

# Observed Changes in Daily Precipitation Intensity in the United States

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Submitted to Geophysical Research Letters

25 February, 2022

## Key Points

- We observe consistent shifts from lower to higher daily precipitation intensity, particularly in the central and eastern United States
- Mean and standard deviation of wet day precipitation intensities increase for nearly all domains within the United States
- Fourteen of seventeen domains show differences in wet day precipitation intensity distributions between 1951-1980 and 1991-2020

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24 **Abstract**

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26 Relative to changes in annual mean and extreme precipitation, the characterization of changes  
27 in the full distribution of precipitation intensities remains overlooked and under-explored,  
28 despite their critical importance to natural hazard, agriculture, and infrastructure risk  
29 assessments. Here, we aggregate daily *in situ* Global Historical Climatology Network  
30 precipitation observations within seventeen internally consistent NEON domains in the United  
31 States for two time periods (1951-1980 and 1991-2020). We find statistically significant changes  
32 in wet day precipitation distributions in fourteen of the domains – changes primarily driven by  
33 a shift from lower to higher wet day intensities. Patterns of robust change are geographically  
34 consistent, with increases in the mean (4.6-7.1%) and standard deviation (20-31%) of wet day  
35 intensity in the eastern U.S., but mixed signals in the western U.S. Beyond their critical  
36 importance to the aforementioned societal impact realms, these observational results can also  
37 inform climate model performance evaluations.

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40 **Plain Language Summary**

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42 Lots of research has been done to see how rainfall and snowfall event totals are affected by  
43 climate change. While most studies look at yearly totals or extreme events, we look at how daily  
44 precipitation is changing at all intensity levels, which has effects on agriculture, infrastructure,

45 landslides, and flooding. We group daily rain gauge measurements within seventeen climate  
46 regions in the United States for two thirty-year time periods. We find changes in daily  
47 precipitation intensity in fourteen regions, changes that are mostly caused by a shift from lower  
48 to higher intensity rain and snow events. We also identify broader areas in the eastern U.S. with  
49 consistent increases in the average and daily variability of precipitation, but changes are mixed  
50 in the western U.S. In addition to the impacts mentioned above, our results can also be used to  
51 assess how well climate models perform.

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#### 54 **Keywords**

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56 precipitation, precipitation variability, precipitation intensity distribution, daily, GHCN, NEON

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#### 59 **1. Introduction**

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61 Anthropogenic climate change is driving shifts in global precipitation patterns (Douville et al.,  
62 2021). Recent studies have characterized these shifts across a diversity of metrics and scales,  
63 including annual totals, frequencies of occurrence, and zonal distributions. At the daily scale,  
64 substantial recent efforts have demonstrated robust changes in extreme precipitation intensities  
65 (i.e., the 95th percentile and above). However, characterization of changes in the full  
66 distribution of precipitation intensities – events which are, by definition, much more common –

67 are often overlooked. While extreme precipitation events can produce outsized damages given  
68 their exceptional nature, changes in non-extreme precipitation have critical impacts on many  
69 Earth systems, including agriculture (Shortridge, 2019), infrastructure (Cook et al., 2019), and  
70 natural hazards (Dinis et al., 2021; Cannon et al., 2008). Here, to more comprehensively  
71 characterize daily precipitation shifts, we explore changes in the full distribution of wet day  
72 precipitation intensities over 17 climatically-distinct regions across the United States.

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#### 74 *1.1. Why is precipitation changing?*

75 Globally, mean annual precipitation is expected to increase  $\sim 2\%/K$  with warming (Trenberth,  
76 2003; Held and Soden, 2006; Wentz 2007; Wood et al., 2021), though considerable observed and  
77 projected regional and temporal variability underlie these projections (e.g., Polade et al. (2014)  
78 globally; Caloiero et al. (2018) in Europe). Changes in precipitation due to anthropogenic  
79 climate change is driven by combinations of thermodynamic and dynamic components. The  
80 thermodynamic component is caused by increases in atmospheric moisture content with  
81 atmospheric warming, which occurs at a rate of  $\sim 6\text{-}7\%/K$  as described by the Clausius-  
82 Clapeyron equation. This increase in atmospheric moisture availability leads to an increase in  
83 globally averaged rainfall, though estimates of the magnitude of increase in precipitation vary  
84 depending on spatial and temporal scales (Westra, 2014; Cannon and Innocenti, 2019; Sun et al.,  
85 2021; Wood and Ludwig, 2000; Wood et al., 2021; Bador et al., 2018, Giorgi et al., 2019). Locally,  
86 the rate of increase of precipitation for smaller-scale events parallels and can even exceed  
87 Clausius-Clapeyron scaling, particularly during convective events (Lenderink and van  
88 Meijgaard, 2008; Guerreiro et al., 2018; Risser and Wehner, 2017) or where local conditions shift

