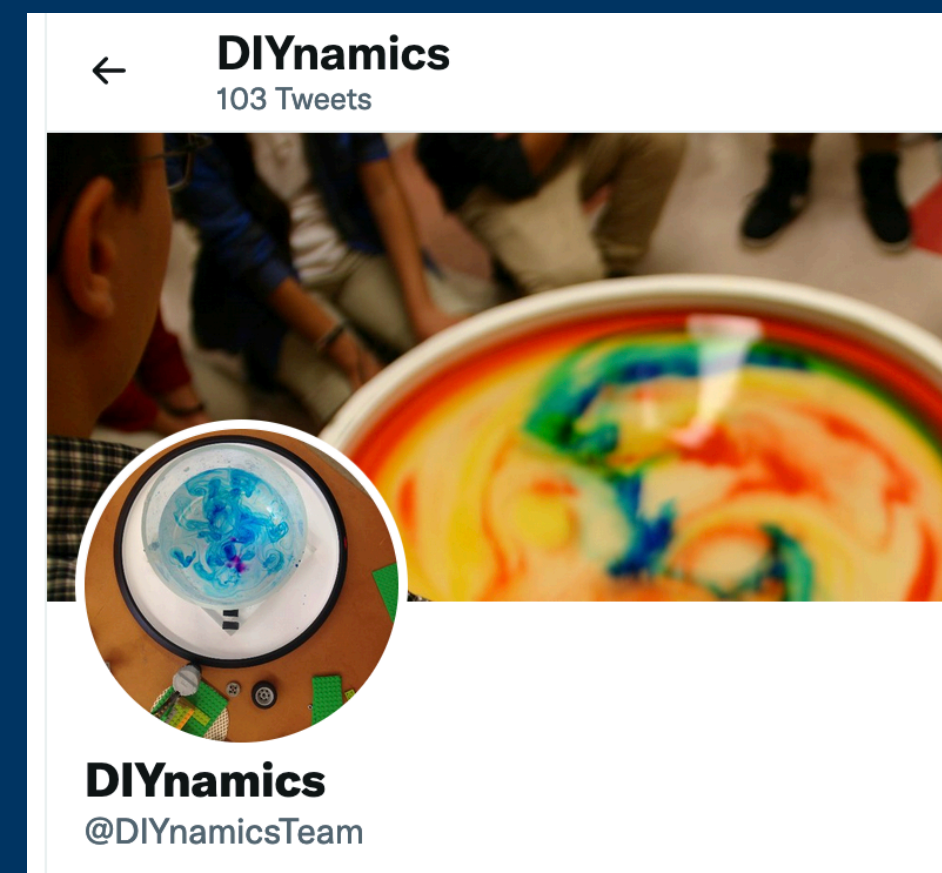


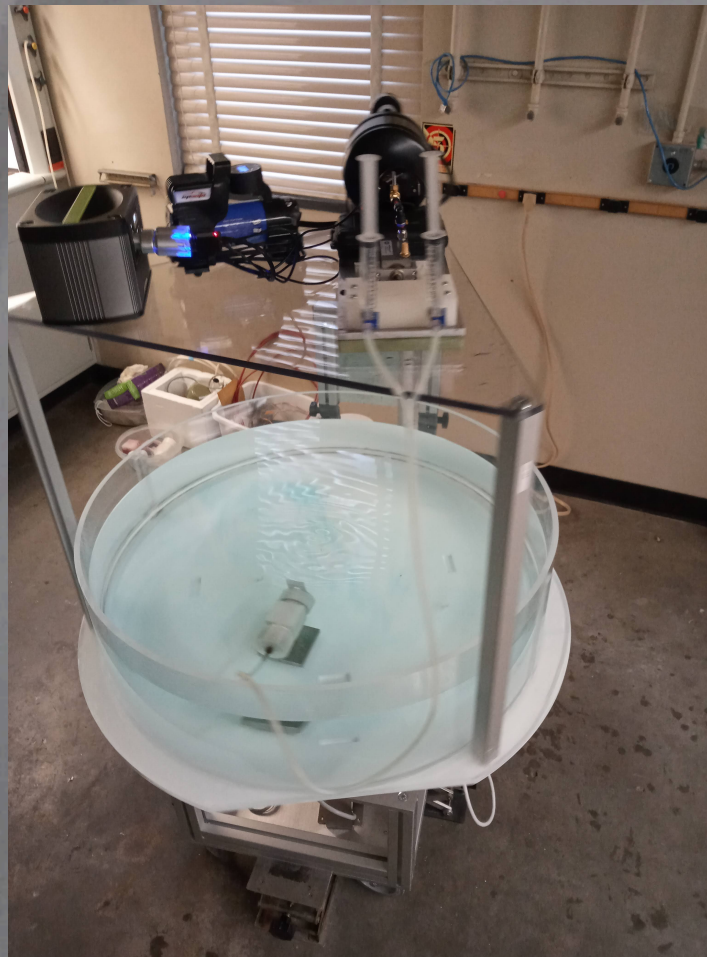
# The Non-traditional Coriolis force drives Westward tilts in Multiscale Theories and Laboratory Experiments of Tropical Dynamics

Joseph A. Biello, Matt R. Igel & Michael D. Toney  
University of California, Davis



