. This pattern of weekly DA outperforming daily DA forecast skill during the
494 mixed period is likely because daily DA led to parameter overfitting, as indicated by the greater
495 short-term variability in parameter estimates over time (Figure 5). Because water temperatures
496 are fairly stable at deeper depths, and thus daily observations can consistently predict tomorrow's
497 water temperature accurately, parameter overfitting was less problematic for daily DA at
498 hypolimnetic depths (Figure 7f). As a result, hypolimnetic forecast skill was best with daily DA
499 during stratified conditions, but this pattern did not extend to other depths or the mixed period
500 (Figure 7).

501 Our work is consistent with studies that have found that the optimal DA frequency often
502 matches that of the forecast model timestep (e.g., Derot et al., 2020, Woelmer et al., 2022). For
503 example, during both the mixed and stratified periods, daily DA was always better for 1-day-
504 ahead forecasts, but was often outperformed by weekly DA at 8-day-ahead forecast horizons

505 (Figure 6). Because water temperatures were homogenous among all depths during the mixed
506 period, water temperature variability among all depths was likely driven by air temperature
507 variability, ultimately making it more challenging to predict water temperature across depths as
508 the forecast horizon increased. During the stratified period, however, less frequent DA could still
509 generate accurate surface and mid-depth water temperature forecasts. The increased importance
510 of daily DA at bottom depths during the stratified period is likely because of the increased
511 thermal stability at bottom depths associated with thermal stratification (Figure 7). This pattern is
512 in contrast with other water temperature forecasting studies that have found daily DA necessary
513 for improving the skill of forecasts in the middle of the water column around the thermocline
514 (Baracchini et al., 2020a), but is likely explained by the overfitting of both the daily longwave
515 radiation and the epilimnetic sediment temperature parameters (Figure 5).

516 We note that there are many ways to quantify skill beyond the 2°C RMSE threshold used
517 here. We chose to use RMSE because it is a commonly used metric by lake modelers to
518 determine the deviation between observed vs. modeled values (Bruce et al., 2018, Read et al.,
519 2014). However, forecast skill could also be quantified by other metrics, such as the correlation
520 coefficient, Nash–Sutcliffe model efficiency coefficient, percent relative error, normalized mean
521 absolute error, or others (Bennett et al., 2013). While 2°C RMSE is a subjective criterion of
522 forecast skill, we note that CRPS results followed similar patterns as our RMSE metric
523 (Appendix S1: Figure S2), further supporting our results when evaluating the full distribution of
524 the forecasts.

525

526 *Q2: How does forecast skill vary across time and space?*

527 Our use of a 35-day forecast horizon allowed us to compare water temperature

528 predictability across multiple horizons at different depths and times of year, thereby elucidating
529 patterns in ecosystem predictability across both space and time. We generally observed expected
530 declines in forecast skill with increasing horizon, as noted in many other studies. However, 1-m
531 forecast skill during the mixed period increased with lower frequency DA (weekly, fortnightly,
532 and monthly), while forecast skill with daily DA leveled off at 19-35-day horizons (Figure 7a).
533 Improved forecast skill for lower frequency DA suggests that our forecasts are capturing surface
534 water temperature dynamics in the mixed period at longer horizons better than other depths,
535 particularly those during the stratified period. This may be due to the smaller range in water
536 temperature variation that occurs in the mixed period relative to the stratified period over a 35-
537 day interval, allowing variance to level off at longer forecast horizons as water temperature
538 observations better matched predicted values (Figures 7-8). While improved forecast skill at
539 longer horizons has been observed in the literature (e.g., Wheeler et al., 2023), this pattern is
540 often associated with variables that have predictable, cyclical patterns at long horizons (e.g.,
541 annual tree leaf-out).

542 Overall, we observed generally high forecast skill across all depths and times of year for
543 most forecast horizons. Across DA frequencies, depths, and times of year, RMSE was only
544 consistently above the 2°C threshold for daily DA at 28-35-day horizons (Figure 6). By the end
545 of the 35-day forecast horizon, daily DA forecast skill for most depths and times of the year was
546 >2°C, except 9 m stratified forecasts, which had a mean RMSE of $1.29 \pm 1.8^\circ\text{C}$ across DA
547 frequencies. The higher forecast skill at 9 m is likely because fluctuations in bottom water
548 temperatures were minimal during stratification (Figure 3).

549 Our findings are similar to other water temperature lake and reservoir forecasting studies.
550 First, the pattern of increased forecast skill in the bottom waters is consistent with Mercado-

551 Bettín et al. (2021) and Thomas et al. (2020), who both found that the bottom water forecasts
552 were more skillful than surface water forecasts. This is likely because bottom waters are not
553 changing as much as surface waters throughout the year due to less atmospheric exchange.
554 However, Clayer et al. (2023) found that surface water temperatures were more accurately
555 simulated than bottom water temperatures, suggesting that the complex lake characteristics that
556 control bottom water temperatures were not captured as well as the air temperature dynamics
557 controlling surface water temperatures. Second, our finding that forecast skill was greater in the
558 stratified period rather than mixed period is similar to the results of Thomas et al. (2020), likely
559 due to the fact that water temperature dynamics were changing less among depths in stratified
560 than mixed periods (Figure 3). Because of the variability in water temperature dynamics among
561 seasons and depths, determining the conditions in which we can most accurately forecast water
562 temperature can improve our understanding of ecosystem processes and functioning. Moreover,
563 accurately forecasting water temperature is critical for forecasting additional lake and reservoir
564 variables that are strongly driven by water temperature, such as phytoplankton biomass,
565 dissolved oxygen concentrations, and greenhouse gas emissions (e.g., McClure et al. 2021).

566

567 *Q3: How does DA frequency influence forecast uncertainty?*

568 We found that initial conditions uncertainty contributed a substantial proportion of total
569 uncertainty for weekly, fortnightly, and monthly DA, but only during the first few days of the
570 forecast horizon. From 6-23 day-ahead horizons, the contribution of initial conditions decreased
571 to <1% across all DA frequencies, depths, and times of year (Figure 9). We observed that high-
572 frequency DA was required for skillful 9 m stratified forecasts, while weekly DA was sufficient
573 for other depths and times. This finding may be because the contribution of initial conditions

574 uncertainty decreases more rapidly within the first few days of the forecast horizon for the daily
575 DA forecasts at 9 m in the stratified period. For all other depths and times of the year, the rate at
576 which initial conditions uncertainty decreases is greater for weekly, fortnightly, and monthly
577 DA, resulting in more similar performance of daily and weekly DA early in the forecast horizon
578 (Figure 9). However, more frequent DA may not always improve forecast performance,
579 especially when initial conditions uncertainty is not the dominant source of uncertainty, as seen
580 at longer horizons. Given that initial conditions uncertainty predominated at the beginning of the
581 forecast horizon, it is likely that total forecast uncertainty at longer horizons was primarily
582 influenced by uncertainty in model process, model parameters, and/or meteorological driver data
583 (Figure 9). Conversely, the dominant source of uncertainty for weather forecasting is typically
584 initial conditions uncertainty given the inherent instability of atmospheric processes (Dietze,
585 2017b), which is why more frequent DA often substantially improves meteorological forecast
586 skill.

587 Other lake and reservoir water quality forecasting studies have found that model driver
588 data and process uncertainty were the dominant sources of total forecast uncertainty (Lofton et
589 al., 2022, McClure et al., 2021, Thomas et al., 2020). Therefore, constraining other sources of
590 uncertainty by using an ensemble approach or different forecasting models would likely further
591 improve water temperature forecast skill. Additionally, using a different DA technique that uses
592 a Bayesian approach to estimate a posterior distribution, rather than assuming that the parameters
593 and model states are normally distributed, may also reduce uncertainty (e.g., particle filter; Wang
594 et al., 2023). Because the dominant source of uncertainty in ecological forecasts will likely differ
595 depending on the variable being forecasted, different DA techniques may not improve forecast
596 skill equally among all ecological variables.

597

598 *Recommendations for setting up DA for other forecasting systems*

599 Determining whether an ecological forecasting application requires high-frequency
600 sensors is necessary for increasing the scalability of ecological forecasting across ecosystems and
601 variables. While high-frequency sensor data may improve forecast skill in some cases, sensor
602 deployment is often costly, which limits the application of high-frequency data in some
603 forecasting systems. Moreover, even if high-frequency sensors are deployed, identifying the
604 minimum frequency of data required to make skillful ecological forecasts can be a useful
605 exercise because high-frequency sensors malfunction and require maintenance, which can result
606 in data gaps (e.g., Herrick et al., 2023). Many water quality forecasting applications to date have
607 relied on high-frequency sensor data for assimilation to produce skillful forecasts of different
608 variables (Cho and Park, 2019, Derot et al., 2020, Page et al., 2018). In this study, we found that
609 daily DA only produced the most skillful 9 m stratified period water temperature forecasts,
610 whereas weekly DA generally produced the most skillful surface and middle layer water
611 temperature forecasts (Figure 7). Our findings indicate that high-frequency sensors may not be
612 needed for accurate mixed period water temperature forecasts or surface layer forecasts in the
613 stratified period.

614 The minimum frequency of DA needed to set up fully operational forecasting systems is
615 likely to vary based on the ecosystem or forecast variable of interest. Depending on the water
616 quality forecast application, different frequencies of data collection may be necessary to fully
617 understand and predict water quality dynamics over time. For example, George and Hurley
618 (2004) found that fortnightly observations were required to discern gradual trends in
619 phytoplankton productivity, but monthly data were adequate for capturing declines in

620 phytoplankton biomass over a 30-year period. Despite many successful applications of high-
621 frequency DA in the literature for forecasting (e.g., Cho et al., 2020, Gottwald and Reich, 2021,
622 Luo et al., 2011, Niu et al., 2014), not all ecological variables benefit from frequent DA, as not
623 all variables are similarly forecastable.

624 In addition to the frequency of data collection, data latency can also affect the frequency
625 of DA. Even for forecasting systems with high-frequency sensor data, data latency may reduce
626 forecast skill if data are not immediately transmitted to forecasting workflows (e.g., they require
627 a manual download) (Dietze et al., 2018). In cases with high data latency of the forecast variable
628 (e.g., microscope counts of phytoplankton requiring laboratory analysis), data fusion approaches
629 that assimilate multiple data sources may improve forecast skill (e.g., Baracchini et al., 2020b,
630 Chen et al., 2021). For example, some studies have assimilated both in-situ measurements and
631 remote sensing data to forecast reservoir water quality variables, including chlorophyll *a* and
632 conductivity (Abdul Wahid and Arunbabu, 2022, Chen et al., 2021).

633 Finally, understanding the contributions of different sources of uncertainty can be useful
634 for determining the DA frequency that generates the most skillful forecasts. Specifically,
635 knowing the relative contribution of initial conditions uncertainty can inform sampling frequency
636 needed to improve ecological forecast skill. For forecasts with total uncertainty dominated by
637 process, parameter, or driver uncertainty, improving forecast skill may require modifying
638 processes used for forecasting the ecological variable of interest, further constraining parameters
639 by collecting more data, or improving weather forecast driver data (e.g., Grönquist et al., 2021).

640

641

642

643 *Study Limitations*

644 Our results suggest that weekly DA may suffice for some lake and reservoir water
645 temperature forecasting applications, with the caveat that more frequent DA often improved
646 water temperature forecast performance at short forecast horizons. However, we only assessed
647 forecast skill for a single reservoir and ecological variable for only one year, and therefore note
648 the limitations of extending these results to other systems and variables. Additionally, updating
649 model parameters and initial conditions too regularly can lead to overprediction biases when
650 forecasting, which may explain why weekly rather than daily DA resulted in more skillful water
651 temperature forecasts in the mixed period (see Lin et al., 2021). Finally, because we did not
652 quantify the contribution of all sources of uncertainty, we can only identify the relative role that
653 DA has on reducing initial conditions uncertainty. Future studies that consider the role of other
654 sources of uncertainty will improve our understanding of DA on total forecast uncertainty.

