

**Suppression of nitrogen deposition on global forest soil CH<sub>4</sub> uptake depends on nitrogen status**

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## **Introduction**

The uploaded Data Set S1 (CH<sub>4</sub>\_exp dataset in main text) was used to derive the response factors of soil CH<sub>4</sub> flux to N input in global forests; Data Set S2 (CH<sub>4</sub>\_obs dataset in main text) was used to estimate the soil CH<sub>4</sub> fluxes in global forests; Data Sets S3–S7 were used to classify the N-limited and N-saturated forests on global level; Data Set S8 contains environmental factors (MAT, MAP, soil texture, etc.) for global estimations; Data Set S9 contains global forest soil CH<sub>4</sub> budgets reported in previous studies. The data analysis process and produced figures can be replicated with the uploaded R script (Code S1).

**Text S1. Nitrogen saturation status of global forests indicated by sensitivity of soil N<sub>2</sub>O emission to N deposition.**

Globally, human-induced increase in atmospheric N deposition is changing forests from a nitrogen-limited to nitrogen-saturated status. In N-limited forests, plants and microbes utilize N conservatively for a lower proportion of input N to be leaked from tight N cycling processes (Chapman et al., 2006; Van Der Heijden et al., 2008). However, when forests become N-saturated, input N exceeds the N demand of plants and microbes, leading to excessive utilization of N, and thus, the N cycle becomes more open (Hietz et al., 2011). Therefore, a higher proportion of input N is lost via leaching or gaseous emission (Aber et al., 1989). This implies that increased gaseous N emissions (N<sub>2</sub>O, NO, and N<sub>2</sub>) per unit of N deposition (i.e., higher sensitivity of gaseous N emissions to N deposition) may indicate forests reaching N saturation (Aber et al., 1998). Coincidentally, studies have measured nitrous oxide (N<sub>2</sub>O) greenhouse gas emissions under different N input levels since the 1980s in global forests, using a controlled experiment design and standard sampling method (Holland et al., 1999). The accumulated experimental data provide an opportunity to quantify the sensitivity of N<sub>2</sub>O emissions to N deposition in various forests, and indicate the N limitation or saturation status of global forests.

*Gathering data*

To quantify the sensitivity of soil N<sub>2</sub>O emissions to N deposition ( $s_N$ ), we compiled soil N<sub>2</sub>O emission data observed in N addition experiments conducted in global forests. On 03/30/2022, we searched for papers and theses published before 01/01/2022 from the Web of Science Core Collection database ([www.webofscience.com](http://www.webofscience.com)) and China National Knowledge Infrastructure Theses and Dissertations Database (<https://oversea.cnki.net/kns?dbcode=CDMD>), using the following keywords: "forest" AND "greenhouse gas" OR "N<sub>2</sub>O" OR "nitrous oxide". The retrieved 7422 papers and 718 theses were then refined manually based on the following criteria: (i) experimental N addition was conducted in forest ecosystem; (ii) literature recorded the location, time, and dose of the experiment(s); (iii) soil N<sub>2</sub>O flux was observed in experimental sites and measured using gas chromatograph technique (Holland et al., 1999). As a result, the compiled "N<sub>2</sub>O\_exp" dataset (Data Set S3) contained 553 observations from 102 sites worldwide (Fig. S7).

Similarly, we compiled data on the soil N<sub>2</sub>O emission rates of global forests observed under natural conditions. We refined from the same papers and theses as above, using a different set of criteria: (i) no nutrients, including N, were artificially added to the forest site so the site only received naturally deposited N; (ii) literature recorded the location, and time of flux measurement; (iii) soil N<sub>2</sub>O flux was observed in the field and measured using gas chromatograph technique (Holland et al., 1999). The compiled "N<sub>2</sub>O\_obs" dataset (Data Set S4) contained 246 observations from 140 sites worldwide (Fig. S7).