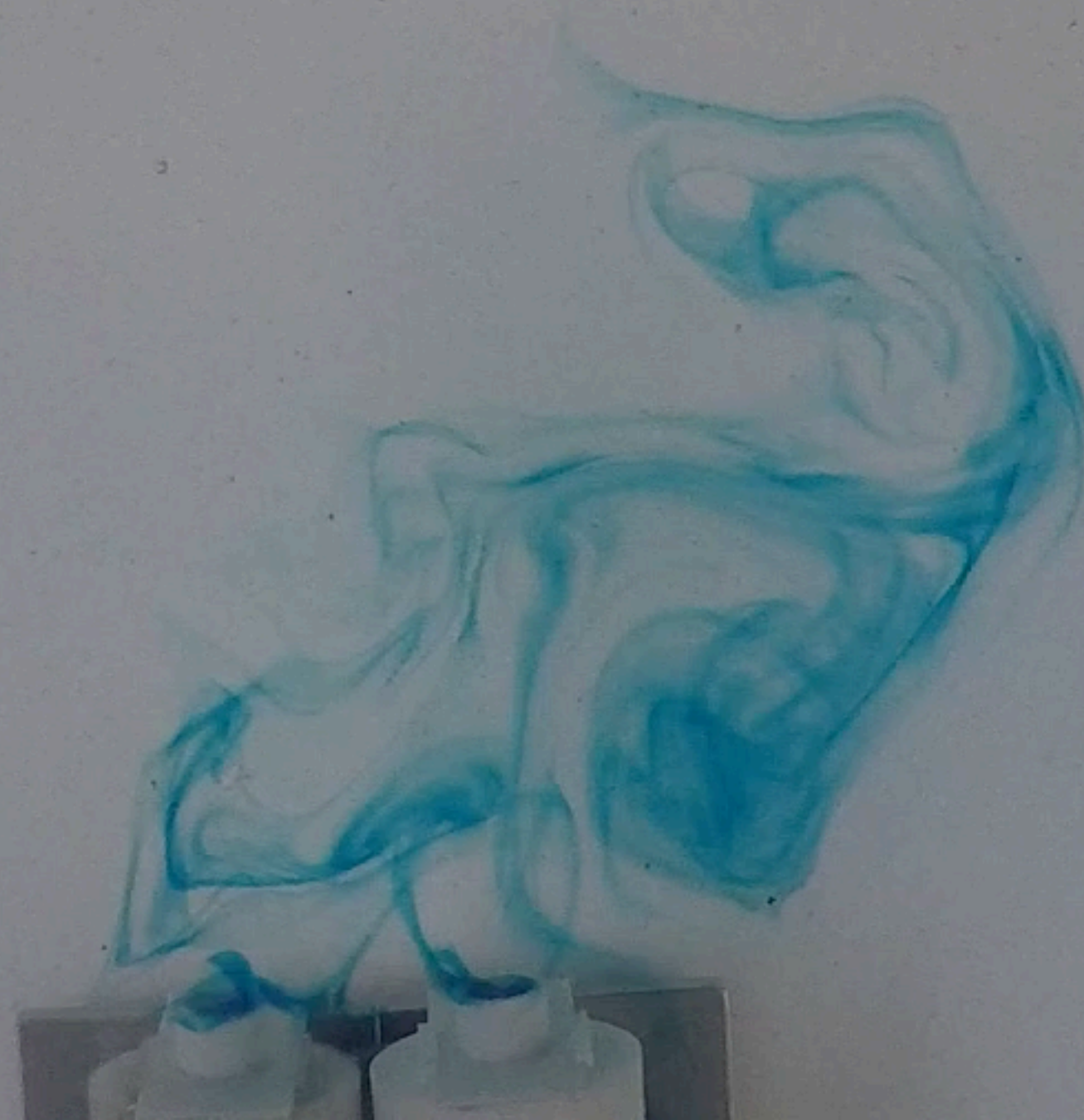


# DONUT



- Counterclockwise rotating tank
- 36 second period
- Seen from above
- 12 cm depth of water
- Horizontal injection of toroidal vortices
- Water + Methyl blue
- 4 pairs of vortex rings

**IPESD + MJO**





# The Multiscale Models of Tropical Dynamics

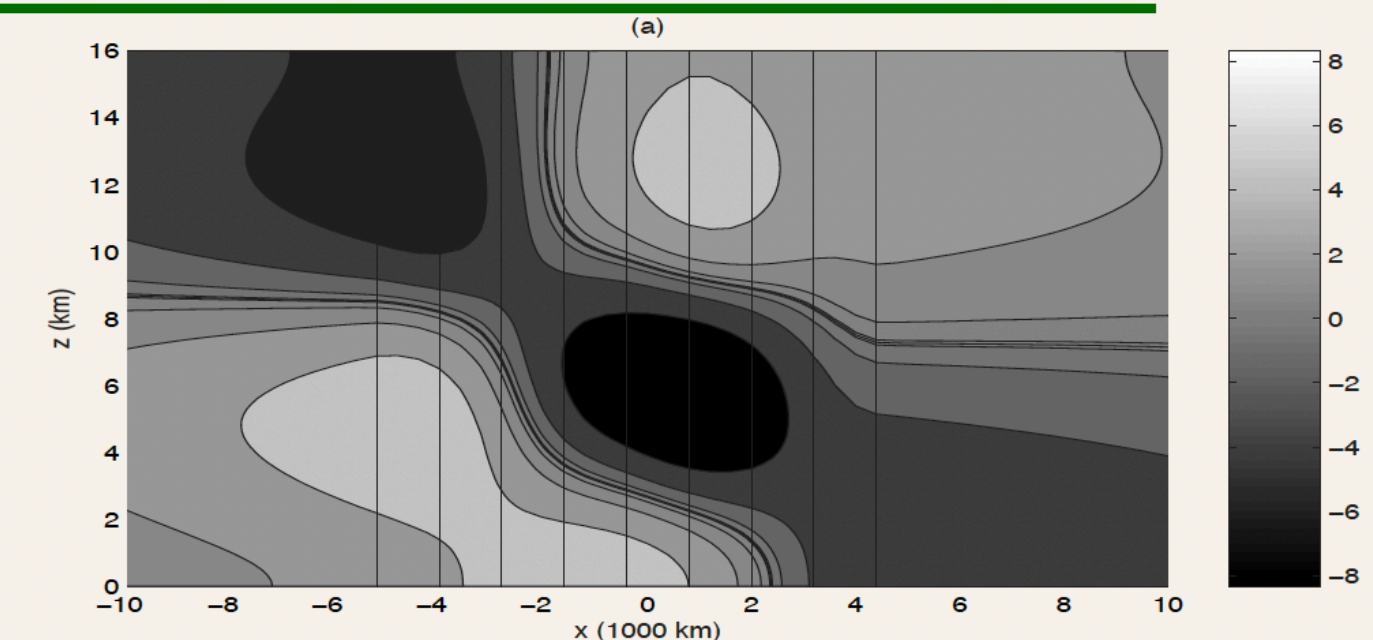
- J.A.B. & Majda (early 2000s) on the MJO
- Majda & Klein (2003)
- The multi scale models are a multiscale system of PDEs wherein diabatic heating on the synoptic scale drives the large scale circulation

$$\begin{aligned}
 \bar{U}_t - y\bar{V} + \bar{P}_X &= F^U - d_0 \bar{U} \\
 y\bar{U} + \bar{P}_y &= 0 \\
 \bar{\Theta}_t + \bar{W} &= F^\theta - d_\theta \bar{\Theta} + \bar{S}_\theta \\
 \bar{P}_z &= \bar{\Theta} \\
 \bar{U}_X + \bar{V}_y + \bar{W}_z &= 0
 \end{aligned}$$

$$\begin{aligned}
 F^U &= -\overline{(v' u')_y} - \overline{(w' u')_z} \\
 F^\theta &= -\overline{(v' \theta')_y} - \overline{(w' \theta')_z}
 \end{aligned}$$

## Equatorial MJO model: Winds above the equator

- Lower troposphere  
congestus  
heating in the east
- Westward tilted anvil  
superclusters in the west