655

656 *Conclusions*

657 This study emphasizes the importance of DA for improving ecological forecast skill and
658 has implications for forecasting efforts among a wide range of ecosystems and ecological
659 variables. We argue that weekly observations of water temperature are likely “good enough” to
660 set up a skillful forecasting system for many reservoir management applications, while daily DA
661 would be most useful for applications requiring high forecast accuracy in the bottom waters or at
662 short (< 5 - 7 day) forecast horizons. Because water temperature dynamics control many
663 biological, chemical, and physical lake processes (Magnuson et al., 1979, Read et al., 2019,
664 Yvon-Durocher et al., 2012), water temperature must be accurately forecasted before we can
665 forecast other water quality variables. Therefore, determining ways to improve water

666 temperature forecasts will have broad utility for advancing the development of many additional
667 water quality forecasting systems.

668 Because near-term, iterative forecasts are particularly well suited to address ecological
669 questions (Carey et al., 2022d, Dietze et al., 2018, White et al., 2019), determining how best to
670 design and deploy ecological near-term, iterative forecasting systems is a pressing need (Diez et
671 al., 2012, Ibáñez et al., 2013, Moustahfid et al., 2021). With the increasing deployment of high-
672 frequency sensor networks (e.g., National Ecological Observatory Network (NEON) and Global
673 Lake Ecological Observatory Network (GLEON); Mantovani et al., 2020, Marcé et al., 2016,
674 Park et al., 2020) comes a growing need to understand how best to use these sensor data for
675 forecasting. In response, we advocate for using DA experiments across ecosystems and
676 ecological variables to determine how best to integrate observational data into iterative
677 forecasting systems.

678

679 **Conflict of Interest Statement:** The authors declare no conflicts of interest.

680

681 **Data Availability Statement:** All data (Wander et al., 2023a) and code (Wander et al., 2023b)
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690

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692 developed the FLARE framework for data assimilation experiments and forecasting workflow
693 used in this study. TNM helped develop early iterations of the forecasting workflow. ABP
694 oversaw sensor data collection. HLW wrote the initial draft of the manuscript with CCC and
695 MEL; all coauthors reviewed the manuscript and approved its final version.

696

697 **References**

698 Abdul Wahid, A. and E. Arunbabu (2022) Forecasting water quality using seasonal ARIMA
699 model by integrating in-situ measurements and remote sensing techniques in Krishnagiri
700 reservoir, India. *Water Pract. Technol.*, **17**, 1230–52.

701 Baracchini, T., P. Y. Chu, J. Šukys, G. Lieberherr, S. Wunderle, A. Wüest, and D. Bouffard
702 (2020a) Data assimilation of in situ and satellite remote sensing data to 3D hydrodynamic
703 lake models: a case study using Delft3D-FLOW v4.03 and OpenDA v2.4. *Geosci. Model*
704 *Dev.*, **13**, 1267–84.

705 Baracchini, T., A. Wüest, and D. Bouffard (2020b) Meteolakes: An operational online three-
706 dimensional forecasting platform for lake hydrodynamics. *Water Res.*, **172**, 115529.

707 Bennett, N. D., B. F. Croke, G. Guariso, J. H. Guillaume, S. H. Hamilton, A. J. Jakeman, S.
708 Marsili-Libelli, L. T. Newham, *et al.* (2013) Characterising performance of
709 environmental models. *Environ. Model. Softw.*, **40**, 1–20.

710 Bett, P. E., K. E. Williams, C. Burton, A. A. Scaife, A. J. Wiltshire, and R. Gilham (2020)

711 Skillful seasonal prediction of key carbon cycle components: NPP and fire risk. *Environ.*
712 *Res. Commun.*, **2**, 055002.

713 Bruce, L. C., M. A. Frassl, G. B. Arhonditsis, G. Gal, D. P. Hamilton, P. C. Hanson, A. L.
714 Hetherington, J. M. Melack, *et al.* (2018) A multi-lake comparative analysis of the
715 General Lake Model (GLM): Stress-testing across a global observatory network. *Environ.*
716 *Model. Softw.*, **102**, 274–91.

717 Carey, C. C. and A. Breef-Pilz (2022) Ice cover data for Falling Creek Reservoir and Beaverdam
718 Reservoir, Vinton, Virginia, USA for 2013-2022. *Environmental Data Initiative*
719 *repository*. DOI: 10.6073/pasta/917b3947d91470eecf979e9297ed4d2e

720 Carey, C. C., A. Breef-Pilz, B. J. Bookout, R. P. McClure, and J. H. Wynne (2023) Time series
721 of high-frequency sensor data measuring water temperature, dissolved oxygen,
722 conductivity, specific conductance, total dissolved solids, chlorophyll a, phycocyanin,
723 fluorescent dissolved organic matter, and turbidity at discrete depths in Beaverdam
724 Reservoir, Virginia, USA in 2016-2022. *Environmental Data Initiative repository*. DOI:
725 10.6073/pasta/4182de376fde52e15d493fdd9f26d0c7

726 Carey, C. C., P. C. Hanson, R. Q. Thomas, A. B. Gerling, A. G. Hounshell, A. S. L. Lewis, M. E.
727 Lofton, R. P. McClure, *et al.* (2022a) Anoxia decreases the magnitude of the carbon,
728 nitrogen, and phosphorus sink in freshwaters. *Glob. Change Biol.*, **28**, 4861–81.

729 Carey, C. C., A. S. L. Lewis, D. W. Howard, W. M. Woelmer, P. A. Gantzer, K. A. Bierlein, J.
730 C. Little, and WVWA (2022b) Bathymetry and watershed area for Falling Creek
731 Reservoir, Beaverdam Reservoir, and Carvins Cove Reservoir. *Environmental Data*
732 *Initiative repository*. DOI: 10.6073/pasta/352735344150f7e77d2bc18b69a22412

733 Carey, C. C., A. S. Lewis, R. P. McClure, A. B. Gerling, A. Breef-Pilz, and A. Das (2022c) Time

734 series of high-frequency profiles of depth, temperature, dissolved oxygen, conductivity,
735 specific conductance, chlorophyll a, turbidity, pH, oxidation-reduction potential,
736 photosynthetic active radiation, and descent rate for Beaverdam Reservoir, Carvins Cove
737 Reservoir, Falling Creek Reservoir, Gatewood Reservoir, and Spring Hollow Reservoir
738 in Southwestern Virginia, USA 2013-2021. *Environmental Data Initiative repository*.
739 DOI: 10.6073/pasta/c4c45b5b10b4cb4cd4b5e613c3effbd0

740 Carey, C. C., W. M. Woelmer, M. E. Lofton, R. J. Figueiredo, B. J. Bookout, R. S. Corrigan, V.
741 Daneshmand, A. G. Hounshell, *et al.* (2022d) Advancing lake and reservoir water quality
742 management with near-term, iterative ecological forecasting. *Inland Waters*, **12**, 107–20.

743 Carpenter, S. R., E. H. Stanley, and M. J. Vander Zanden (2011) State of the world’s freshwater
744 ecosystems: physical, chemical, and biological changes. *Annu. Rev. Environ. Resour.*, **36**,
745 75–99.

746 Chen, C., Q. Chen, G. Li, M. He, J. Dong, H. Yan, Z. Wang, and Z. Duan (2021) A novel multi-
747 source data fusion method based on Bayesian inference for accurate estimation of
748 chlorophyll-a concentration over eutrophic lakes. *Environ. Model. Softw.*, **141**, 105057.

749 Cho, H. and H. Park (2019) Merged-LSTM and multistep prediction of daily chlorophyll-a
750 concentration for algal bloom forecast. *IOP Conf. Ser. Earth Environ. Sci.*, **351**, 012020.

751 Cho, K. H., Y. Pachepsky, M. Ligaray, Y. Kwon, and K. H. Kim (2020) Data assimilation in
752 surface water quality modeling: A review. *Water Res.*, **186**, 116307.

753 Clark, J. S., S. R. Carpenter, M. Barber, S. Collins, A. Dobson, J. A. Foley, D. M. Lodge, M.
754 Pascual, *et al.* (2001) Ecological forecasts: an emerging imperative. *science*, **293**, 657–
755 60.

756 Clark, P., N. Roberts, H. Lean, S. P. Ballard, and C. Charlton-Perez (2016) Convection-

757 permitting models: a step-change in rainfall forecasting. *Meteorol. Appl.*, **23**, 165–81.

758 Clayer, F., L. Jackson-Blake, D. Mercado-Bettín, M. Shikhani, A. French, T. Moore, J. Sample,
759 M. Norling, *et al.* (2023) Sources of skill in lake temperature, discharge and ice-off
760 seasonal forecasting tools. *Hydrol. Earth Syst. Sci.*, **27**, 1361–81.

761 Corbari, C., R. Salerno, A. Ceppi, V. Telesca, and M. Mancini (2019) Smart irrigation forecast
762 using satellite LANDSAT data and meteo-hydrological modeling. *Agric. Water Manag.*,
763 **212**, 283–94.

764 Daneshmand, V., A. Breef-Pilz, C. C. Carey, Y. Jin, Y.-J. Ku, K. C. Subratie, R. Q. Thomas, and
765 R. J. Figueiredo (2021) Edge-to-cloud Virtualized Cyberinfrastructure for Near Real-time
766 Water Quality Forecasting in Lakes and Reservoirs. *2021 IEEE 17th International*
767 *Conference on eScience (eScience)*. IEEE, pp. 138–48.

768 DeChant, C. M. and H. Moradkhani (2011) Improving the characterization of initial condition
769 for ensemble streamflow prediction using data assimilation. *Hydrol. Earth Syst. Sci.*, **15**,
770 3399–3410.

771 Derot, J., H. Yajima, and F. G. Schmitt (2020) Benefits of machine learning and sampling
772 frequency on phytoplankton bloom forecasts in coastal areas. *Ecol. Inform.*, **60**, 101174.

773 Dietze, M. C. (2017a) *Ecological forecasting*. Princeton University Press.

774 Dietze, M. C. (2017b) Prediction in ecology: a first-principles framework. *Ecol. Appl.*, **27**, 2048–
775 60.

776 Dietze, M. C., A. Fox, L. M. Beck-Johnson, J. L. Betancourt, M. B. Hooten, C. S. Jarnevich, T.
777 H. Keitt, M. A. Kenney, *et al.* (2018) Iterative near-term ecological forecasting: Needs,
778 opportunities, and challenges. *Proc. Natl. Acad. Sci.*, **115**, 1424–32.

779 Diez, J. M., I. Ibáñez, A. J. Miller-Rushing, S. J. Mazer, T. M. Crimmins, M. A. Crimmins, C. D.

780 Bertelsen, and D. W. Inouye (2012) Forecasting phenology: from species variability to
781 community patterns. *Ecol. Lett.*, **15**, 545–53.

782 Doubek, J. P., K. L. Campbell, M. E. Lofton, R. P. McClure, and C. C. Carey (2019)
783 Hypolimnetic hypoxia increases the biomass variability and compositional variability of
784 crustacean zooplankton communities. *Water*, **11**, 2179.

785 Duc, L., K. Saito, and D. Hotta (2021) Analysis and design of covariance inflation methods
786 using inflation functions. Part 2: adaptive inflation. *Q. J. R. Meteorol. Soc.*, **147**, 2375–
787 94.

788 Engelhardt, C. and G. Kirillin (2014) Criteria for the onset and breakup of summer lake
789 stratification based on routine temperature measurements. *Fundam. Appl. Limnol.*, **184**,
790 183–94.

791 Evensen, G. (2003) The Ensemble Kalman Filter: theoretical formulation and practical
792 implementation. *Ocean Dyn.*, **53**, 343–67.

793 Francy, D. S., J. L. Graham, E. A. Stelzer, C. D. Ecker, A. M. Brady, P. Struffolino, and K. A.
794 Loftin (2015) Water quality, cyanobacteria, and environmental factors and their relations
795 to microcystin concentrations for use in predictive models at ohio lake erie and inland
796 lake recreational sites, 2013-14. US Geological Survey Scientific Investigations Report
797 2015–5120, 58 p.

798 Georgakakos, K. P., N. E. Graham, T. M. Carpenter, and H. Yao (2005) Integrating climate-
799 hydrology forecasts and multi-objective reservoir management for northern California.
800 *Eos Trans. Am. Geophys. Union*, **86**, 122–27.

801 George, D. G. and M. A. Hurley (2004) The influence of sampling frequency on the detection of
802 long-term change in three lakes in the English Lake District. *Aquat. Ecosyst. Health*

803 *Manag.*, **7**, 1–14.