89 from favoring stratiform to convective precipitation (Berg and Haerter, 2013; Berg et al., 2013;  
90 Ivancic and Shaw, 2016). Generally, the Clausius-Clapeyron relationship leads to increases in  
91 extreme precipitation frequency and intensity with rising temperatures in moist, energy-limited  
92 environments, but abrupt decreases in dry, moisture-limited environments (Prein et al., 2017).  
93 Conversely, the dynamic component of precipitation changes is composed of shifts in  
94 atmospheric circulation (e.g., Swain et al, 2016) . Examples of these mechanisms include shifts in  
95 the climatologies of cyclone and anticyclone tracks, baroclinic zones, and jets, which are all  
96 driven by the reduction in the equator-pole temperature gradient, a poleward expansion of the  
97 descending branch of Hadley cells, and increases in land-sea temperature gradients (Polade et  
98 al., 2014). Increased precipitation totals can also be caused by more subtle changes, such as  
99 reductions in storm speeds (Kahraman et al., 2021) and weakening landfalling tropical cyclones  
100 (Touma et al, 2019). The relative importance of these factors varies widely depending on  
101 location.

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### 103 *1.2 How is daily precipitation variability changing?*

104 Increases in the frequency and intensity of extreme daily precipitation have been widely  
105 observed around the globe (Westra et al., 2014; Donat et al., 2016; Asadieh and Krakauer, 2015;  
106 Sun et al., 2021; Wood et al., 2021; Alexander et al., 2006; Myrhe et al., 2019) and generally agree  
107 with increases projected by climate model simulations (Moustakis et al., 2021; Toreti et al., 2013;  
108 Groisman et al., 2005; Fischer and Knutti, 2014; Fischer and Knutti, 2016; Myrhe et al., 2019; Min  
109 et al., 2011; O’Gorman, 2015). For example, Lehmann et al. (2015) found that record-breaking  
110 rainfall events occurred 12% more often than expected from 1981-2010 with an estimated 26%

111 chance that a record-setting rainfall event is due to long-term climate change. Min et al. (2011)  
112 examined observed and modeled changes and found that climate change has contributed to the  
113 observed intensification of heavy precipitation events found over about two-thirds of the  
114 Northern Hemisphere. Sub-daily extreme precipitation is both observed and projected to  
115 increase at an even faster rate than daily extremes (Prein et al., 2017; Lenderink and van  
116 Meijgaard, 2008; Westra et al., 2014).

117         Despite widespread research into precipitation extremes, changes over the full  
118 distribution of precipitation intensities are less well-characterized. For instance, Chou et al.  
119 (2012) find an increase in heavy precipitation events relative to light in the global tropics in  
120 model simulations, and Giorgi et al. (2019) find similar results over global extratropical land,  
121 including an overall reduction in lower intensity event frequency and increase in higher  
122 intensity event frequency. Hennessy et al. (1997) modeled changes in daily precipitation and  
123 found distribution shifts from low to high intensity at high latitudes along with increased  
124 heavier precipitation events coincident with a reduction of moderate events in the mid-  
125 latitudes. Despite the identification of changes in precipitation intensity distributions at broad  
126 global or zonal scales, studies at regional and local scales are sparse.

127         In the United States, increases in mean annual precipitation and extreme precipitation  
128 have been observed, though changes are not uniform (Easterling et al., 2017). Increases in heavy  
129 to extreme precipitation are well established in the central and eastern portions of the country  
130 (Groisman et al., 2012; Sun et al., 2021; Kunkel et al., 2013; Guilbert et al., 2015; Karl and Knight,  
131 1998; Pryor et al., 2008; Groisman et al., 2001; Villarini et al., 2013; Contractor et al., 2021;  
132 Groisman et al., 2005). In addition, increases in light to moderate precipitation frequency are

133 driving a general increase in precipitation frequency in the U.S. (Pal et al., 2013; Goodwell and  
134 Kumar, 2019; Karl and Knight, 1998; Roque and Kumar, 2021). However, the evolution of the  
135 proportion of lower vs higher intensity wet days is less understood with contradictory findings  
136 reported. For example, Groisman et al. (2012) focused on the central U.S. and found that  
137 moderate precipitation has become less frequent compared to heavy and extreme events, with  
138 higher intensity events increasing in frequency and moderate intensity events remaining  
139 unchanged. In contrast, Karl and Knight (1998) found an increasing frequency of events across  
140 most percentiles and U.S. regions, including an increase in moderate intensity events. While  
141 findings focused on the eastern and central U.S. are relatively consistent, studies focused on the  
142 western U.S. disagree and often unveil changes of a different sign. For example, Contractor et  
143 al. (2021) and Higgins and Kousky (2013) find generally increasing frequency and intensity of  
144 wet day events over the majority of the U.S. but decreasing moderate to heavy intensity events  
145 along the Pacific coast. This is inconsistent with findings of increasing or insignificant extreme  
146 precipitation on the U.S. west coast by Kunkel et al. (2013). Many previous analyses used  
147 gridded precipitation products (e.g., Contractor et al., 2021) that possess known inconsistencies  
148 across products (Alexander et al., 2020) and center on heavy-to-extreme precipitation or  
149 arbitrary light or moderate thresholds (e.g., 50th percentile or 10mm; Higgins and Kousky,  
150 2013). To overcome methodological limitations and reconcile disparate findings, we examine  
151 changes over the complete distribution of precipitation intensities by spatially aggregating a  
152 large number of *in-situ* station observations across a high number of empirically determined,  
153 disparate local U.S. climate regimes.