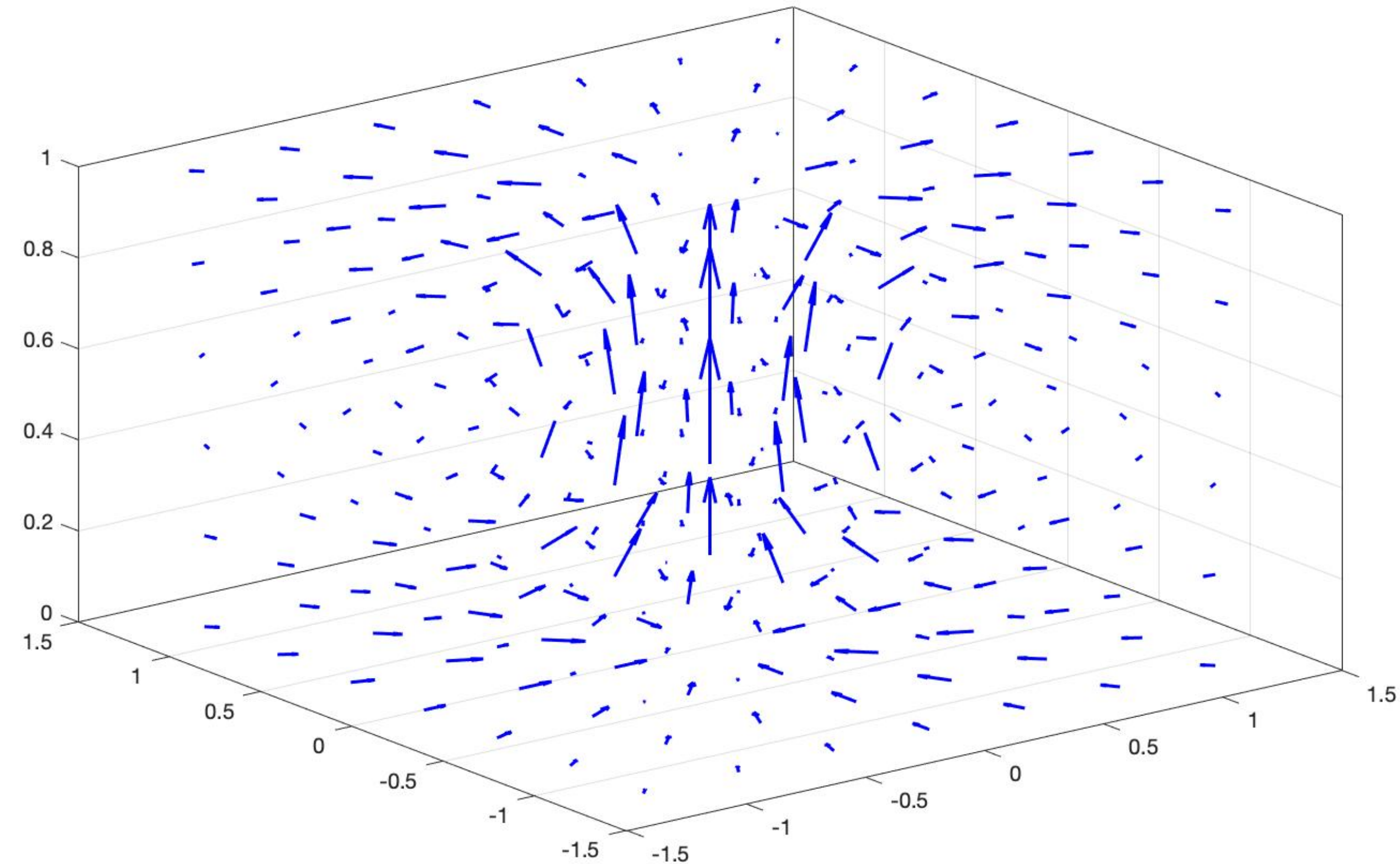
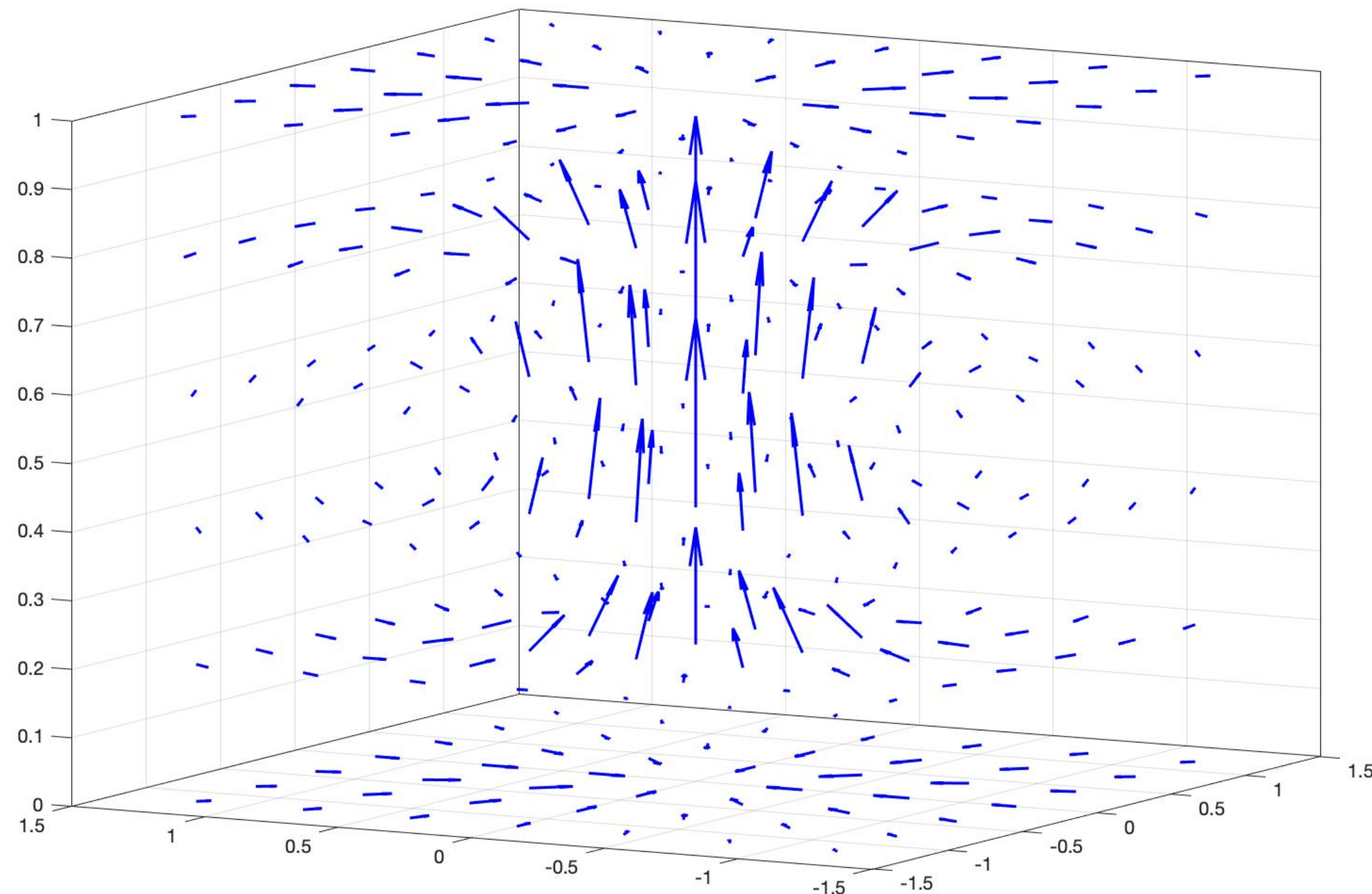
- Vertical/Westward tilted convection ( $w'u'$ ) is essential, but unexplained



# Dynamics of Non-rotating Updraft Tori



- w/ Igel, we are developing dynamical convective models consisting of poloidal circulations
- poloidal circulations are **tori**



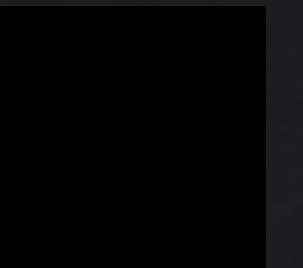


# Some Other DoNUT Talks/Posters

- ◆ A25D-1709: The Divergence-free Net Coriolis Force creates Vertical Geostrophic Balance and Westward tilted Atmospheric Convection
- ◆ NG33A-06: The Non-Traditional Coriolis Terms and Their Impact on the Convective Weak Temperature Gradient

