

# Why are Mountain-Tops Cold?

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## The Decorrelation of Surface Temperature and Topography Due to the Decline of Greenhouse Effect on Early Mars

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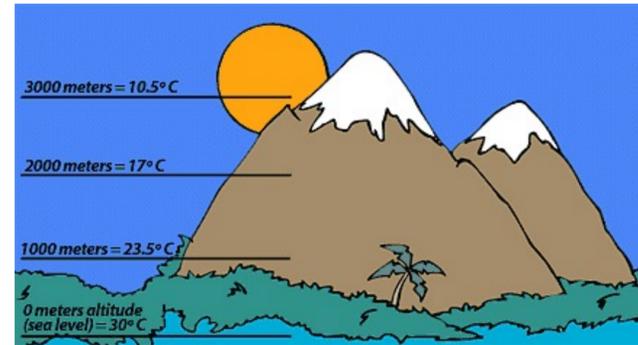
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### Motivation

On Earth, mountain-tops are cold because the temperature follows atmospheric lapse rate.

- Mars lost its CO<sub>2</sub>-dominated atmosphere over time, from  $\leq 2$  bar around 4 Ga to 6 mbar today (Jakosky et al., 2018; Warren et al., 2019).
- The atmospheric evolution of Mars was accompanied by climate change, which was recorded by shifts in the spatial distribution of rivers and lakes (Kite, 2019).
- Climate models find shifts in surface temperature pattern with decreasing atmospheric CO<sub>2</sub> (Wordsworth, 2016). When the CO<sub>2</sub> atmosphere is thick, T<sub>s</sub> decreases with height (correlated with topography); when the atmosphere is thin, T<sub>s</sub> only depends on insolation (decorrelation with topography).
- **Question: What mechanism is responsible for this decorrelation? What are the implications for Mars' climate change?**



What about Mars?

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### Surface Energy Budget Indicates that the Greenhouse Effect Controls the Decorrelation

$$SW + LW_a = LW_s + SH$$

- Under steady state, the surface energy budget is the balance between **shortwave heating from the sun (SW)**, **longwave heating from the atmosphere (LW<sub>a</sub>)**, **surface cooling by emission (LW<sub>s</sub>)**, and **the cooling by sensible heat flux (SH)**.
- For the 1<sup>st</sup> time, we find that **the decorrelation of T<sub>s</sub> (scales with emission) with topography is due to the decrease of CO<sub>2</sub> greenhouse effect**, not due to **the sensible heat flux** as proposed by Wordsworth (2016) and Kite (2019).

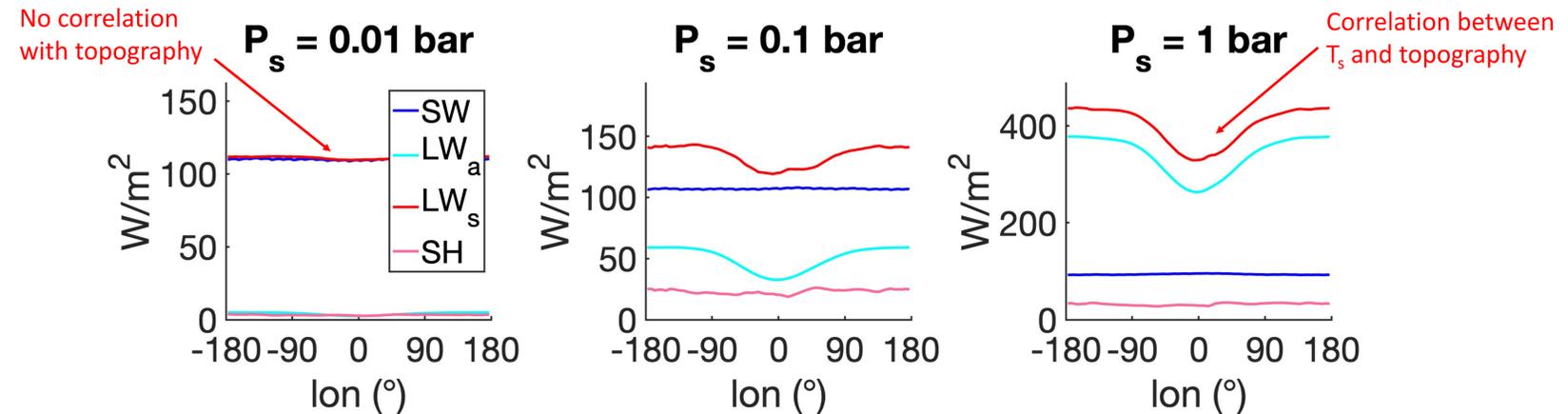


Fig. 2: Time-averaged surface energy budgets in runs. Each term is averaged within the tropics (20°N - 20°S). The red curve with a dip represents the correlation between T<sub>s</sub> and topography (lower T<sub>s</sub>/emission over the mountain), which is controlled by the decrease of greenhouse heating. These examples are performed with Gaussian topography, correlated-k CO<sub>2</sub> scheme, and diurnal-mean insolation, but changing topography, radiation scheme, or diurnal cycle do not change our conclusion (Fig. 3).