In addition, we compiled data on total N loss (N leaching and gaseous N emission combined), N leaching, and change in soil N pool, from N addition experiments in global forests. We searched in the aforementioned databases using the following keywords: "forest" AND "nitrogen addition" OR "fert\*" AND "nitrogen loss" OR "nitrogen leaching" OR "nitrogen budget". Retrieved 2693 papers and theses were then refined based on the following criteria: (i) literature recorded the location, time, and dose of experimental N addition in forests; (ii) total N loss rate, N leaching rate, or change rate of soil N pool was observed or estimated in the

experiments. The compiled "Ncycle\_exp" dataset (Data Set S5) contained 169 observations from 37 sites (Fig. S7).

To analyze the relationship between  $s_N$  and N saturation status, we compiled data on field-observed N-limited and N-saturated forests indicated by N leaching. On 10/31/2022, we searched for literature in the aforementioned databases using the following keywords: "forest" AND "leaching" AND "nitrogen limit\*" OR "nitrogen saturat\*". Retrieved 823 papers and theses were then refined based on the following criteria: (i) literature recorded whether the forest was N-limited or N-saturated, and its location; (ii) literature used nitrogen leaching as an indicator of N limitation or saturation status. The compiled "Nleach" dataset (Data Set S6) contains 136 observations from 92 sites worldwide (Fig. S6). We also used data on field-observed N-limited and N-saturated ecosystems indicated by plant growth response to N input ("NuLi" dataset; Data Set S7) from a published database by Du et al. (2020). It covers 106 sites worldwide, 65 of which are forest sites (Fig. S6).

Moreover, we extracted auxiliary information from the literature on environmental factors (including mean annual temperature, MAT; mean annual precipitation, MAP; mean annual N deposition rate,  $N_{depo}$ ; coefficients of temporal variation, MAT.cv, MAP.cv, and  $N_{depo}.cv$ ; soil sand content, soil clay content, and other soil properties) for the forest sites in the datasets. However, the literature did not provide the necessary auxiliary information for all sites; therefore, spatial datasets were used to fill in the missing data based on the location of the sites. Global temperature and precipitation datasets were obtained from the Climatic Research Unit, University of East Anglia ([https://crudata.uea.ac.uk/cru/data/hrg/cru\\_ts\\_4.03/](https://crudata.uea.ac.uk/cru/data/hrg/cru_ts_4.03/)). The soil C:N ratio was obtained from a published database (Shangguan et al., 2014). Other soil properties were obtained from the HWSD dataset (<https://www.fao.org/soils-portal/data-hub/soil-maps-and-databases/harmonized-world-soil-database-v12/en/>). N deposition rate and forest cover data were from published databases (Ackerman et al., 2019; Liu et al., 2020). The forest biome map was derived from the Global Forest Monitoring project (Hansen et al., 2010).

#### *Quantifying the sensitivity of soil N<sub>2</sub>O emissions to N deposition*

Under low N input, the soil N<sub>2</sub>O emission rate responds almost linearly to N input, whereas high N input may induce non-linear responses (Aber et al., 1998; D.-G. Kim et al., 2013). High N input may change ecosystem properties, leading to a deviation from the natural response of ecosystems to environmental change. Therefore, we used a linear model (Eq. S1) to define and quantify the sensitivity ( $s_N$ ) of soil N<sub>2</sub>O emissions to N deposition (or low N input), for  $s_N$  to reflect ecosystem properties (i.e., N saturation status).

$$R_{N_2O} = s_N \times N_{depo} + R_0 \quad (\text{Eq. S1})$$

where  $R_{N_2O}$  is the soil N<sub>2</sub>O emission rate (kgN<sub>2</sub>O-N ha<sup>-1</sup> yr<sup>-1</sup>),  $N_{depo}$  is the atmospheric N deposition rate (kg N ha<sup>-1</sup> yr<sup>-1</sup>),  $s_N$  is the sensitivity of soil N<sub>2</sub>O emission to N deposition, quantified as soil N<sub>2</sub>O emission per unit of low N input (kgN<sub>2</sub>O-N kgN<sup>-1</sup>), and  $R_0$  is the background soil N<sub>2</sub>O emission rate when there is no N deposition or artificial N addition (kgN<sub>2</sub>O-N ha<sup>-1</sup> yr<sup>-1</sup>).