\* A22A-08: Modeling Tropical Convective Clouds and Circulations with the DoNUT

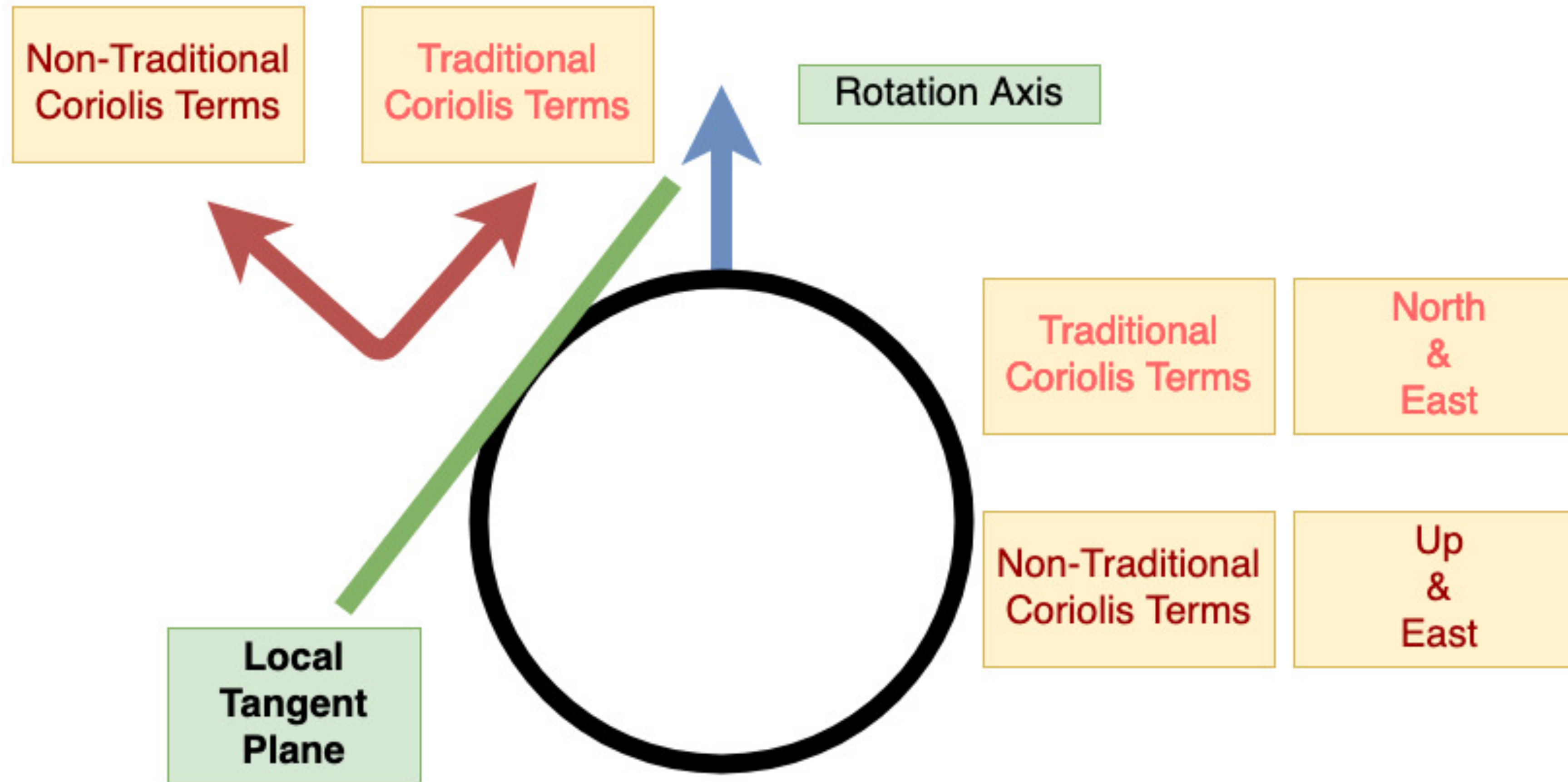
- ◆ Focus in these talks is in using the DoNUT (or a conceptually simplified version) to understand the impact of physics on convective circulations.





# The Coriolis Force

The traditional simplification and the non-traditional terms





# The Coriolis Force

## The traditional simplification and the non-traditional terms

- The non-traditional terms are underlined
- The traditional terms vanish at the equator
- The non-traditional terms are negligible (and rightfully neglected) for flows which are horizontally much larger than the height of the troposphere
- The non-traditional terms, NCT, are significant for convective scale flows at the equator.
- Igel & B., JAS, 77, 2020

$$\frac{Du}{Dt} + \frac{\partial p}{\partial x} = 2\Omega_0 \sin(\phi)v - \underline{2\Omega_0 \cos(\phi)w},$$

$$\frac{Dv}{Dt} + \frac{\partial p}{\partial y} = -2\Omega_0 \sin(\phi)u,$$

$$\frac{Dw}{Dt} + \frac{\partial p}{\partial z} = \underline{2\Omega_0 \cos(\phi)u} + B.$$



# The Net Coriolis Force

## Using the Leray projection to determine the pressure

- The net Coriolis force is the divergence-free force resulting from the pressure gradient in the Leray projection

$$\vec{F} = -\nabla p - 2\vec{\Omega} \times \vec{u}, \quad \nabla \cdot \vec{F} = 0 \quad \vec{F} \cdot \hat{n} = 0 \quad \text{on rigid boundaries,}$$