804 Gilarranz, L. J., A. Narwani, D. Odermatt, R. Siber, and V. Dakos (2022) Regime shifts, trends,
805 and variability of lake productivity at a global scale. *Proc. Natl. Acad. Sci.*, **119**,
806 e2116413119.

807 Gneiting, T., A. E. Raftery, A. H. Westveld, and T. Goldman (2005) Calibrated probabilistic
808 forecasting using ensemble model output statistics and minimum CRPS estimation. *Mon.*
809 *Weather Rev.*, **133**, 1098–1118.

810 Gottwald, G. A. and S. Reich (2021) Supervised learning from noisy observations: Combining
811 machine-learning techniques with data assimilation. *Phys. Nonlinear Phenom.*, **423**,
812 132911.

813 Grönquist, P., C. Yao, T. Ben-Nun, N. Dryden, P. Dueben, S. Li, and T. Hoefler (2021) Deep
814 learning for post-processing ensemble weather forecasts. *Philos. Trans. R. Soc. Math.*
815 *Phys. Eng. Sci.*, **379**, 20200092.

816 Hamre, K. D., M. E. Lofton, R. P. McClure, Z. W. Munger, J. P. Doubek, A. B. Gerling, M. E.
817 Schreiber, and C. C. Carey (2018) In situ fluorometry reveals a persistent, perennial
818 hypolimnetic cyanobacterial bloom in a seasonally anoxic reservoir. *Freshw. Sci.*, **37**,
819 483–95.

820 Harris, D. J., S. D. Taylor, and E. P. White (2018) Forecasting biodiversity in breeding birds
821 using best practices. *PeerJ*, **6**, e4278.

822 He, H., L. Lei, J. S. Whitaker, and Z.-M. Tan (2020) Impacts of Assimilation Frequency on
823 Ensemble Kalman Filter Data Assimilation and Imbalances. *J. Adv. Model. Earth Syst.*,
824 **12**, e2020MS002187.

825 Heilman, K. A., M. C. Dietze, A. A. Arizpe, J. Aragon, A. Gray, J. D. Shaw, A. O. Finley, S.

826 Klesse, *et al.* (2022) Ecological forecasting of tree growth: Regional fusion of tree-ring
827 and forest inventory data to quantify drivers and characterize uncertainty. *Glob. Change*
828 *Biol.*, **28**, 2442–60.

829 Herrick, C., B. G. Steele, J. A. Brentrup, K. L. Cottingham, M. J. Ducey, D. A. Lutz, M. W.
830 Palace, M. C. Thompson, *et al.* (2023) lakeCoSTR: A tool to facilitate use of Landsat
831 Collection 2 to estimate lake surface water temperatures. *Ecosphere*, **14**, e4357.

832 Hipsey, M. R., C. Boon, L. C. Bruce, Q. Thomas, M. Weber, L. Winslow, J. S. Read, and D. P.
833 Hamilton (2022) AquaticEcoDynamics/glm-aed: v3.3.0.

834 Hipsey, M. R., L. C. Bruce, C. Boon, B. Busch, C. C. Carey, D. P. Hamilton, P. C. Hanson, J. S.
835 Read, *et al.* (2019) A General Lake Model (GLM 3.0) for linking with high-frequency
836 sensor data from the Global Lake Ecological Observatory Network (GLEON). *Geosci.*
837 *Model Dev.*, **12**, 473–523.

838 Hounshell, A. G., R. P. McClure, M. E. Lofton, and C. C. Carey (2021) Whole-ecosystem
839 oxygenation experiments reveal substantially greater hypolimnetic methane
840 concentrations in reservoirs during anoxia. *Limnol. Oceanogr. Lett.*, **6**, 33–42.

841 Ibáñez, I., E. S. Gornish, L. Buckley, D. M. Debinski, J. Hellmann, B. Helmuth, J.
842 HilleRisLambers, A. M. Latimer, *et al.* (2013) Moving forward in global-change ecology:
843 capitalizing on natural variability. *Ecol. Evol.*, **3**, 170–81.

844 Jolliffe, I. T. and D. B. Stephenson (2012) *Forecast verification: a practitioner's guide in*
845 *atmospheric science*. John Wiley & Sons.

846 Kehoe, M. J., K. P. Chun, and H. M. Baulch (2015) Who Smells? Forecasting Taste and Odor in
847 a Drinking Water Reservoir. *Environ. Sci. Technol.*, **49**, 10984–92.

848 Kirchner, J. W. and C. Neal (2013) Universal fractal scaling in stream chemistry and its

849 implications for solute transport and water quality trend detection. *Proc. Natl. Acad. Sci.*,
850 **110**, 12213–18.

851 LaDeau, S. L., B. A. Han, E. J. Rosi-Marshall, and K. C. Weathers (2017) The Next Decade of
852 Big Data in Ecosystem Science. *Ecosystems*, **20**, 274–83.

853 Ladwig, R., L. A. Rock, and H. A. Dugan (2021) Impact of salinization on lake stratification and
854 spring mixing. *Limnol. Oceanogr. Lett.*, **8**, 93-102.

855 Lewis, A. S. L., W. M. Woelmer, H. L. Wander, D. W. Howard, J. W. Smith, R. P. McClure, M.
856 E. Lofton, N. W. Hammond, *et al.* (2022) Increased adoption of best practices in
857 ecological forecasting enables comparisons of forecastability. *Ecol. Appl.*, **32**, e2500.

858 Lin, E., Y. Yang, X. Qiu, Q. Xie, R. Gan, B. Zhang, and X. Liu (2021) Impacts of the radar data
859 assimilation frequency and large-scale constraint on the short-term precipitation forecast
860 of a severe convection case. *Atmospheric Res.*, **257**, 105590.

861 Lindegren, M., C. Möllmann, A. Nielsen, K. Brander, B. R. MacKenzie, and N. Chr. Stenseth
862 (2010) Ecological forecasting under climate change: the case of Baltic cod. *Proc. R. Soc.*
863 *B Biol. Sci.*, **277**, 2121–30.

864 Liu, H., Y. D. Chen, T. Liu, and L. Lin (2019) The river chief system and river pollution control
865 in China: A case study of Foshan. *Water*, **11**, 1606.

866 Lofton, M. E., J. A. Brentrup, W. S. Beck, J. A. Zwart, R. Bhattacharya, L. S. Brighenti, S. H.
867 Burnet, I. M. McCullough, *et al.* (2022) Using near-term forecasts and uncertainty
868 partitioning to inform prediction of oligotrophic lake cyanobacterial density. *Ecol. Appl.*,
869 **32**, e2590.

870 Lofton, M. E., D. W. Howard, R. Q. Thomas, and C. C. Carey (2023) Progress and opportunities
871 in advancing near-term forecasting of freshwater quality. *Glob. Change Biol.*, **29**, 1691–

872 1714.

873 Luo, Y., K. Ogle, C. Tucker, S. Fei, C. Gao, S. LaDeau, J. S. Clark, and D. S. Schimel (2011)

874 Ecological forecasting and data assimilation in a data-rich era. *Ecol. Appl.*, **21**, 1429–42.

875 Machete, R. L. and L. A. Smith (2016) Demonstrating the value of larger ensembles in

876 forecasting physical systems. *Tellus Dyn. Meteorol. Oceanogr.*, **68**, 28393.

877 Magnuson, J. J., L. B. Crowder, and P. A. Medvick (1979) Temperature as an ecological

878 resource. *Am. Zool.*, **19**, 331–43.

879 Malhi, Y., J. Franklin, N. Seddon, M. Solan, M. G. Turner, C. B. Field, and N. Knowlton (2020)

880 Climate change and ecosystems: threats, opportunities and solutions. *Philos. Trans. R.*

881 *Soc. B Biol. Sci.*, **375**, 20190104.

882 Mantovani, C., L. Corgnati, J. Horstmann, A. Rubio, E. Reyes, C. Quentin, S. Cosoli, J. L.

883 Asensio, *et al.* (2020) Best Practices on High Frequency Radar Deployment and

884 Operation for Ocean Current Measurement. *Front. Mar. Sci.*, **7**, 210.

885 Marcé, R., G. George, P. Buscarinu, M. Deidda, J. Dunalska, E. de Eyto, G. Flaim, H.-P.

886 Grossart, *et al.* (2016) Automatic High Frequency Monitoring for Improved Lake and

887 Reservoir Management. *Environ. Sci. Technol.*, **50**, 10780–94.

888 Marj, A. F. and A. M. J. Meijerink (2011) Agricultural drought forecasting using satellite

889 images, climate indices and artificial neural network. *Int. J. Remote Sens.*, **32**, 9707–19.

890 Massoud, E. C., J. Huisman, E. Benincà, M. C. Dietze, W. Bouten, and J. A. Vrugt (2018)

891 Probing the limits of predictability: data assimilation of chaotic dynamics in complex

892 food webs. *Ecol. Lett.*, **21**, 93–103.

893 McClure, R. P., R. Q. Thomas, M. E. Lofton, W. M. Woelmer, and C. C. Carey (2021) Iterative

894 Forecasting Improves Near-Term Predictions of Methane Ebullition Rates. *Front.*

895 *Environ. Sci.*, **9**, 756603.

896 Mercado-Bettín, D., F. Clayer, M. Shikhani, T. N. Moore, M. D. Frías, L. Jackson-Blake, J.

897 Sample, M. Iturbide, *et al.* (2021) Forecasting water temperature in lakes and reservoirs

898 using seasonal climate prediction. *Water Res.*, **201**, 117286.

899 Meyer, J. L., M. J. Sale, P. J. Mulholland, and N. L. Poff (1999) Impacts of climate change on

900 aquatic ecosystem functioning and health 1. *JAWRA J. Am. Water Resour. Assoc.*, **35**,

901 1373–86.

902 Mi, C., T. Shatwell, J. Ma, Y. Xu, F. Su, and K. Rinke (2020) Ensemble warming projections in

903 Germany’s largest drinking water reservoir and potential adaptation strategies. *Sci. Total*

904 *Environ.*, **748**, 141366.

905 Moustahfid, H., L. C. Hendrickson, A. Arkhipkin, G. J. Pierce, A. Gangopadhyay, H. Kidokoro,

906 U. Markaida, C. Nigmatullin, *et al.* (2021) Ecological-fishery forecasting of squid stock

907 dynamics under climate variability and change: review, Challenges, and

908 Recommendations. *Rev. Fish. Sci. Aquac.*, **29**, 682–705.

909 Niu, S., Y. Luo, M. C. Dietze, T. F. Keenan, Z. Shi, J. Li, and F. S. C. Iii (2014) The role of data

910 assimilation in predictive ecology. *Ecosphere*, **5**, 1-16.

911 O’Reilly, C. M., S. Sharma, D. K. Gray, S. E. Hampton, J. S. Read, R. J. Rowley, P. Schneider,

912 J. D. Lenters, *et al.* (2015) Rapid and highly variable warming of lake surface waters

913 around the globe. *Geophys. Res. Lett.*, **42**, 10–773.

914 Paerl, H. W. and V. J. Paul (2012) Climate change: links to global expansion of harmful

915 cyanobacteria. *Water Res.*, **46**, 1349–63.

916 Page, T., P. J. Smith, K. J. Beven, I. D. Jones, J. A. Elliott, S. C. Maberly, E. B. Mackay, M. De

917 Ville, *et al.* (2018) Adaptive forecasting of phytoplankton communities. *Water Res.*, **134**,

918 74–85.

919 Page, T., P. J. Smith, K. J. Beven, I. D. Jones, J. A. Elliott, S. C. Maberly, E. B. Mackay, M. De
920 Ville, *et al.* (2017) Constraining uncertainty and process-representation in an algal
921 community lake model using high frequency in-lake observations. *Ecol. Model.*, **357**, 1–
922 13.

923 Park, J., K. T. Kim, and W. H. Lee (2020) Recent Advances in Information and Communications
924 Technology (ICT) and Sensor Technology for Monitoring Water Quality. *Water*, **12**, 510.

925 Parrish, M. A., H. Moradkhani, and C. M. DeChant (2012) Toward reduction of model
926 uncertainty: Integration of Bayesian model averaging and data assimilation. *Water*
927 *Resour. Res.*, **48**, W03519.