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### Methods

- Model: MarsWRF Mars GCM (Richardson et al., 2007; Toigo et al., 2012).
- Solar luminosity: 85% of the modern value.
- Obliquity and eccentricity set to zero.
- Ice-albedo feedback disabled.
- Atmospheric thickness varied from 0.01 bar to 3 bar.
- Radiation: CO<sub>2</sub> (correlated-k scheme) or gray gas (absorption in the IR wavelength).
- Topography: 2-D Gaussian topography (Fig. 1)

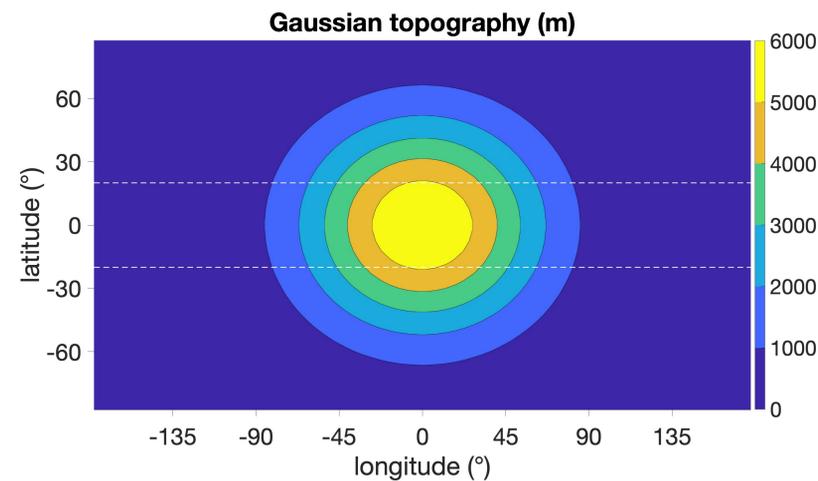


Fig. 1: The idealized Gaussian topography. A 6000-meter-high mountain is placed at the equator (comparable to Tharsis on Mars). The white dash lines are the boundary for tropical averaging in Fig. 2. This topography is labelled as "Gaussian" in Fig. 3 (represents the case when highlands only cover a small fraction of the surface). We also simulated with "iGaussian" topography, which means the surface elevation is "Gaussian  $\times -1$ ". The "iGaussian" case represents when lowlands only cover a small fraction of the surface.

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### Reducing the Complexities from GCM

- The decorrelation can be reproduced by decreasing IR opacity under the gray scheme, without changing atmospheric thickness.
- **→ Changes in Mars' fluvial/temperature patterns may arise from changes in non-CO<sub>2</sub> GH gases (e.g., H<sub>2</sub>), rather than the lost of CO<sub>2</sub>-dominated atmosphere.**
- A 3D model is necessary to determine the correct strength of GH forcing for the decorrelation.
- **→ Atmospheric circulation is important to redistribute the energy between highlands and lowlands.**
- Other factors (diurnal cycle, topography) do not matter within our idealized simulations.
- Next step: a conceptual framework to explain the decorrelation.

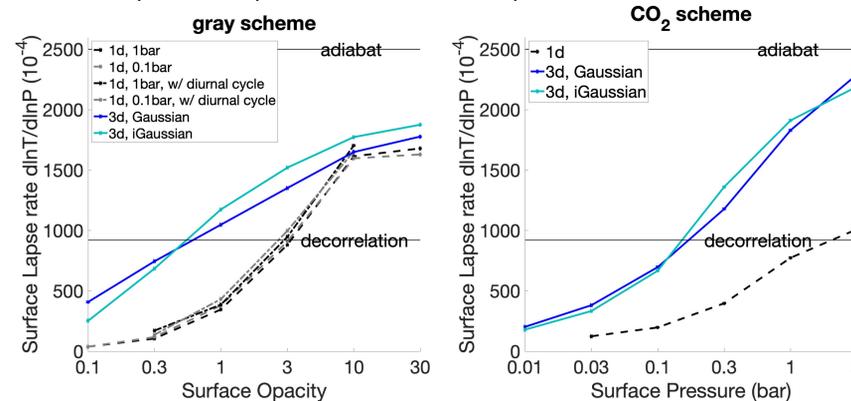


Fig. 3: The relation between T<sub>s</sub> and topography in simulations with gray gas scheme (left) and CO<sub>2</sub> scheme (right). The relation is quantified as surface lapse rate  $\left(\frac{d\ln T_s}{d\ln P_s}\right)$ . When the greenhouse effect is strong, the lapse rate approaches to the atmospheric lapse rate (adiabat); when the greenhouse effect is weak, the lapse rate in T<sub>s</sub> approaches to zero. We define the decorrelation point as the adiabat divided by Euler's number e.

### References

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