- Using the Leray projection we can compute the pressure and net force

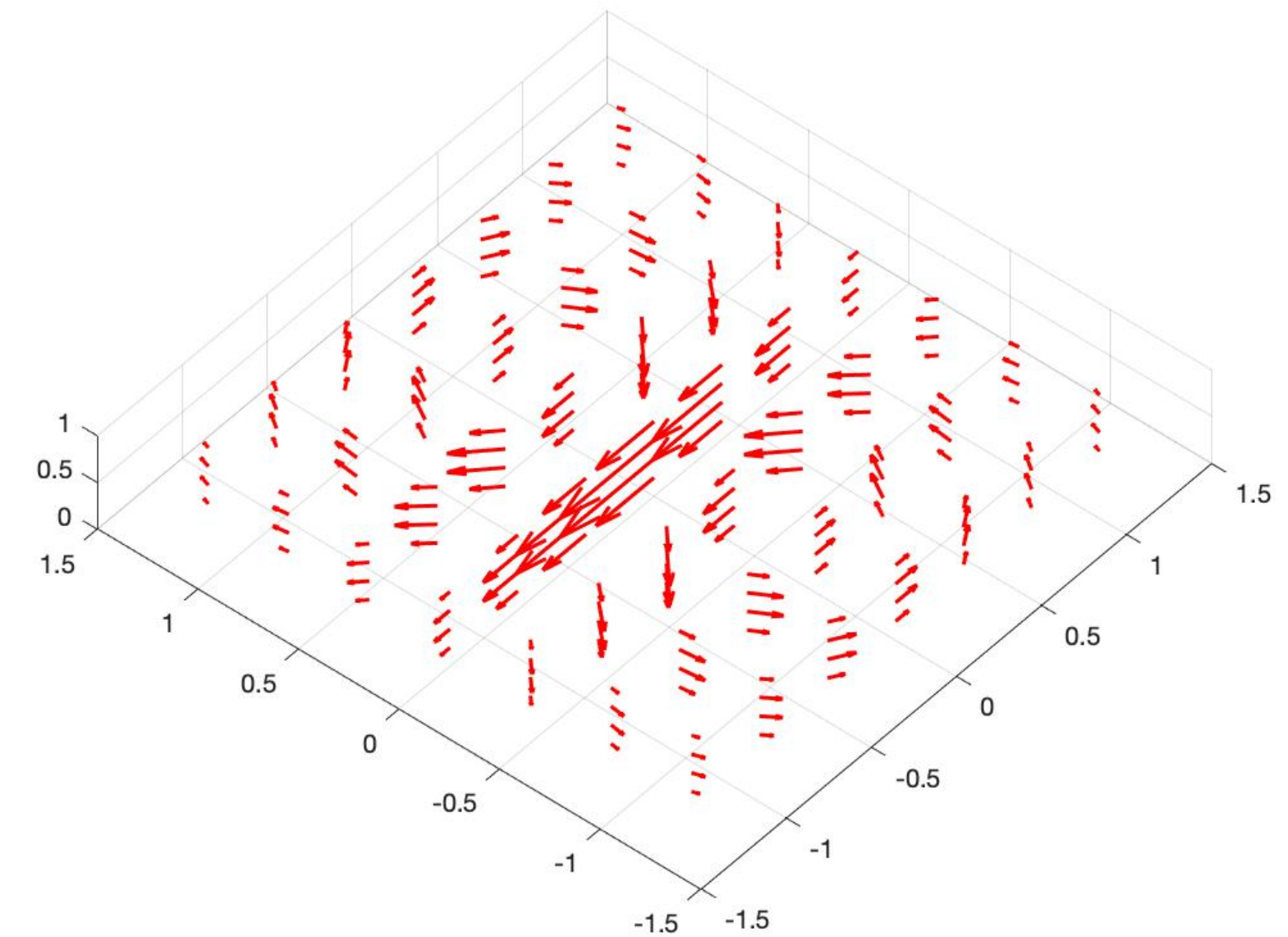
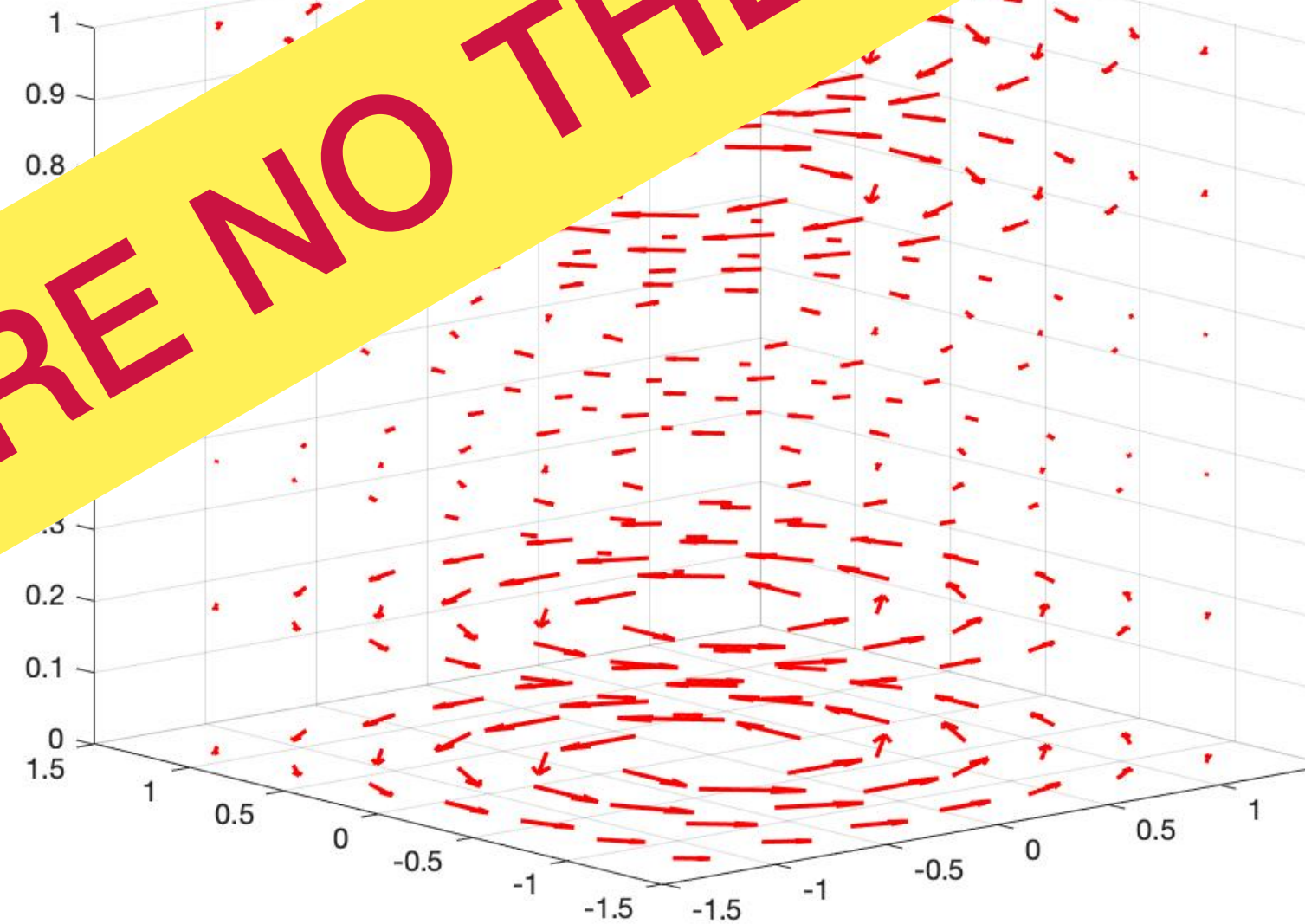
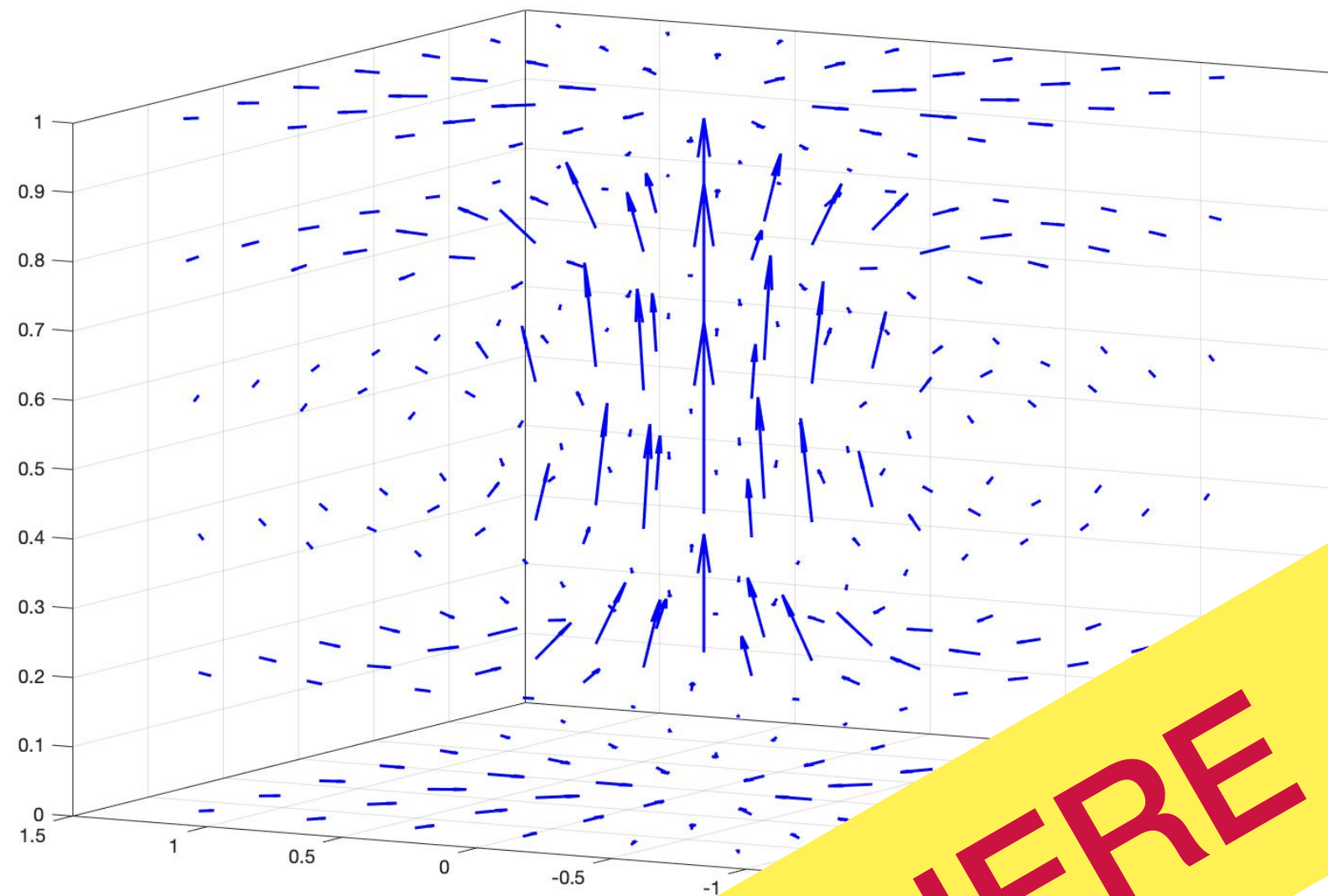
$$\begin{aligned} \nabla^2 p &= 2\vec{\Omega} \cdot \vec{\omega}, \\ \hat{k} \cdot \nabla p &= 2 \left( \vec{u} \times \vec{\Omega} \right) \cdot \hat{k}, \quad \text{on } z = 0. \end{aligned}$$