928 Piazzzi, G., G. Thirel, L. Campo, and S. Gabellani (2018) A particle filter scheme for multivariate
929 data assimilation into a point-scale snowpack model in an Alpine environment. *The*
930 *Cryosphere*, **12**, 2287–2306.

931 R Core Team (2022) R: A language and environment for statistical computing.

932 Read, J. S., X. Jia, J. Willard, A. P. Appling, J. A. Zwart, S. K. Oliver, A. Karpatne, G. J. A.
933 Hansen, *et al.* (2019) Process-Guided Deep Learning Predictions of Lake Water
934 Temperature. *Water Resour. Res.*, **55**, 9173–90.

935 Read, J. S., L. A. Winslow, G. J. A. Hansen, J. Van Den Hoek, P. C. Hanson, L. C. Bruce, and C.
936 D. Markfort (2014) Simulating 2368 temperate lakes reveals weak coherence in
937 stratification phenology. *Ecol. Model.*, **291**, 142–50.

938 Romero, J. R., I. Kagalou, J. Imberger, D. Hela, M. Kotti, A. Bartzokas, T. Albanis, N.
939 Evmirides, *et al.* (2002) Seasonal water quality of shallow and eutrophic Lake Pamvotis,
940 Greece: implications for restoration. *Hydrobiologia*, **474**, 91-105

941 Rousso, B. Z., E. Bertone, R. Stewart, and D. P. Hamilton (2020) A systematic literature review
942 of forecasting and predictive models for cyanobacteria blooms in freshwater lakes. *Water*
943 *Res.*, **182**, 115959.

944 Simonin, D., C. Pierce, N. Roberts, S. P. Ballard, and Z. Li (2017) Performance of Met Office
945 hourly cycling NWP-based nowcasting for precipitation forecasts. *Q. J. R. Meteorol.*
946 *Soc.*, **143**, 2862–73.

947 Steere, D. C., A. Baptista, D. McNamee, C. Pu, and J. Walpole (2000) Research challenges in
948 environmental observation and forecasting systems. *Proceedings of the 6th annual*
949 *international conference on Mobile computing and networking - MobiCom '00*. ACM
950 Press, Boston, Massachusetts, United States, pp. 292–99.

951 Tanut, B., R. Waranusast, and P. Riyamongkol (2021) High accuracy pre-harvest sugarcane yield
952 forecasting model utilizing drone image analysis, data mining, and reverse design
953 method. *Agriculture*, **11**, 682.

954 Thomas, R. Q., C. Boettiger, C. C. Carey, M. C. Dietze, L. R. Johnson, M. A. Kenney, J. S.
955 McLachlan, J. A. Peters, *et al.* (2023a) The NEON Ecological Forecasting Challenge.
956 *Front. Ecol. Environ.*, **21**, 112–13.

957 Thomas, R. Q., R. J. Figueiredo, V. Daneshmand, B. J. Bookout, L. K. Puckett, and C. C. Carey
958 (2020) A Near-Term Iterative Forecasting System Successfully Predicts Reservoir
959 Hydrodynamics and Partitions Uncertainty in Real Time. *Water Resour. Res.*, **56**,
960 e2019WR026138.

961 Thomas, R. Q., R. P. McClure, T. N. Moore, W. M. Woelmer, C. Boettiger, R. J. Figueiredo, R.
962 T. Hensley, and C. C. Carey (2023b) Near-term forecasts of NEON lakes reveal gradients
963 of environmental predictability across the US. *Front. Ecol. Environ.*, **n/a**.

964 Wander, H. L., R. Q. Thomas, T. N. Moore, M. E. Lofton, A. Breef-Pilz, and C. C. Carey
965 (2023a) Data assimilation experiments inform monitoring needs for near-term ecological
966 forecasts in a eutrophic reservoir: data, forecasts, and scores [Data set]. *Zenodo*.
967 <https://doi.org/10.5281/zenodo.7951402>

968 Wander, H. L., R. Q. Thomas, T. N. Moore, M. E. Lofton, A. Breef-Pilz, and C. C. Carey
969 (2023b) hlwander/BVRE-forecast-code: Data assimilation experiments inform
970 monitoring needs for near-term ecological forecasts in a eutrophic reservoir: Code (v1.1).
971 *Zenodo*. <https://doi.org/10.5281/zenodo.7958471>

972 Wang, S., N. Flipo, and T. Romary (2023) Which filter for data assimilation in water quality
973 models? Focus on oxygen reaeration and heterotrophic bacteria activity. *J. Hydrol.*, **620**,
974 129423.

975 Wang, X., J. Zhang, and V. Babovic (2016) Improving real-time forecasting of water quality
976 indicators with combination of process-based models and data assimilation technique.
977 *Ecol. Indic.*, **66**, 428–39.

978 Weng, E. and Y. Luo (2011) Relative information contributions of model vs. data to short- and
979 long-term forecasts of forest carbon dynamics. *Ecol. Appl.*, **21**, 1490–1505.

980 Wheeler, K., M. Dietze, D. LeBauer, J. Peters, A. D. Richardson, R. Q. Thomas, K. Zhu, U.
981 Bhat, *et al.* (2023) Predicting Spring Phenology in Deciduous Broadleaf Forests: An
982 Open Community Forecast Challenge. Available at SSRN: 10.2139/ssrn.4357147

983 White, E. P., G. M. Yenni, S. D. Taylor, E. M. Christensen, E. K. Bledsoe, J. L. Simonis, and S.
984 K. M. Ernest (2019) Developing an automated iterative near-term forecasting system for
985 an ecological study. *Methods Ecol. Evol.*, **10**, 332–44.

986 Williamson, C. E., E. P. Overholt, J. A. Brentrup, R. M. Pilla, T. H. Leach, S. G. Schladow, J. D.

987 Warren, S. S. Urmy, *et al.* (2016) Sentinel responses to droughts, wildfires, and floods:
988 effects of UV radiation on lakes and their ecosystem services. *Front. Ecol. Environ.*, **14**,
989 102–9.

990 Woelmer, W. M., R. Q. Thomas, M. E. Lofton, R. P. McClure, H. L. Wander, and C. C. Carey
991 (2022) Near-term phytoplankton forecasts reveal the effects of model time step and
992 forecast horizon on predictability. *Ecol. Appl.*, **32**, e2642.

993 Woolway, R. I., E. Jennings, T. Shatwell, M. Golub, D. C. Pierson, and S. C. Maberly (2021)
994 Lake heatwaves under climate change. *Nature*, **589**, 402–7.

995 Yvon-Durocher, G., J. M. Caffrey, A. Cescatti, M. Dossena, P. del Giorgio, J. M. Gasol, J. M.
996 Montoya, J. Pumpanen, *et al.* (2012) Reconciling the temperature dependence of
997 respiration across timescales and ecosystem types. *Nature*, **487**, 472–76.

998 Ziliani, M. G., R. Ghostine, B. Ait-El-Fquih, M. F. McCabe, and I. Hoteit (2019) Enhanced flood
999 forecasting through ensemble data assimilation and joint state-parameter estimation. *J.*
1000 *Hydrol.*, **577**, 123924.

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1004 **Figure Captions**

1005 **Figure 1:** Forecasting Lake And Reservoir Ecosystems (FLARE) workflow showing the step-by-
1006 step process for generating daily water temperature forecasts, starting with data collection from
1007 thermistors deployed in the reservoir (step 1), then data access for running the forecast model
1008 (step 2), then generation of forecasts with data assimilation (step 3), and ending with forecast
1009 skill assessment (step 4). During the data assimilation steps (3a-b), data assimilation experiments
1010 were performed with four different data assimilation frequencies (daily, weekly, fortnightly, and
1011 monthly; see dashed line box). Steps 1-4 occurred throughout the entire forecast period (1
1012 January - 31 December 2021). Buoy figure via NexSens Technology Inc., CC by 2.0
1013 <https://creativecommons.org/licenses/by/2.0/>

1014
1015 **Figure 2:** Map of Beaverdam Reservoir, Vinton, VA (37.31° N, 79.82° W). The map shows the
1016 surrounding forested watershed; the point represents the reservoir monitoring site where high-
1017 frequency sensor data were collected.

1018
1019 **Figure 3:** Observed water temperature for all depths with high-frequency sensors during the
1020 forecasting period of 1 January - 31 December 2021 in Beaverdam Reservoir (BVR). The gray
1021 background indicates the mixed period (1 January - 11 March, 8 November - 31 December
1022 2021), while the white background indicates the thermally-stratified period (12 March - 7
1023 November 2021), defined by a $<0.1 \text{ kg/m}^3$ density differential between surface and bottom
1024 layers.

1025

1026 **Figure 4:** Example of water temperature forecasts at 1 m (a), 5 m (b), and 9 m (c) generated for
1027 1-35 days into the future in Beaverdam Reservoir. Data assimilation (DA) frequencies are
1028 depicted by colors; shading shows 95% confidence intervals around the mean predicted
1029 temperature for each day. Black points represent water temperature observations. Colored points
1030 represent the most recent day that data was assimilated for each DA frequency. In this example,
1031 data were most recently assimilated on the day that the forecasts were generated: 25 June for the
1032 monthly DA scenario, 9 July for the fortnightly DA scenario, 16 July for the weekly DA
1033 scenario, and 22 July for the daily DA scenario.

1034
1035 **Figure 5:** Parameter evolution during the forecast period (1 January - 31 December 2021) for
1036 daily, weekly, fortnightly, and monthly data assimilation (DA) frequencies at 1-day-ahead
1037 forecast horizons. Longwave (a) is the longwave radiation scaling parameter, hypo_sed_temp (b)
1038 is the hypolimnetic sediment temperature parameter, and epi_sed_temp (c) is the epilimnetic
1039 sediment temperature parameter.

1040
1041 **Figure 6:** Root mean square error (RMSE) of mean forecasted water temperature compared to
1042 observations for 1-35 day-ahead forecast horizons in Beaverdam Reservoir, aggregated for all
1043 depths in the water column and days within the 365-day forecast period. RMSE for each forecast
1044 horizon was averaged from forecasts generated during 1 January - 31 December 2021. Colored
1045 lines represent different data assimilation (DA) frequencies. The dotted line depicts the 2°C
1046 threshold for skillful water temperature forecasts.