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## 156 2. Methods

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158 Our analysis uses daily *in-situ* observations of precipitation from the Global Historical  
159 Climatology Network Daily (GHCN-D). The GHCN-D database is compiled by NOAA's  
160 National Centers for Environmental Information and consists of records from over 80,000  
161 stations and 180 countries and territories, including the most complete collection of U.S. daily  
162 data available (Menne et al., 2012). Observations in GHCN-D have a sensitivity of 0.1 mm and  
163 undergo a series of nineteen quality control tests to flag duplicate data, climatological outliers,  
164 and other inconsistencies as detailed in Durre et al. (2010).

165 We filter available U.S. station records to curate a set of station observations of sufficient  
166 length and completeness for trend analysis. To do so, we impose a minimum length  
167 requirement of 50 years of complete data, where a complete year is defined as containing 95%  
168 or more of all available daily records after removal of any flagged entries. These thresholds are  
169 in general agreement with similar analyses (e.g., Anderson et al., 2015). In addition, a station  
170 record must consist of 90% or more complete data-years for inclusion. Applying these filters  
171 reduces the number of U.S. records available from an initial 63,571 to 934 that are potentially  
172 suitable for our analysis. Supporting information Figure S1 depicts station locations and  
173 additional summary statistics.

174 Our analysis focuses on regional changes in precipitation intensity distributions. To  
175 partition the U.S. into climatologically-distinct regions, we adopt the National Ecological  
176 Observatory Network (NEON) domains. These twenty domains were designed to be

177 climatically homogeneous within-domains but distinct across-domains and were created using  
178 a multivariate geographic clustering analysis incorporating nine different temperature and  
179 precipitation variables (National Ecological Observatory Network, 2022; Schimel et al., 2011;  
180 Keller et al., 2008). We center our analysis on the seventeen domains that exist entirely or  
181 predominantly within the contiguous United States (Figure S1). Rather than analyze GHCN  
182 stations individually, we employ spatial aggregation to provide a larger sample size and better  
183 view of change over time given the inherent limitations of individual station statistics and  
184 internal climate variability. Spatial aggregation has frequently been employed in precipitation  
185 analyses (e.g., Fischer et al., 2013; Groisman et al., 2005; Kunkel et al., 2013).

186         To examine changes in the distribution of wet day precipitation intensities, we aggregate  
187 all wet day precipitation observations for all stations within each NEON domain, where a wet  
188 day is defined as a station-day observing 1 mm or more of precipitation. This is done for two  
189 thirty-year periods, 1951-1980 and 1991-2020, for all seventeen NEON domains under  
190 consideration. We choose the early time period (1951-1980) due to the proliferation of GHCN-D  
191 stations that peaked in this interval (see Fig. 3b, Menne et al., 2012); we selected the late time  
192 period (1991-2020) as it is the most recent 30-year interval with available data. The distributions  
193 are built around 30-year periods of reference to align with World Meteorological Organization  
194 guidelines for climate (World Meteorological Organization, 2017) and overcome known impacts  
195 of interannual modes of climate variability (e.g., Groisman et al., 2012). Data from a station is  
196 included if the station record spans the entirety of both the early and late periods, in addition to  
197 meeting the aforementioned quality control measures. Though this definition results in the  
198 exclusion of some stations that may otherwise have been included in either the early or late

199 period but not the other, we err on the side of consistency, though results based on analysis that  
200 include stations that covered only the early or late period are included in the supporting  
201 information (Figure S2). The majority of stations span both periods (546, 58%), with 23% (161)  
202 and 18% (116) of stations covering either the early or late period, respectively. Each wet day  
203 observation is then aggregated into a corresponding early or late period daily precipitation  
204 intensity probability distribution.