**Theorem:** The vertical component of the net Coriolis force is identically zero for Poloidal Flow

**Consequence:** The pressure is computed as a vertical integral

**THERE ARE NO THEOREMS AT AGU!!**





# Theorem: The vertical component of the net Coriolis Force is identically zero for Poloidal Flows

**Consequence:** The pressure is computed as a simple vertical integral

$$p = -2\Omega_0 \cos(\lambda) \int_z^\infty u(x, y, z') dz';$$

$$\Phi = \int_z^\infty \phi(x, y, z') dz'.$$

$$\vec{F}_{\text{TCT}} = -2\Omega_0 \sin(\lambda) U \hat{\theta}$$

$$\vec{F}_{\text{NCT}} = 2\Omega_0 \cos(\lambda) \left[ \Phi_{yy} \hat{i} - \Phi_{xy} \hat{j} \right]$$

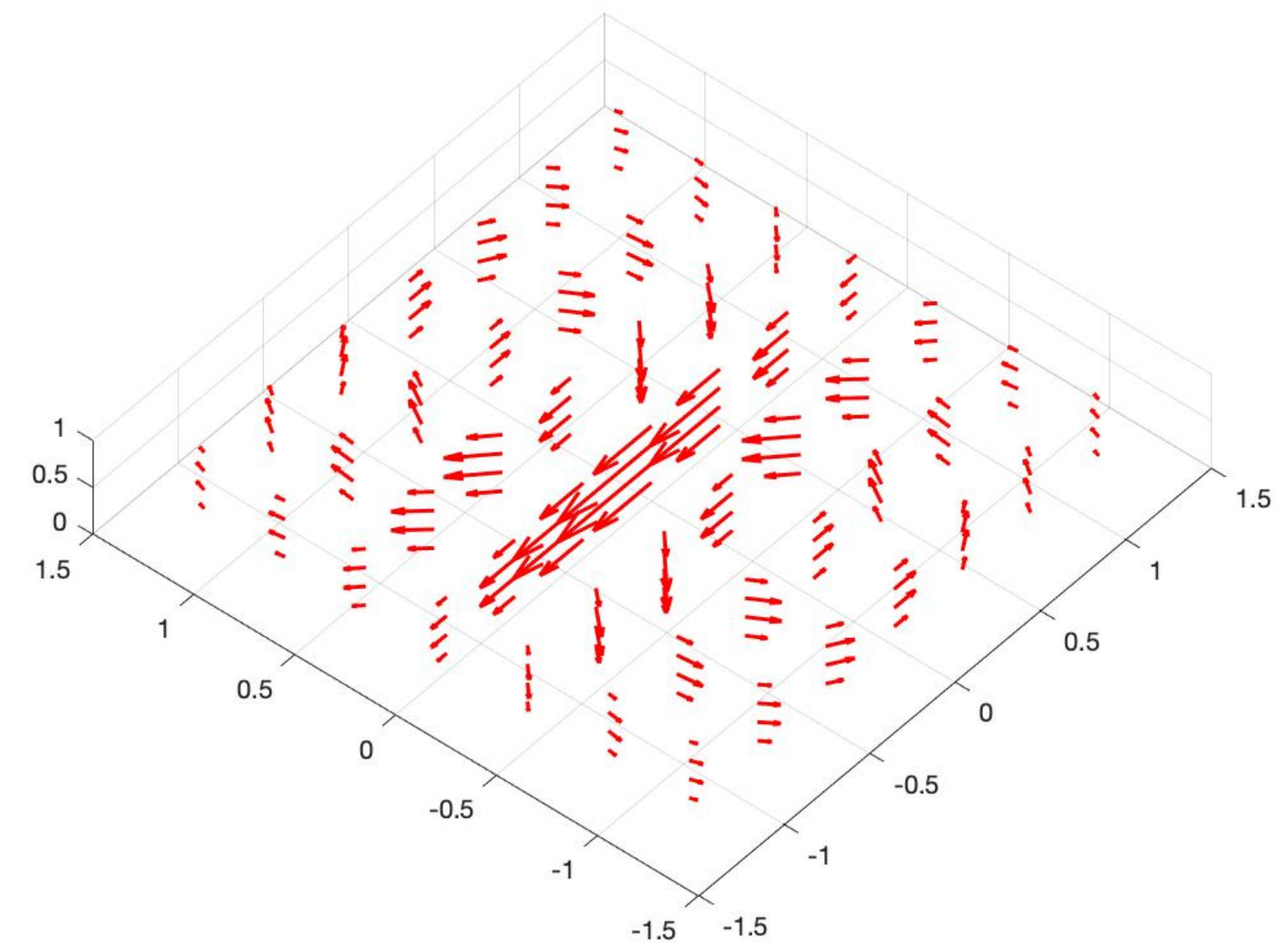
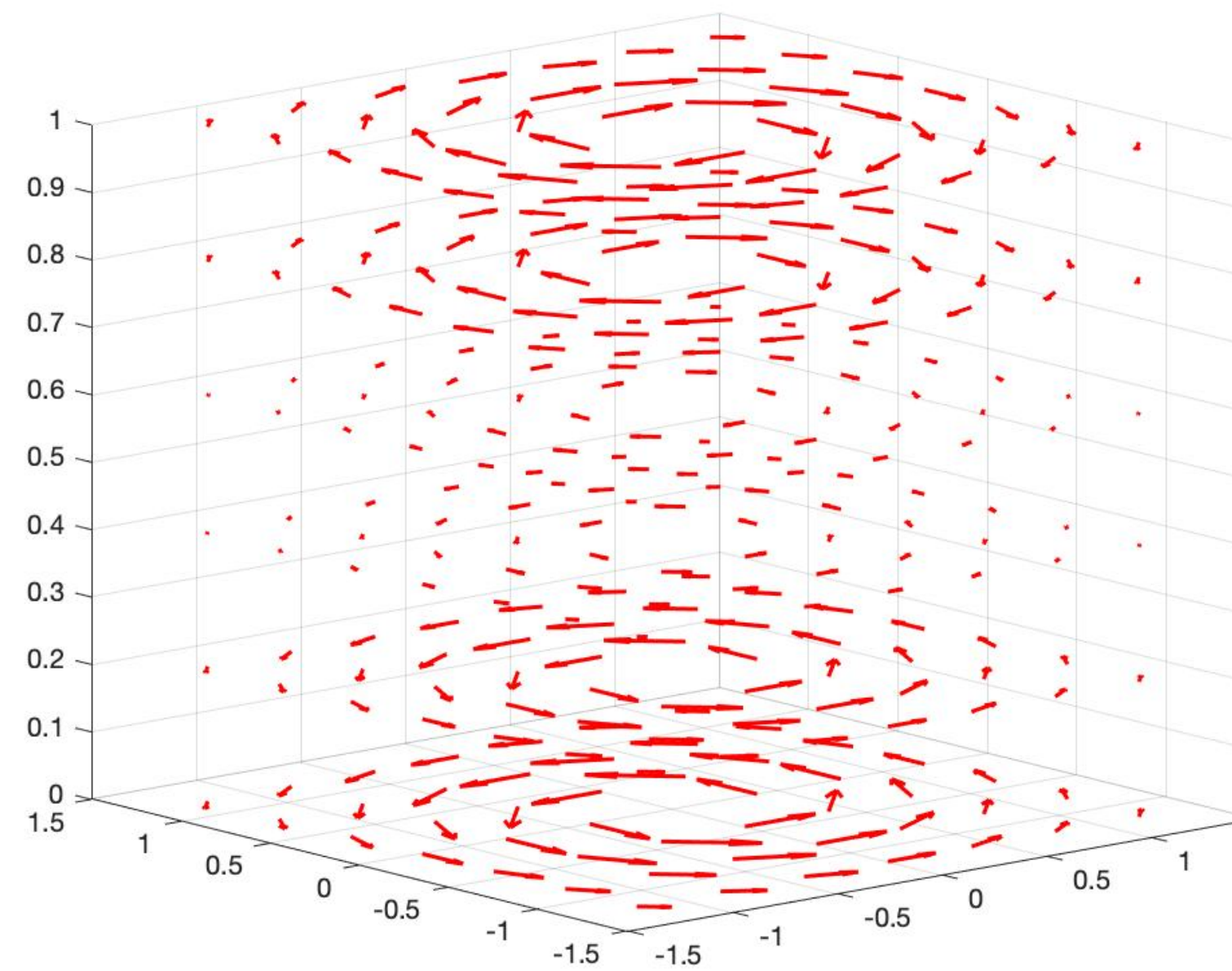
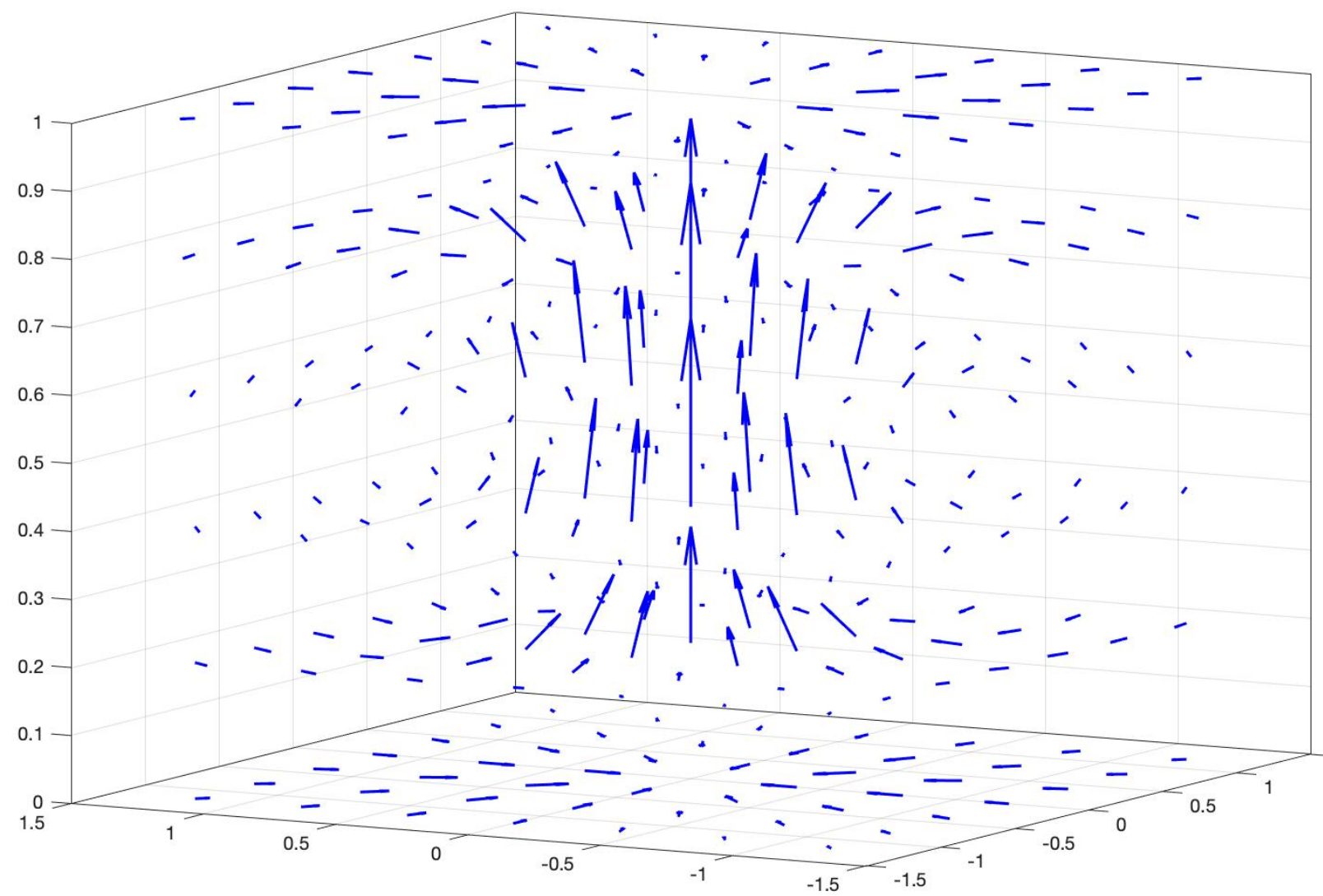
Tropical Cyclogenesis