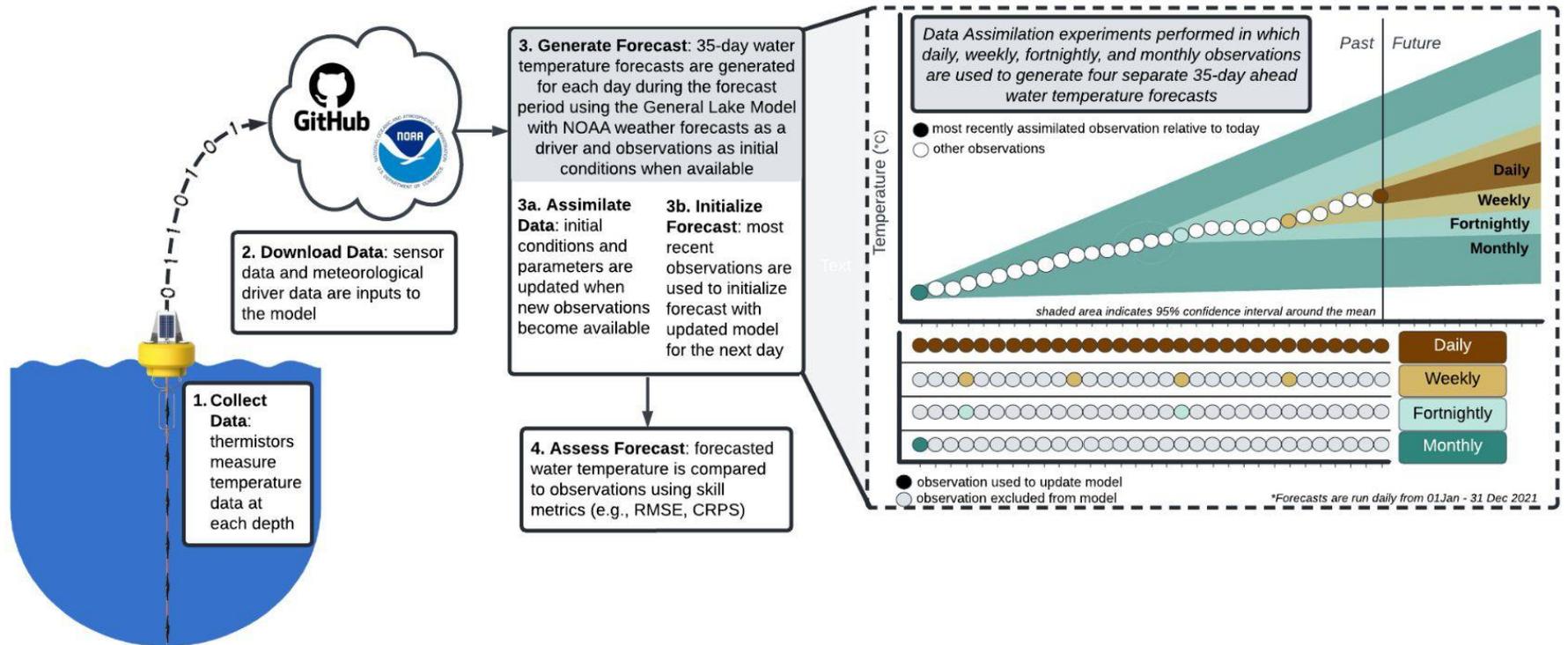
1047

1048 **Figure 7:** Root mean square error (RMSE) of mean forecasted water temperature compared to
1049 observations for 1-35-day-ahead forecast horizons in Beaverdam Reservoir during the mixed
1050 (panels a, c, e) vs. stratified (panels b, d, f) periods at 1 m (a, b), 5 m (c, d), and 9 m (e, f). RMSE
1051 for each forecast horizon was averaged across the 365-day forecast period (1 January - 31
1052 December 2021). Colored lines correspond to different data assimilation (DA) frequencies
1053 (daily, weekly, fortnightly, and monthly); dotted horizontal lines depict the 2°C threshold for
1054 skillful forecasts.

1055
1056 **Figure 8:** Mean water temperature forecast variance across horizons (1-35 days ahead) in
1057 Beaverdam Reservoir during the mixed (panels a, c, e) vs. stratified (panels b, d, f) periods for 1
1058 m (a, b), 5 m (c, d), and 9 m (e, f). Variance for each forecast horizon was averaged from all 365
1059 forecasts generated during the forecast period (1 January - 31 December 2021). Colored lines
1060 correspond to different data assimilation (DA) frequencies (daily, weekly, fortnightly, and
1061 monthly).

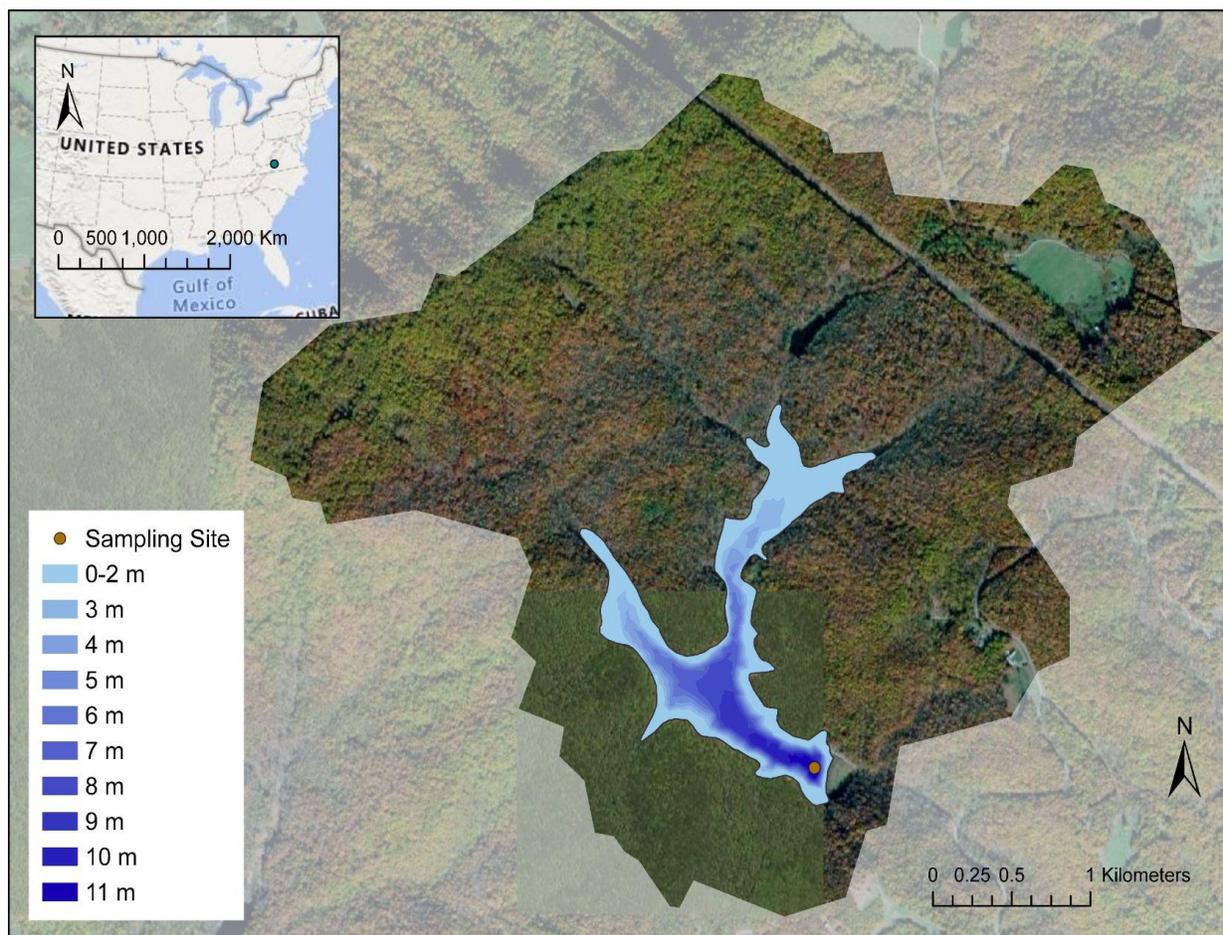
1062
1063 **Figure 9:** Proportion of initial conditions uncertainty relative to total forecast uncertainty
1064 averaged across all forecasts generated with each DA frequency calculated using forecast
1065 variance across 1 January - 31 December 2021. Colored lines depict data assimilation (DA)
1066 frequencies (daily, weekly, fortnightly, and monthly). Panels a, c, and e represent mixed period
1067 forecasts, panels b, d, and f represent stratified period forecasts. Depths (1, 5, and 9 m) are
1068 indicated by gray facet labels.