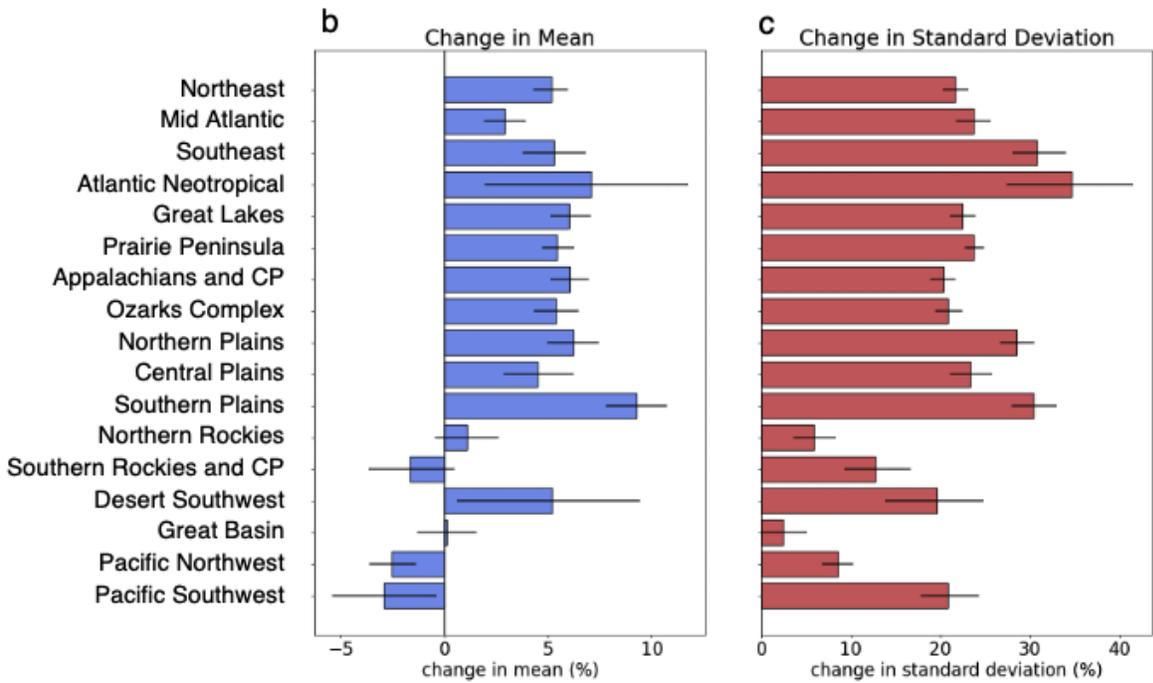
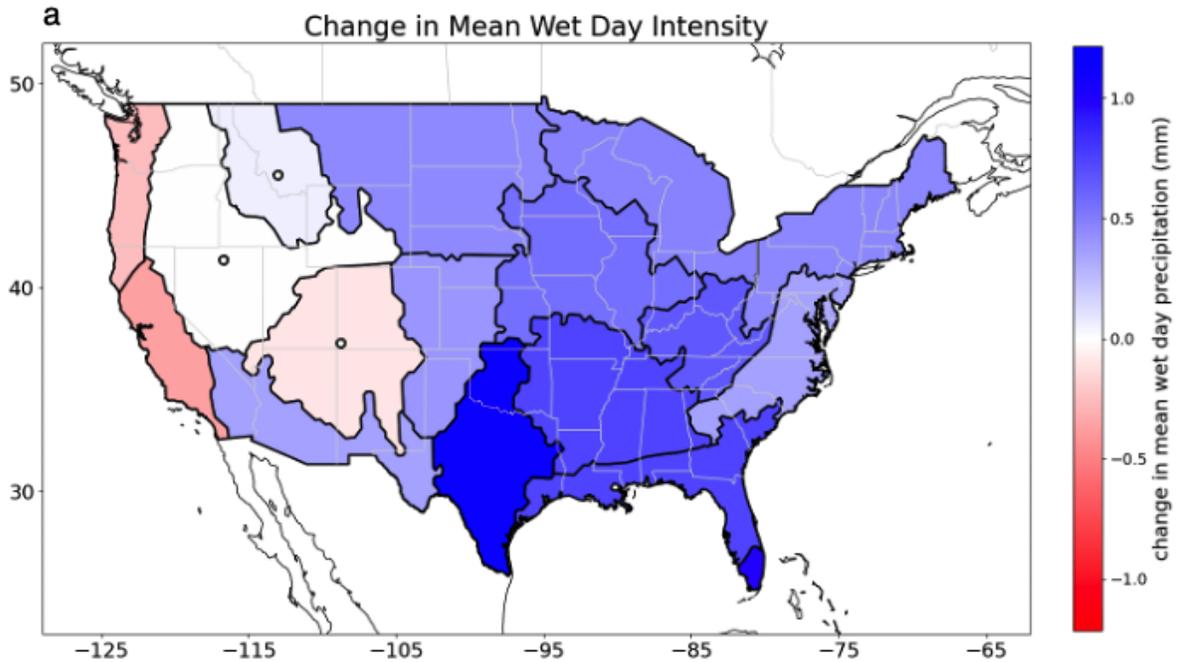
205         The resultant probability distribution functions are then directly compared through two-  
206 sample Kolmogorov-Smirnov and Anderson-Darling tests, both of which are suitable for  
207 nonparametric analysis and are not sensitive to the number of events in the distributions  
208 (Chakravarti et al., 1967; Stephens, 1974). Despite the similarity of the two tests and the more  
209 common use of the Kolmogorov-Smirnov test in hydrometeorological analyses, we employed  
210 the Anderson-Darling test as well due to its higher sensitivity to extreme values, though results  
211 ultimately proved largely consistent regardless of test. While the two tests are able to determine  
212 if distributions are distinct, the Kolmogorov-Smirnov and Anderson-Darling tests do not  
213 provide descriptive information as to how the distributions differ. We thus also examine the  
214 statistical moments of each distribution (mean, standard deviation, skew, kurtosis) and employ  
215 a bootstrapping methodology to assess statistical confidence intervals. Our bootstrapping  
216 process resamples each distribution with replacement to create new distributions of the same  
217 size as the original distributions. This process is replicated 1,000 times and differences in each  
218 statistical moment between the distributions are recalculated for each resample to produce  
219 confidence intervals for each statistical moment change.