MJO-genesis?



# **Theorem:** The vertical component of the net Coriolis Force is identically zero for Poloidal Flows

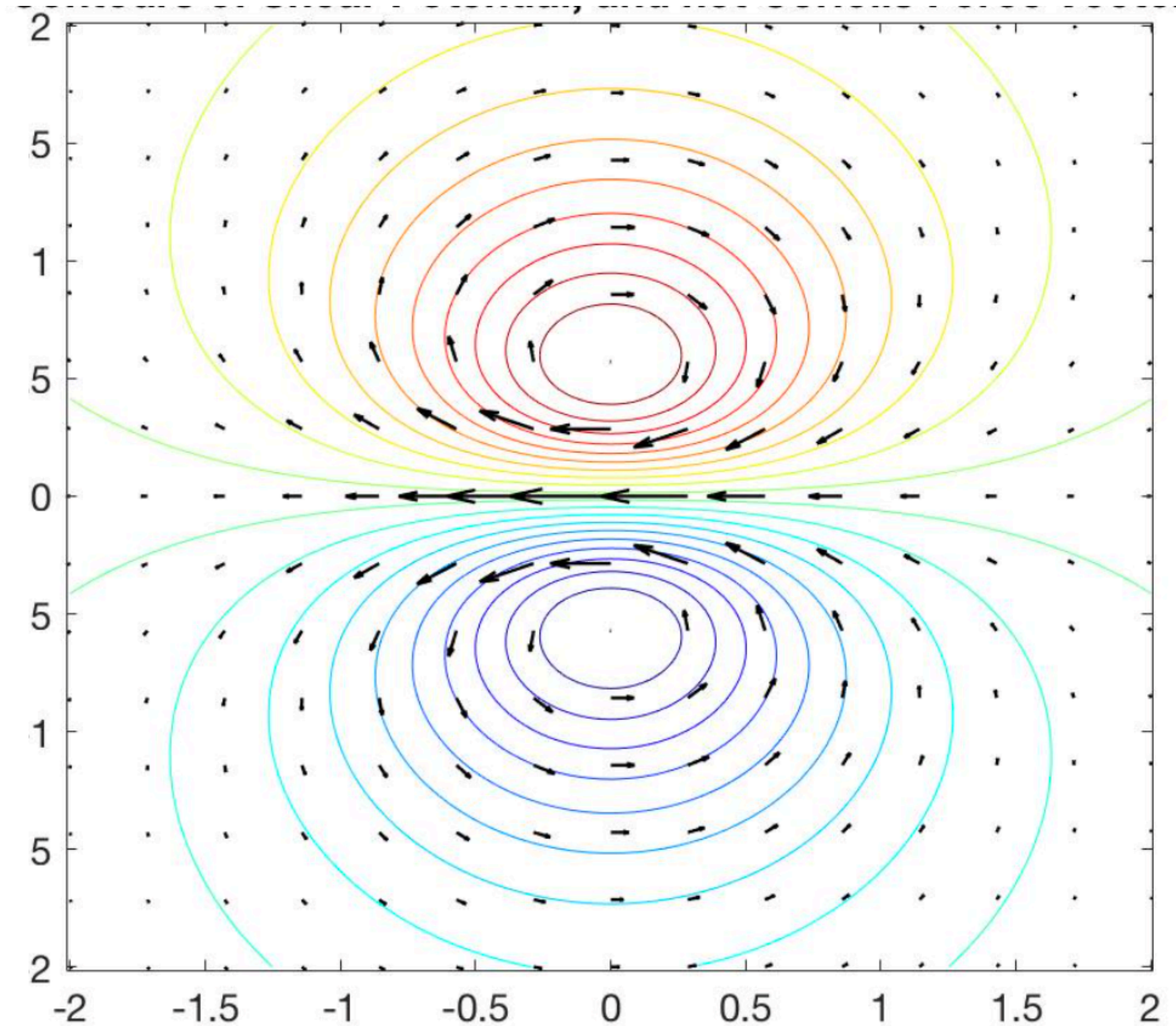
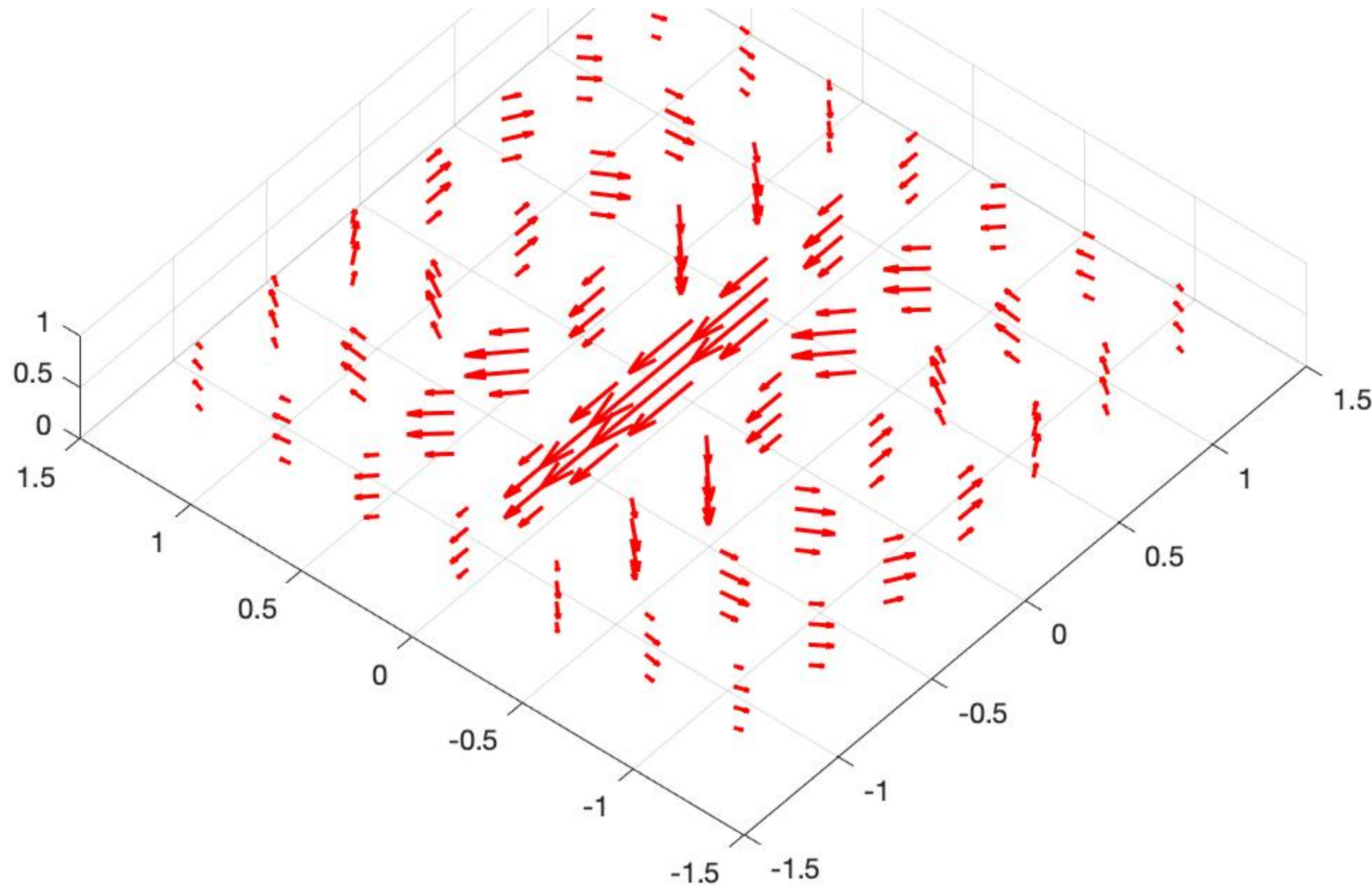
**Consequence:** The pressure is computed as a simple vertical integral





# The Net NCT for Poloidal flows

Westward in the up flow, eastward in the down flow



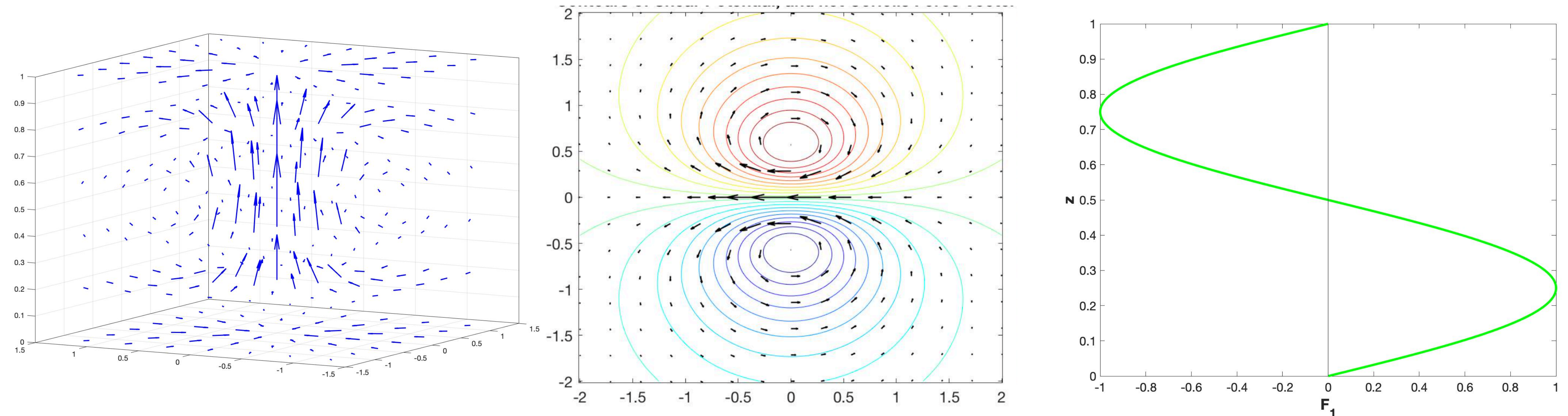


# Vertical Upscale Flux of Zonal Momentum

Lower tropospheric westerlies, upper tropospheric easterlies

$$\vec{F} = -(\overline{uw})_z \hat{i}$$

$$\overline{F}_1 = \frac{f}{H} \left( \frac{\Omega}{D} \right) \left( \frac{S_0^2}{\Gamma^2} \right) \sin \left( \frac{2z}{H} \right)$$



**The DoNUT circulation + Net Non-Traditional Coriolis force provide the vertical upscale fluxes of zonal momentum which were needed in the Multiscale MJO Models of Majda & B.**





Equatorial Non-  
Traditional  
Coriolis Force

**IPESD + MJO**

Large Scale Lower  
tropospheric  
Westerly Wind