1069 Figure 1.
1070



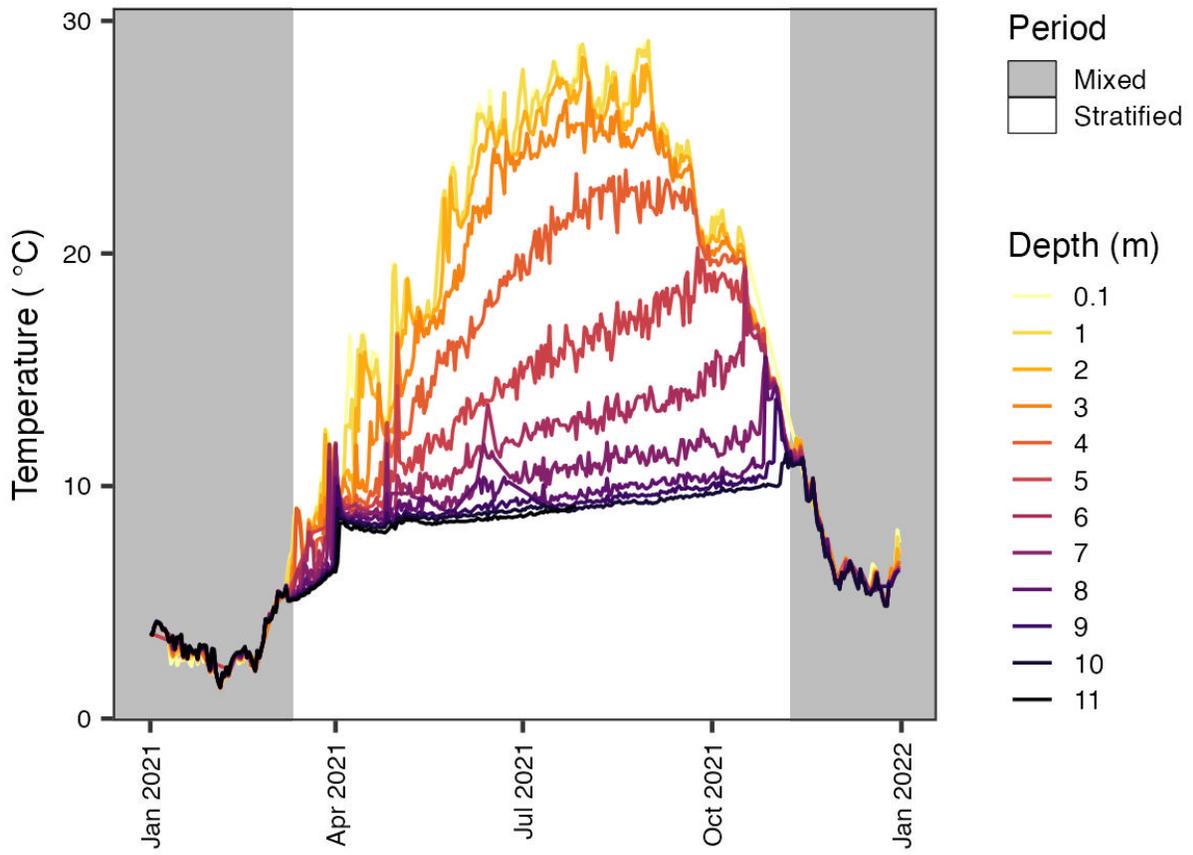
1071

1072 Figure 2.



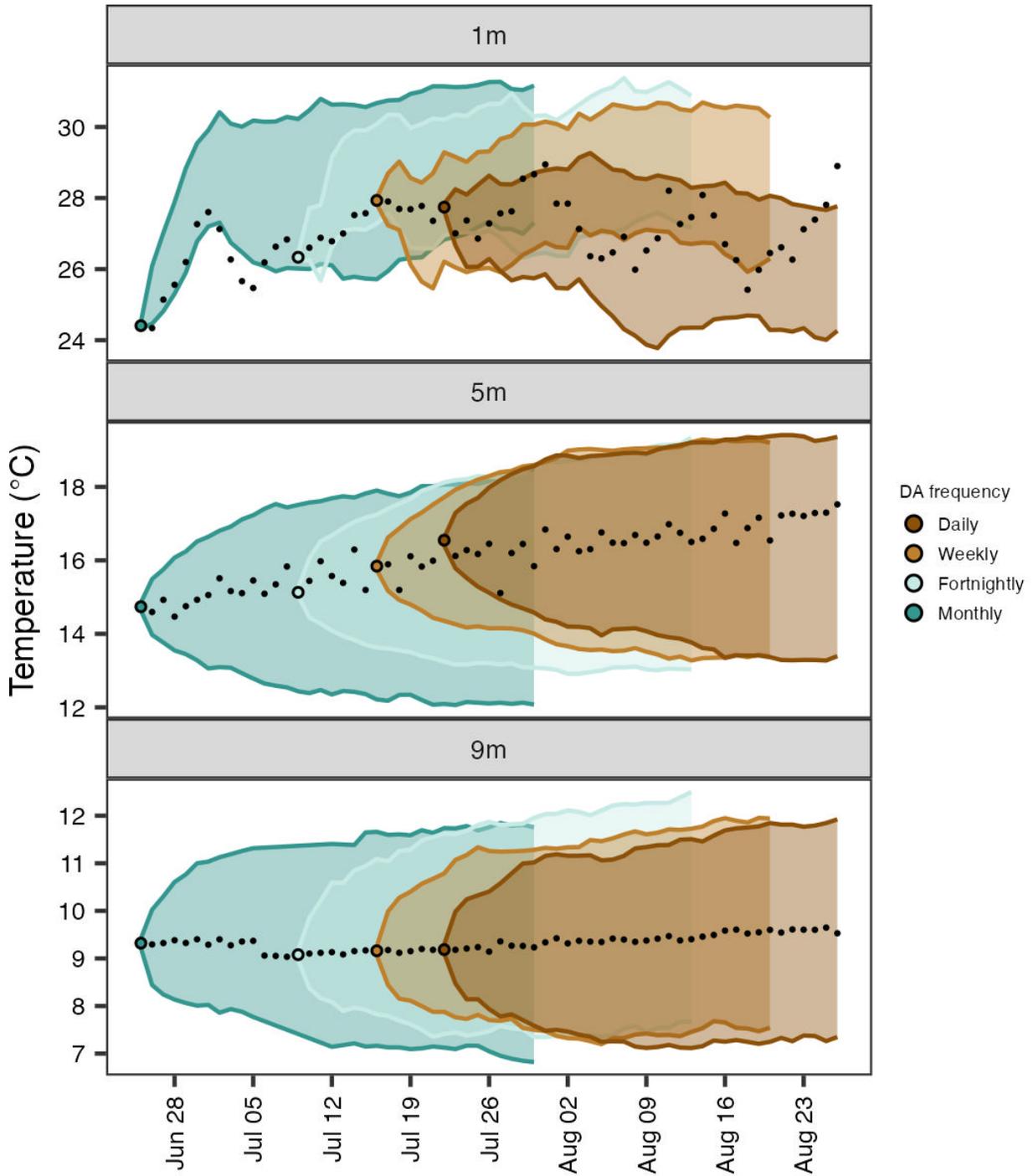
1073

1074 Figure 3.



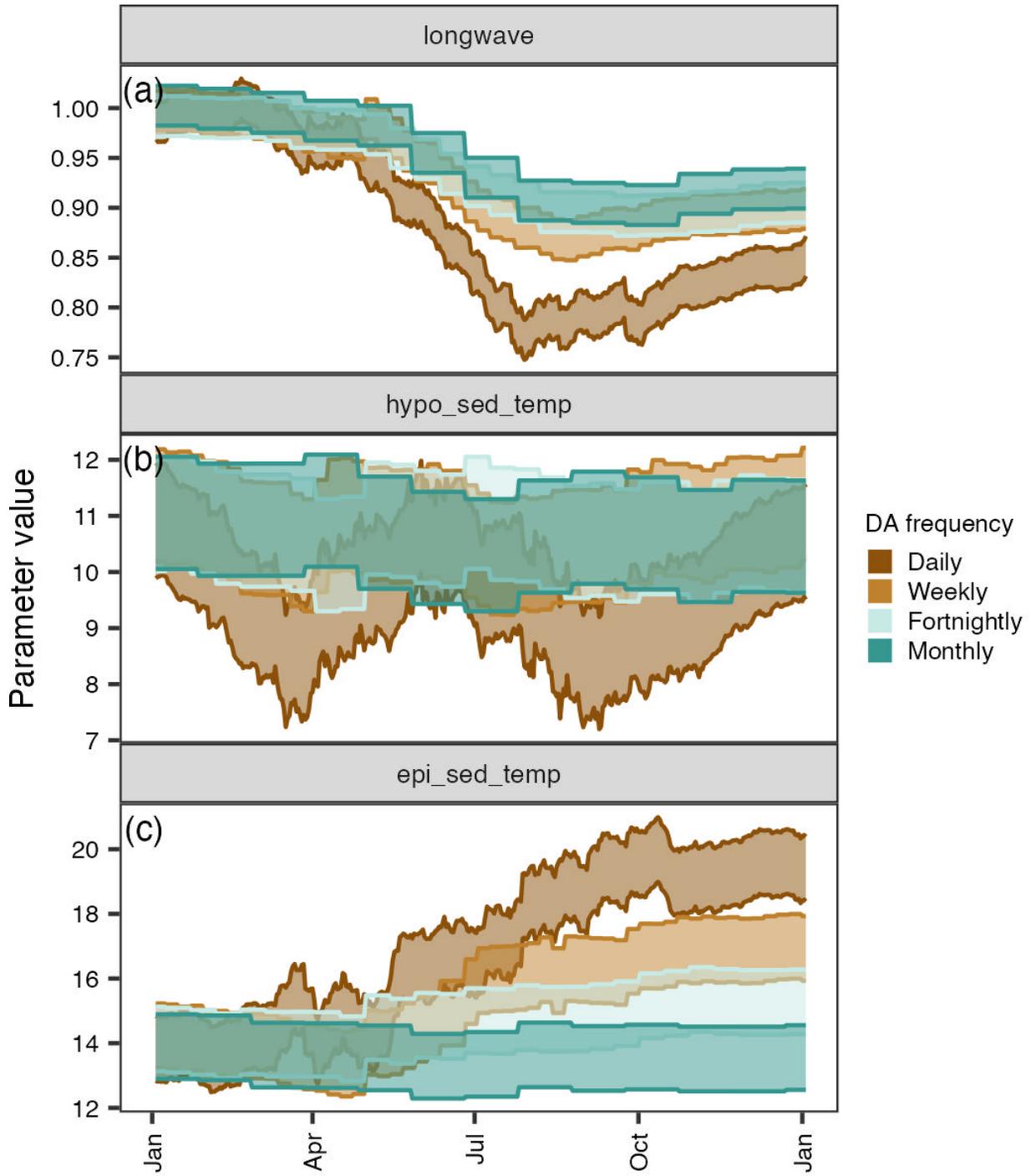
1075

1076 Figure 4.



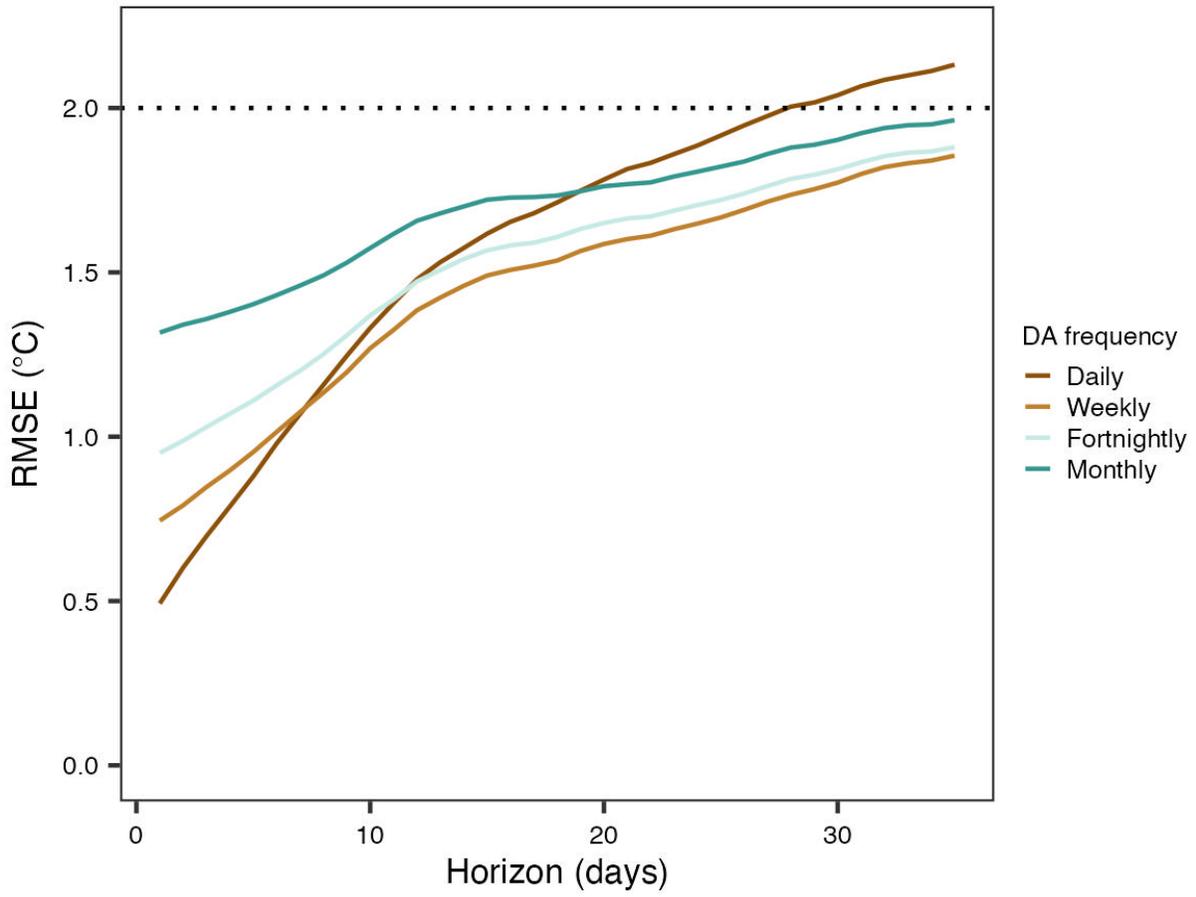
1077

1078 Figure 5.



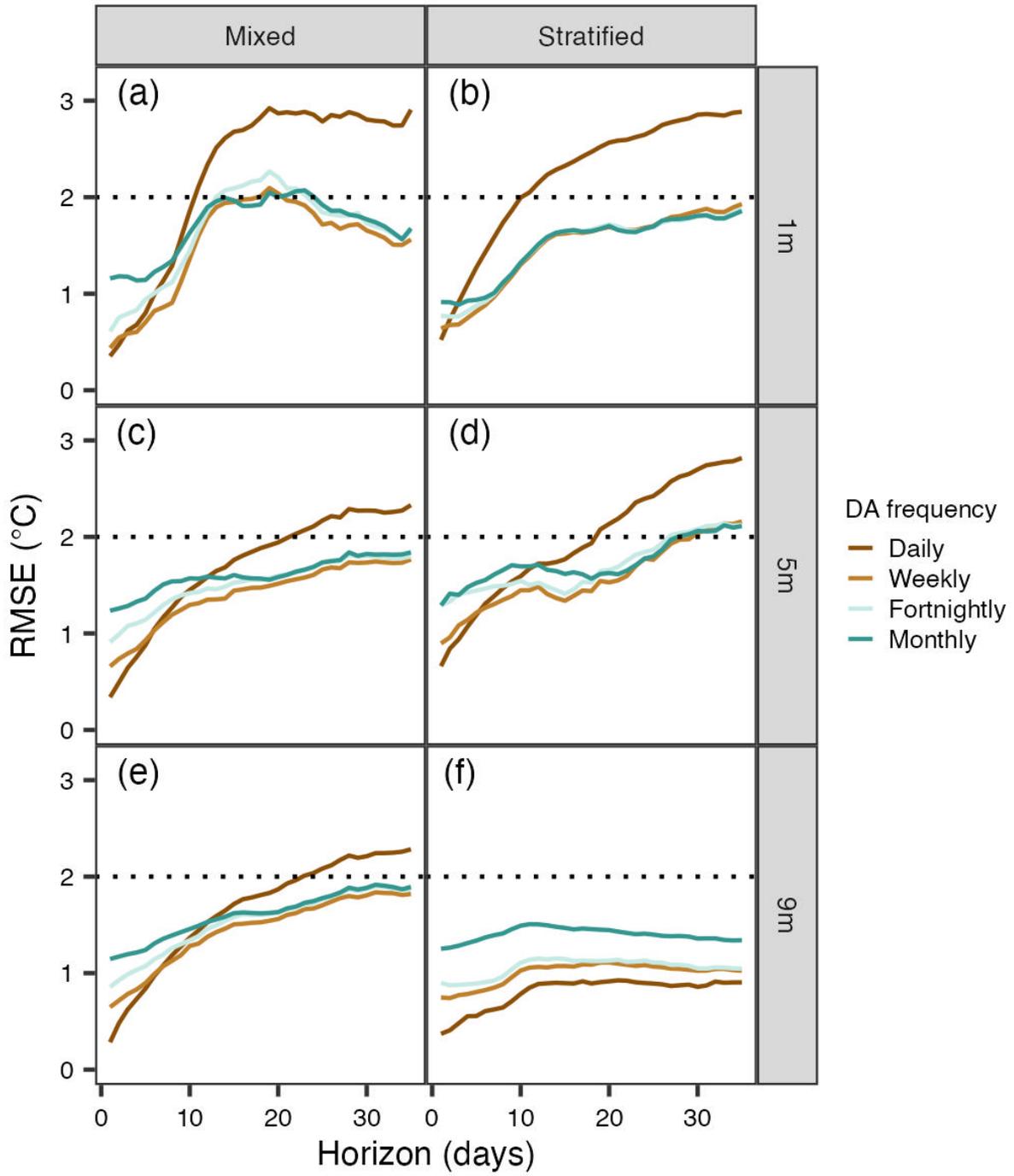
1079

1080 Figure 6.



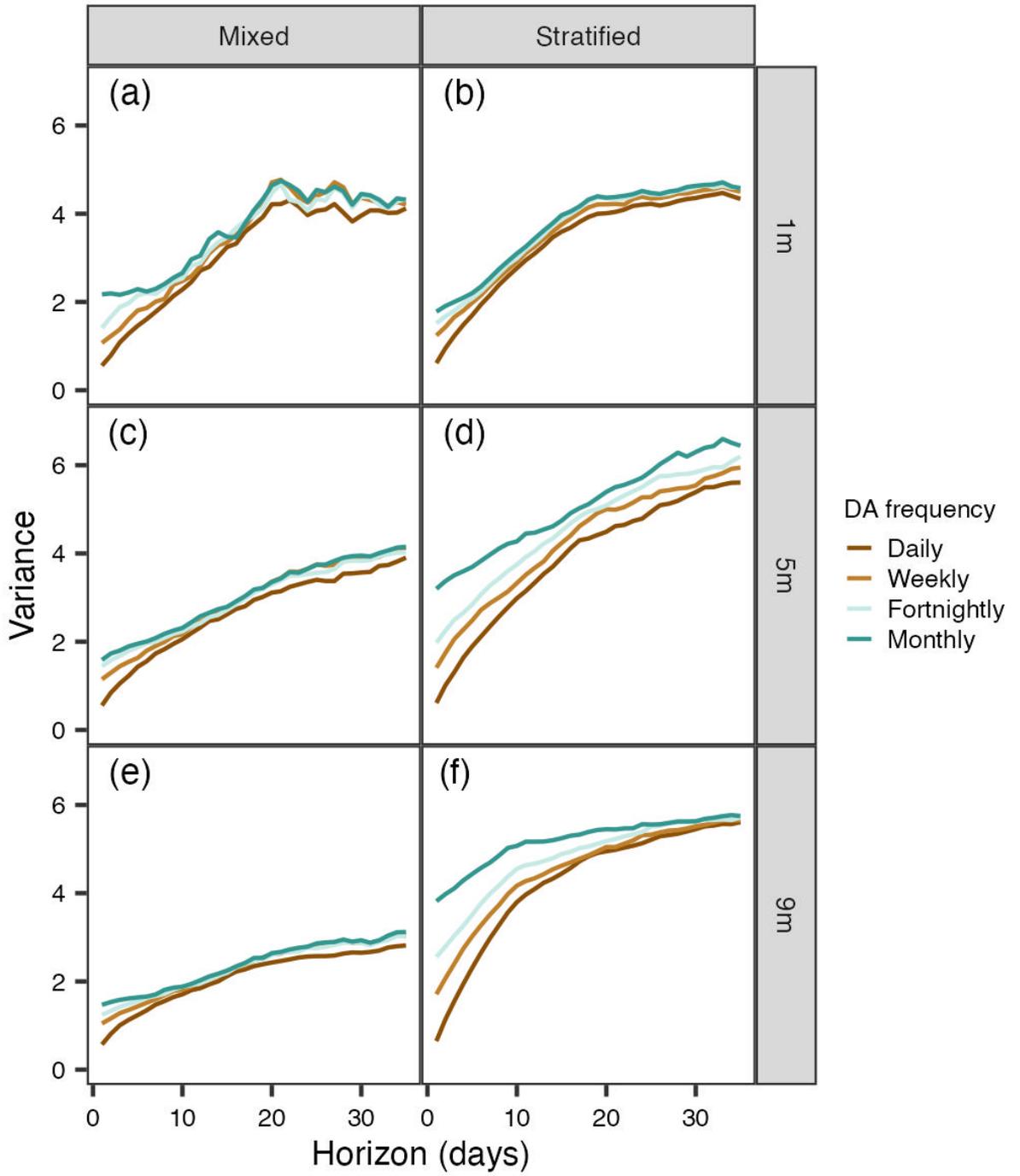
1081

1082 Figure 7.



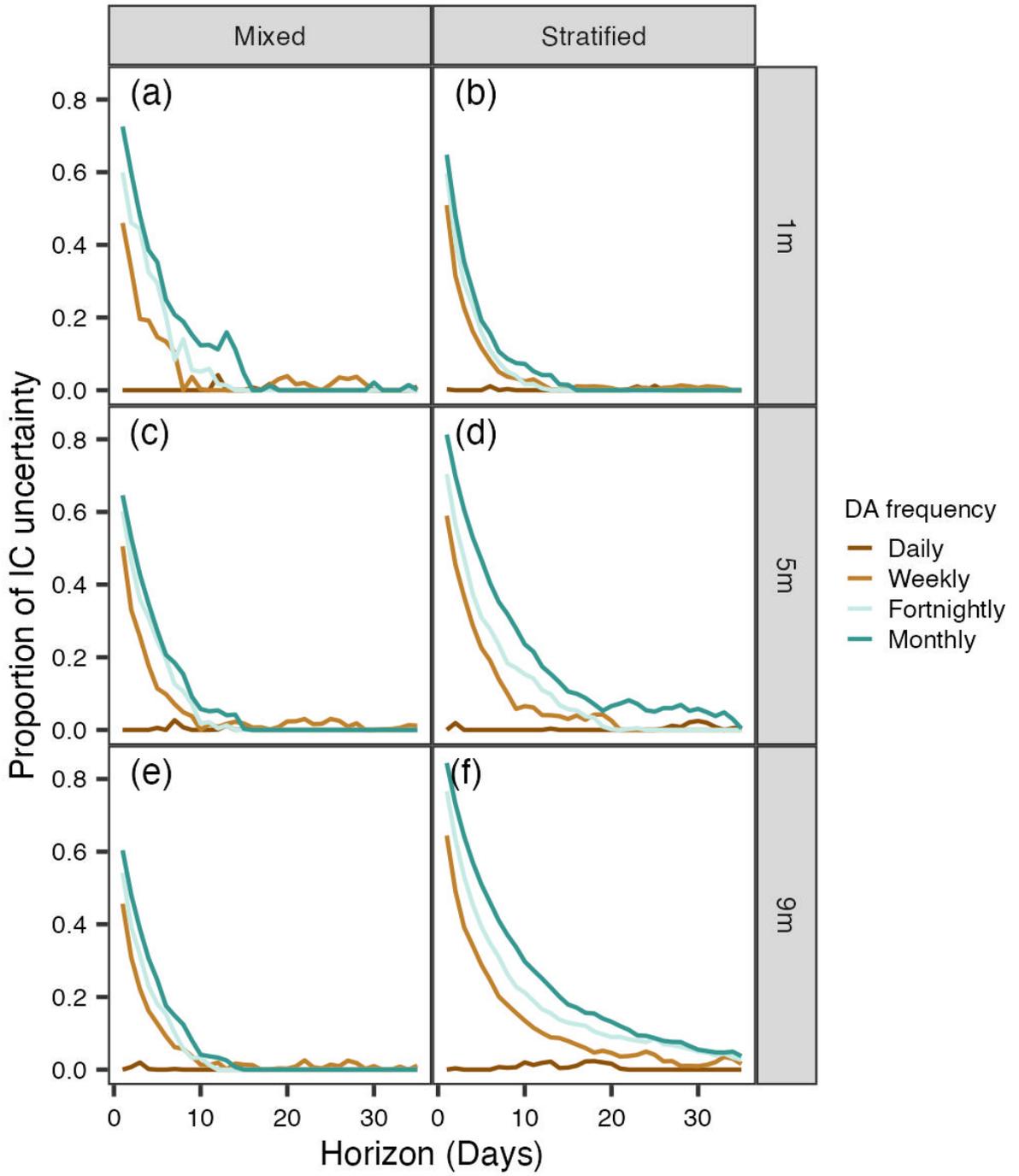
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1084 Figure 8.



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1086 Figure 9.