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### 3. Results

Early and late period wet day precipitation intensity distributions are statistically significantly different for all but three NEON domains (Figure 1a; Northern Rockies, Great Basin, and Southern Rockies and Colorado Plateau) with broadly consistent changes observed in central and eastern domains. Specifically, mean wet day precipitation increases in all domains east or the Rocky Mountains (Figure 1a), with an intensification in mean wet day precipitation between 4.6-7.1% for all but two of these domains (Figure 1b). Similarly, the standard deviation of wet day precipitation intensity increased between 20-31% for each domain in this grouping (Figure 1c). Changes are less consistent for western domains, with an increase in mean wet day precipitation in one domain (Desert Southwest), a decrease in mean wet day precipitation in two domains (Pacific Northwest and Pacific Southwest), and no statistically significant change in the remaining domains. Despite the inconsistencies, the standard deviation of wet day precipitation intensity increased in all western domains, though the magnitude of increases vary more widely than in the eastern U.S. Supporting information Table S1 shows the differences in mean, standard deviation, skew, and kurtosis across all domains.



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240 *Figure 1: Changes in Wet Day Precipitation Intensity. (a) Map of changes in mean wet day precipitation*

241 *for NEON domains. Red-blue fill indicates change in precipitation intensity (mm/day) within domains*

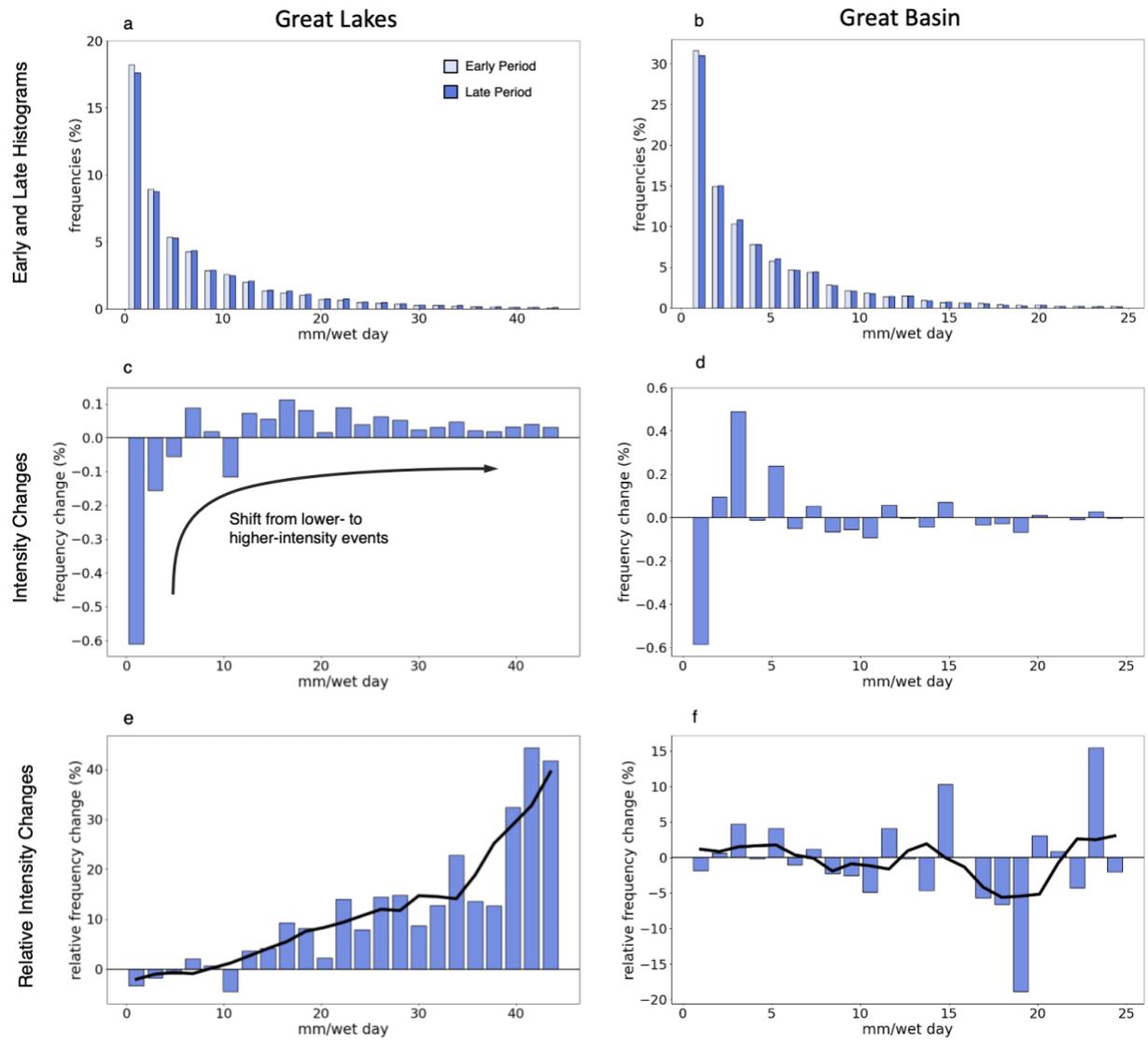
242 *(dark grey borders) on top of state boundaries (light grey borders). Grey circle with white fill denotes*

243 *domains without a statistically significant change. (b) Changes in mean wet day precipitation for NEON*  
244 *domains. Blue bars show percentage change of mean and horizontal black line shows 95% confidence*  
245 *interval. (c) Same as (b) but with red bars and standard deviation of wet day precipitation.*

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247 In addition to changes in mean and standard deviation, we also quantified shifts in the  
248 underlying distributions between the early and late periods, allowing for a more nuanced  
249 characterization of observed distribution changes (Figures 2, 3). Figure 2 illustrates how the  
250 precipitation intensity distribution changes between the early and late periods for two example  
251 domains. We characterize absolute differences in wet day intensities in Figures 3b and 3e along  
252 with relative differences in Figures 2c and 2f. For example, in Figure 2b, we demonstrate that  
253 the Great Lakes domain has experienced a robust shift from lower to higher precipitation  
254 intensities across the full distribution of intensities, which becomes clearer when compared  
255 against the initial frequencies in the early period (Figure 2c). To illustrate, the likelihood of 40  
256 mm or greater events in the Great Lakes domain is roughly 30-40% greater in the later period of  
257 observation. Changes over the Great Lakes domain contrast with the lack of a consistent  
258 response in the Great Basin domain (Figures 2d-f). The observed shift from lower- to higher-  
259 intensity wet day totals in the Great Lakes domain is broadly consistent with findings across the  
260 central and eastern U.S. (see blue filled regions, Figure 3).

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264 *Figure 2: Change in Precipitation Intensity between Early and Late Periods. (a) Histograms of wet day*

265 *precipitation intensity in the Great Lakes domain for the early (light blue; 1951-1980) and late (dark blue;*

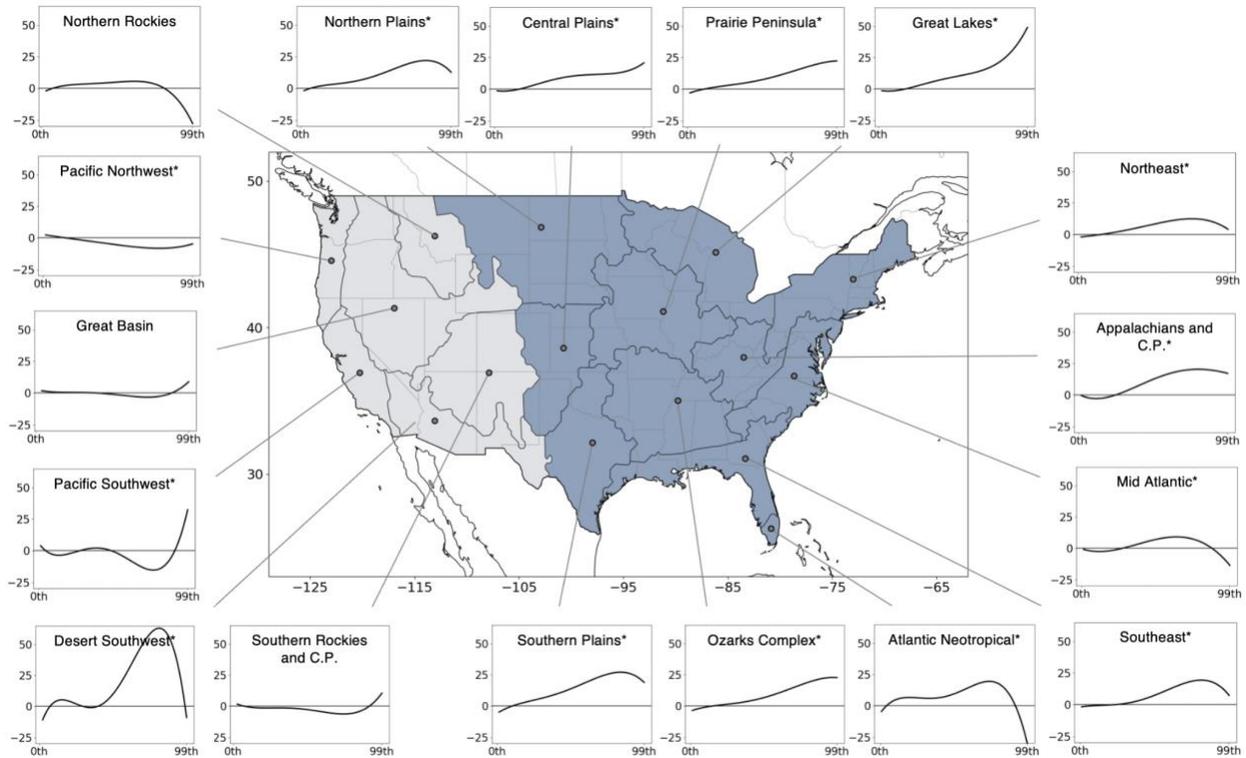
266 *1991-2020) periods for sub-99th percentile daily precipitation totals. Histogram values represent the*