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Supplement to: Data assimilation experiments inform monitoring needs for near-term ecological forecasts in a eutrophic reservoir

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Appendix S1

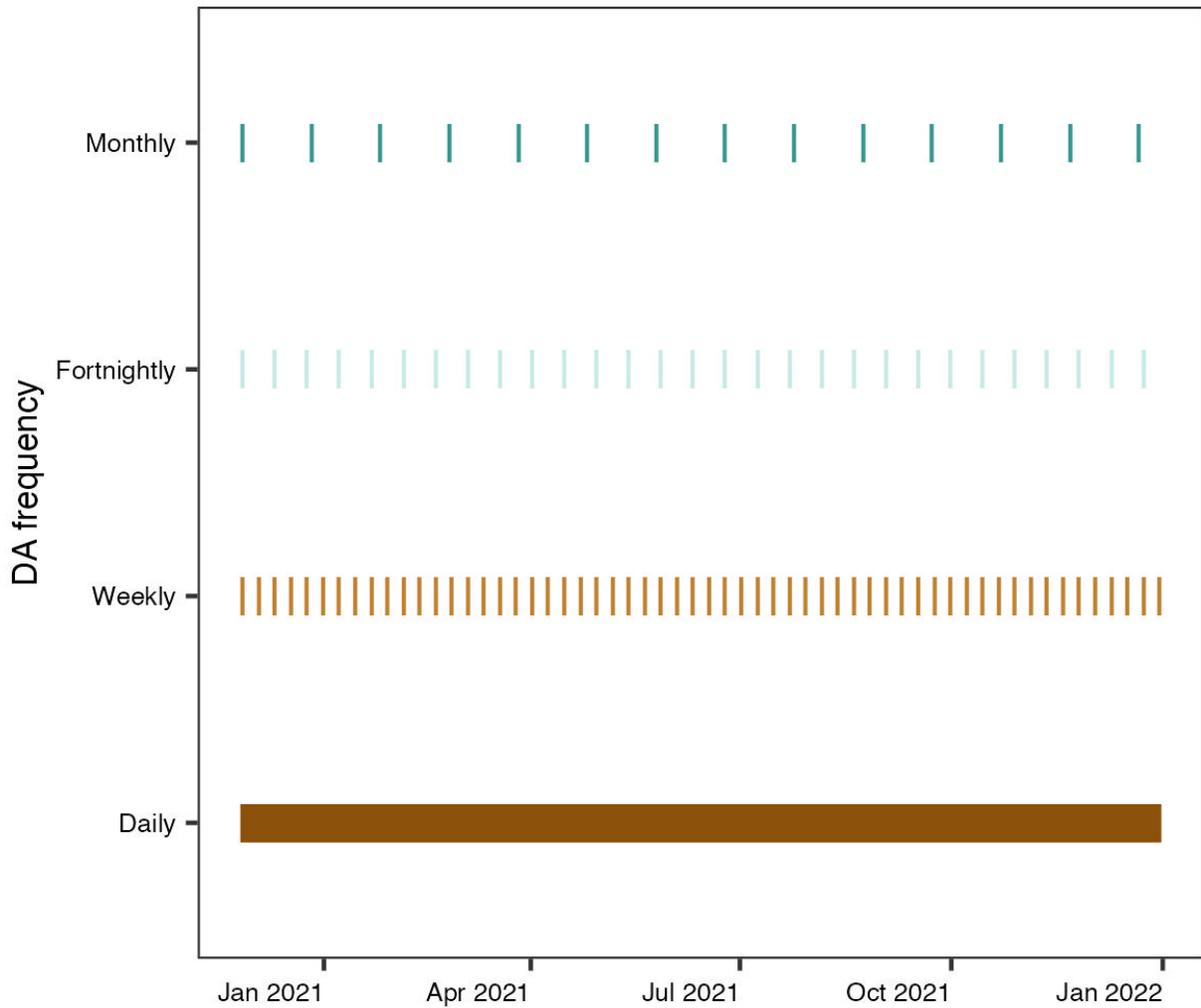


Figure S1: Frequencies for daily, weekly, fortnightly, and monthly data assimilation (DA); lines indicate the dates when DA occurred. For example, daily DA occurred every day from 27 November 2020 to 31 December 2021, whereas monthly DA occurred 14 times during the 14-month period of November 2020 to January 2022.