267 *percentage of all wet-day events within the binned intensity. (b) Absolute difference in wet day*

268 *precipitation intensity frequency between the late and early periods for the Great Lakes NEON domain*

269 *for sub-99th percentile daily precipitation events. (c) Same as (b) but the change has been normalized by*

270 *dividing by the early period frequency. Thick black line denotes average change of the five nearest bins*  
 271 *centered on a given bin. (d-f) Same as (a-c) but for the Great Basin domain. This figure is replicated for all*  
 272 *domains in supporting information Figures S3-S17.*  
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274  
 275 *Figure 3: Smoothed Relativized Frequency Change for Each Domain. (map) The United States with*  
 276 *NEON domain boundaries (thick dark grey) and state borders (thin light grey). Blue fill denotes the*  
 277 *cluster of central and eastern domains with a predominantly consistent change in frequency across*  
 278 *intensities. Conversely, grey fill denotes the cluster of western domains with inconsistent changes in*  
 279 *frequency across intensities. (domain subplots) Smoothed change in frequency of intensity for each*  
 280 *domain. See Figures 2c, 2f for demonstration of underlying data and calculations. Displayed change is for*

281 *the 0th-99th percentile of wet day intensities. Asterisked name indicates domains with a statistically*  
282 *significant change in their distributions.*

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284         The shift from lower- to higher-intensity events is largely consistent in the central and  
285 eastern U.S., with lower-intensity events decreasing in relative frequency for all eleven domains  
286 (blue filled regions, Figure 3). However, while higher-intensity events generally increase for all  
287 central and eastern domains and intensities, this change is not uniform, with a reduction in the  
288 most intense events in the Atlantic Neotropical and Mid Atlantic domains. Similarly, while  
289 across-domain responses converge on an approximately linear increase peaking at an ~25%  
290 increase in the relative frequency of highest intensity events, there is a lesser increase in the  
291 relative frequency of highest intensity events compared to heavy precipitation events in several  
292 domains (e.g., the Northeast domain). Similar to the mixed responses in wet day mean  
293 precipitation changes, changes across distribution frequencies vary between domains in the  
294 western U.S. (see grey filled regions, Figure 3). For example, the greatest increase in higher-  
295 intensity, non-extreme event frequency of all domains is in the Desert Southwest. This change is  
296 juxtaposed against nearby regions such as the Pacific Southwest, Southern Rockies and  
297 Colorado Plateau, Great Basin, and Pacific Northwest, with generally consistent shifts  
298 consisting of an increase in lowest and highest-intensity events but decreases in moderate and  
299 heavy precipitation events. It should be noted that the muted changes within the Great Basin,  
300 Southern Rockies and Colorado Plateau, and Northern Rockies domains are not statistically  
301 significant.

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304 **4. Discussion**

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306 Here, we examine the full extent of wet day precipitation intensity distributions and reveal  
307 statistically robust changes throughout the United States. Broadly, our analysis reveals an  
308 increase in mean wet day precipitation from 1951-1980 to 1991-2020 driven by a shift from lower  
309 to higher intensity wet day events in the central and eastern U.S. Changes in mean wet day  
310 precipitation and underlying wet day intensity distribution shifts are varied in the western U.S.,  
311 though standard deviation of wet day precipitation totals increases for all domains investigated.  
312 This combination of changes leads to a statistically significant change in the underlying wet day  
313 precipitation intensity distributions for all but three of the domains analyzed across the U.S.  
314 (Northern Rockies, Southern Rockies and Colorado Plateau, and Great Basin). In summary, our  
315 analysis generally reveals observed increases in both wet day precipitation intensity and  
316 intensity variability.

317 While this work does not assess mechanistic drivers of observed changes, we note that  
318 the consistency of the increase in mean wet day precipitation across an area as broad as the  
319 eastern two-thirds of the U.S. (Figures 1, 3) is consistent with an underlying thermodynamic  
320 shift. The varying response along the U.S. west coast does not clearly align with thermodynamic  
321 explanations and we do not attempt to further explain underlying mechanisms here. However,  
322 we do note earlier work by Pfahl et al. (2017) found an omnipresent increase in thermodynamic  
323 influences across the country, but an east-west dichotomy in dynamic influences with dynamic

324 mechanisms enhancing precipitation in the eastern U.S. and diminishing precipitation along the  
325 Pacific coast.

326         Though existing literature largely focuses on heavy-to-extreme precipitation or arbitrary  
327 light or moderate thresholds, our findings largely coincide and complement earlier findings  
328 such as the east-west division of changes in extreme precipitation described in the Fourth  
329 National Climate Assessment (Easterling et al., 2017). The relative increases in moderate and  
330 heavy precipitation in the eastern U.S. mirrors well-established increases in precipitation  
331 extremes, as well as annual precipitation, over the central and northeastern portions of the  
332 country (e.g., Groisman et al., 2012). We highlight the high consistency in the shift in  
333 precipitation intensities across the precipitation intensity distributions in this area (Figure 3) as  
334 well as the rising mean (~4.6-7.1%) and standard deviation (20-31%) of wet day precipitation.  
335 The inconsistent pattern of results for the western U.S. mirrors earlier results as well (Contractor  
336 et al., 2021; Higgins and Kousky, 2013; Rosenberg et al., 2010), though our analysis builds off  
337 earlier work by using a large number of *in situ* measurements instead of gridded observational  
338 products or studies that focus on a limited number of stations.

339         There are some potential limitations of this study, beginning with the underlying  
340 assumption that NEON domains are internally consistent. While NEON domains are  
341 empirically designed to possess internally homogeneous climates, there exists some measure of  
342 variability within domains and inconsistent station availability may impact domain-level  
343 findings. Further, we find slight differences in domain-level changes if all qualifying stations  
344 are included in the analysis as opposed to only stations which span both the early and late  
345 periods, though the vast majority of changes remain the same regardless of this methodological

346 decision (see supporting information Figure S2 for findings with all stations). While not a  
347 limitation of our work, it should be explicitly noted that our analysis focuses on changes in *wet*  
348 *day* precipitation intensity and, therefore, does not consider underlying changes in precipitation  
349 frequency. This distinction is important for considering the impacts of these findings in the  
350 scope of annual precipitation totals, for instance. Finally, although we examine trends in  
351 precipitation through a period of time of increasing greenhouse gas emissions and resultant  
352 climate impacts, the analysis presented here is insufficient to directly attribute changes to  
353 ongoing anthropogenic climate change, although our findings are largely consistent with  
354 expected changes (e.g. Pfahl et al., 2017) and such an analysis could be performed using a  
355 robust attribution methodology (e.g., Diffenbaugh et al, 2017) .

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## 358 5. Conclusion

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360 We use curated daily *in situ* precipitation measurements from the GHCN to examine regional  
361 trends in wet day precipitation distributions from 1951-1980 to 1991-2020. We reveal significant  
362 changes in wet day intensity distributions for fourteen of seventeen domains around the United  
363 States, including all domains in the central and eastern U.S. These nearly ubiquitous changes  
364 are driven by a general shift from lower to higher intensity wet day precipitation totals around  
365 the U.S. and are largely manifested as increases in the mean wet day precipitation intensity and  
366 in the standard deviation of wet day precipitation intensity, though we identify an east-west  
367 split. Our findings can help inform the understanding of how natural hazard risk has changed

368 over time, such as how shifting rates of moderate and heavy intensity precipitation may be  
369 interacting with increased wildfire burn areas in the western U.S. to affect water resources and  
370 landslide risk (Williams et al., 2022). These results also have important implications for research  
371 and applications surrounding agricultural yields and infrastructure design. Additionally, these  
372 results can be compared with climate model output to examine the abilities of climate models to  
373 reproduce the spatial patterns in changes in daily precipitation intensities.

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### 376 **Acknowledgements**

377 We would like to acknowledge and thank the National Centers for Environmental Information  
378 for the use of their publicly available Global Historical Climatology Network Daily dataset. We  
379 also express our gratitude for the Ubben Program for Carbon and Climate Science at  
380 Northwestern University for supporting and facilitating this work through a postdoctoral  
381 fellowship to Ryan D. Harp. Finally, this research was supported in part through the  
382 computational resources and staff contributions provided for the Quest high performance  
383 computing facility at Northwestern University, which is jointly supported by the Office of the  
384 Provost, the Office for Research, and Northwestern University Information Technology.

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### 387 **Open Research and Availability Statement**

388 Global Historical Climatology Network Daily data is publicly available through the National  
389 Centers for Environmental Information at <https://www.ncei.noaa.gov/products/land-based->

390 [station/global-historical-climatology-network-daily](#). Code developed by the authors to conduct  
391 the analysis and produce the figures within this study is available at  
392 [https://github.com/ryandharp/Observed Changes in Daily Precipitation Intensity in the Uni](https://github.com/ryandharp/Observed_Changes_in_Daily_Precipitation_Intensity_in_the_United_States)  
393 [ted\\_States](#). This code will be archived on Zenodo upon completion of the peer review process,  
394 at which time the finalized link to archive, DOI, and data citation will be added to this  
395 statement.

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