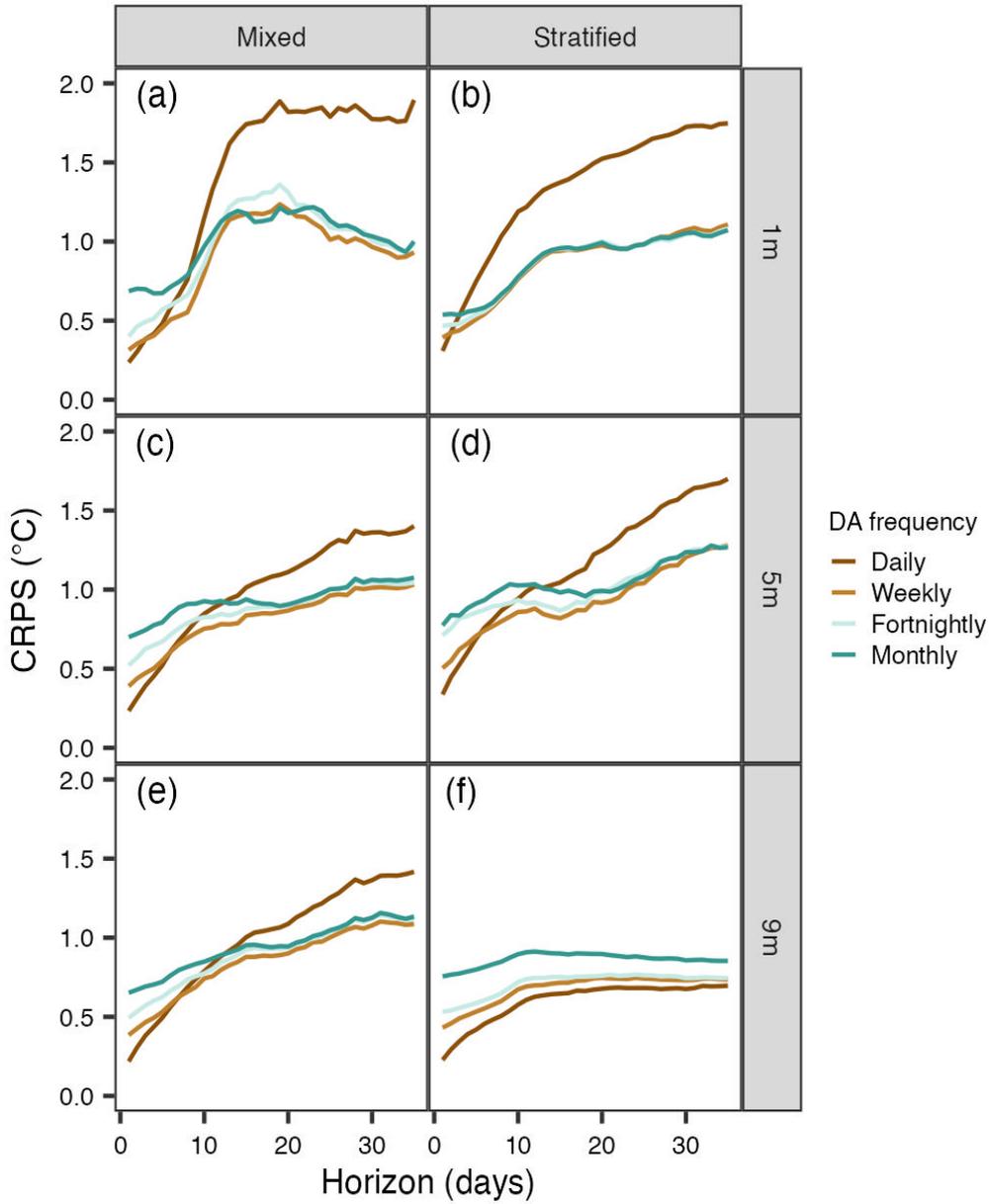


Figure S2: Water temperature forecast continuous ranked probability score (CRPS) across different depths (1 m: a, b; 5 m: c, d; and 9 m: e, f) and horizons in Beaverdam Reservoir during the mixed (a, c, e) vs. stratified (b, d, f) periods. Each grouping of bars represents a different data assimilation (DA) frequency (daily, weekly, fortnightly, monthly). Depths are indicated in the right y-axis labels; horizons are depicted by colored boxplots.

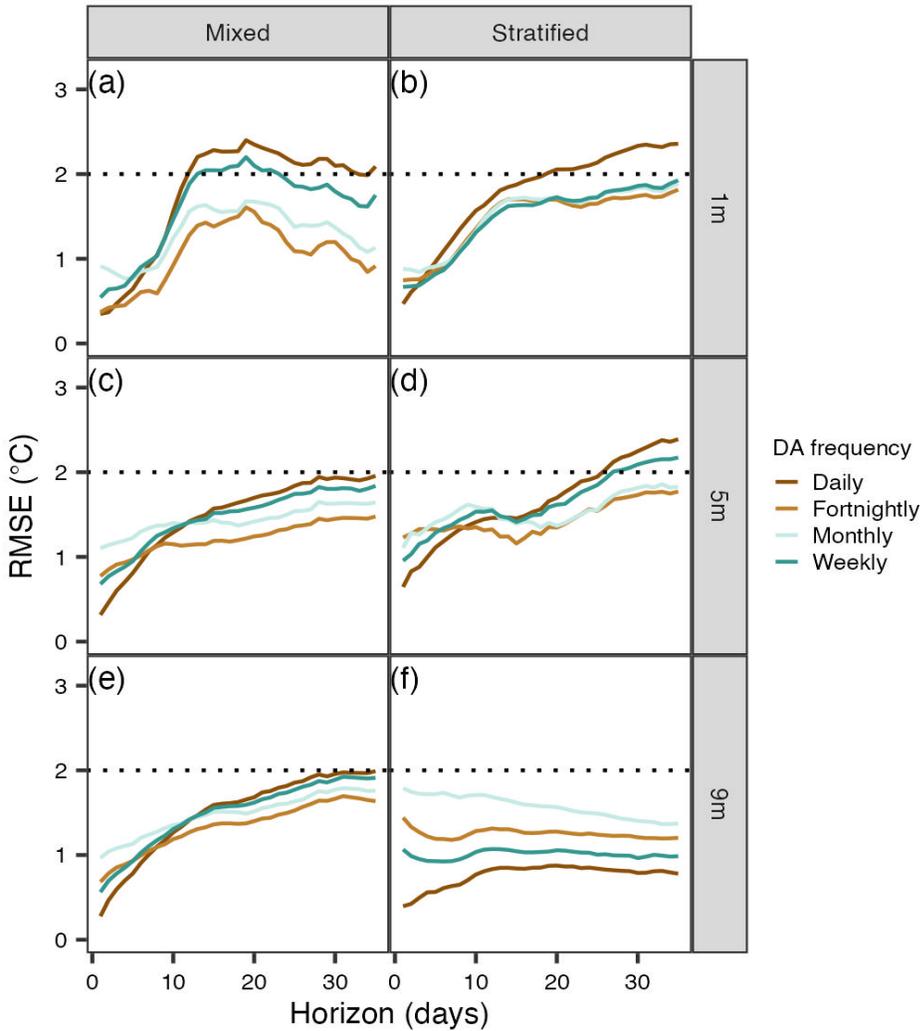


Figure S3: Root mean square error (RMSE) calculated from comparing water temperature observations with forecasts that did not include initial conditions uncertainty for 1-35-day-ahead forecasts in Beaverdam Reservoir during the mixed (a, c, e) vs. stratified (b, d, f) periods at 1 m (a, b), 5 m (c, d), and 9 m (e, f). RMSE for each forecast horizon was averaged across the 365-day forecast period (1 January - 31 December 2021). Colored lines correspond to different data assimilation (DA) frequencies (daily, weekly, fortnightly, and monthly); dotted horizontal lines depict the 2°C RMSE threshold for skillful forecasts.

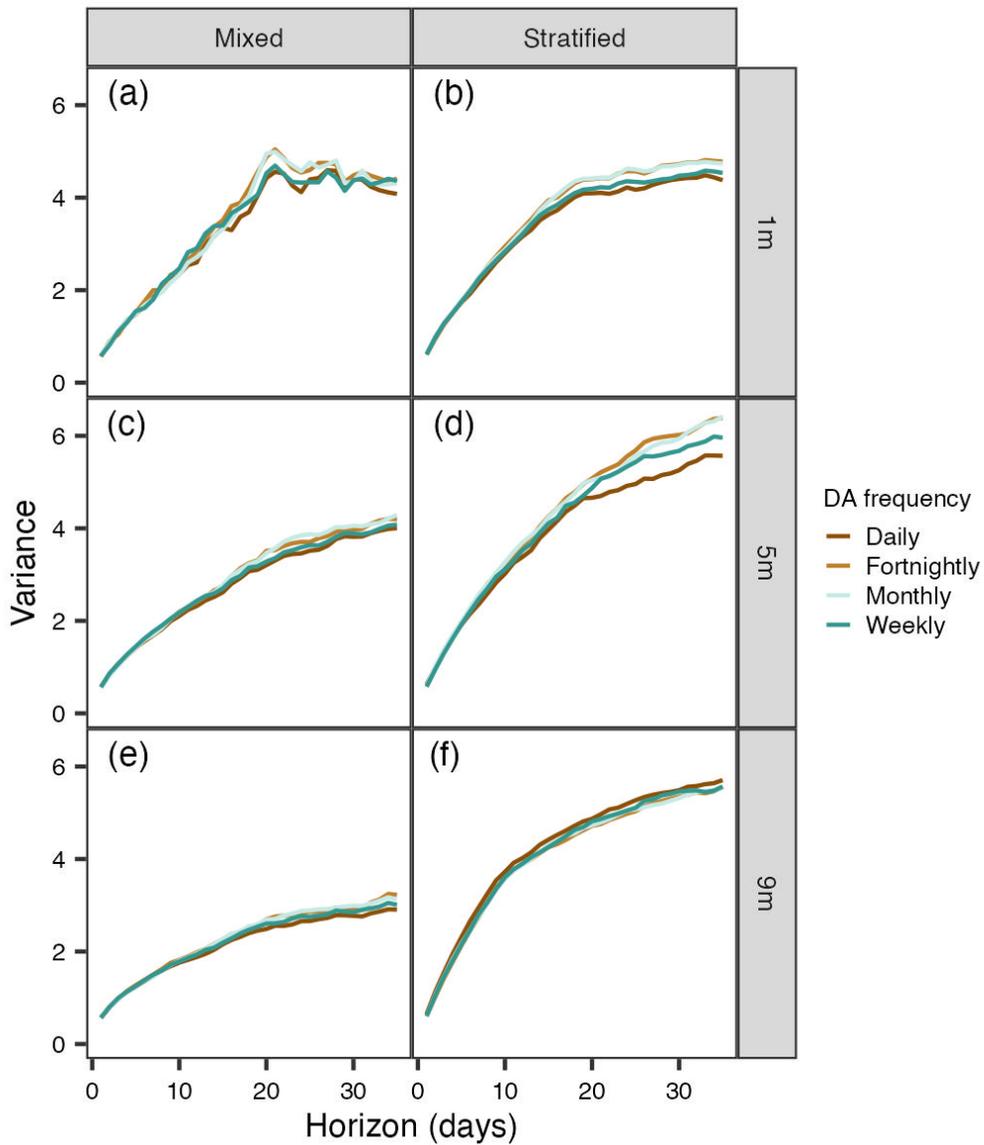


Figure S4: Variance for forecasts that did not include initial conditions uncertainty across horizons (1-35 days ahead) in Beaverdam Reservoir during the mixed (a, c, e) vs. stratified (b, d, f) periods for 1 m (a, b), 5 m (c, d), and 9 m (e, f). Variance for each forecast horizon was averaged from all 365 forecasts generated during the forecast period (1 January - 31 December 2021). Colored lines correspond to different data assimilation (DA) frequencies (daily, weekly, fortnightly, and monthly).