A segmented regression analysis on N<sub>2</sub>O\_exp dataset showed that there is one change point in the linear relationship between N input rate and  $R_{N_2O}$ , which is  $174.70 \pm 19.73$  kgN ha<sup>-1</sup> yr<sup>-1</sup>. That is in line with change points estimated or used in previous studies (Bouwman et al., 2002; Hoben et al., 2011; M. Lu et al., 2022; McSwiney & Robertson, 2005; Shcherbak et al.,

2014). Conservatively, experimental data with N addition rates not exceeding  $150 \text{ kg N ha}^{-1} \text{ yr}^{-1}$  were used as “low N input” data in further analysis. The N deposition rates in global forests were lower than the level (Ackerman et al., 2019). For all the low-N input data in the  $\text{N}_2\text{O\_exp}$  dataset, we aggregated them to  $0.5^\circ \times 0.5^\circ$  grids based on their coordinates to match the spatial resolution with environmental factors and reduce random errors in sampling. A linear model (Model:  $R_{\text{N}_2\text{O}} \sim \text{N input rate}$ ) was built for each grid with low-N input data. The slope of the linear model was the estimated  $s_{\text{N}}$  of the grid (Table S3).

Based on the estimated  $s_{\text{N}}$  of all grids and the corresponding environmental factors, we built a generalized linear model to simulate  $s_{\text{N}}$  (Table S4). In addition, another generalized linear model was built to simulate  $R_0$ .

To validate  $s_{\text{N}}$ , we firstly used the modeled  $s_{\text{N}}$ , together with the modeled  $R_0$  and  $N_{\text{depo}}$  datasets, to estimate  $R_{\text{N}_2\text{O}}$  (Eq. S1). The estimated  $R_{\text{N}_2\text{O}}$  values were compared with  $R_{\text{N}_2\text{O}}$  observations ( $\text{N}_2\text{O\_obs}$  dataset) and indirectly validated the intermediate variable  $s_{\text{N}}$  (Fig. S8). In addition,  $s_{\text{N}}$  was validated using a second approach. The sensitivity of N loss to N input ( $c_1$ ), the sensitivity of N leaching to N input ( $c_2$ ), and the end-product ratio of nitrification and denitrification processes ( $c_3$ ) were either derived from the  $\text{Ncycle\_exp}$  dataset or extracted from the literature;  $s_{\text{N}}$  was then calculated from these parameters (Eq. S2).

$$s_{\text{N}} = c_3 \times (c_1 - c_2) \quad (\text{Eq. S2})$$

The limited observations allowed us to calculate  $s_{\text{N}}$  on a biome scale (Fig. S7), which was then compared with the biome-mean value of the modeled  $s_{\text{N}}$  to validate it. The good agreement also validated the modeled  $s_{\text{N}}$  ( $r = 0.998$ ).

#### *Determining N saturation status of global forests using $s_{\text{N}}$*

We tested whether  $s_{\text{N}}$  can distinguish between N-limited and N-saturated forests using data from forests having field-observed N saturation status data. First, we combined  $\text{Nleach}$  and  $\text{NuLi}$  datasets to enlarge the sample size and derive a universal classification. Excluding three duplicate sites in both datasets, the combined dataset had 154 sites with field-observed N saturation status (86 N-limited and 68 N-saturated sites).

We modeled the  $s_{\text{N}}$  of the 154 sites using environmental factors (Table S4). We then analyzed the  $s_{\text{N}}$  of N-limited and N-saturated forests and verified if there were significant differences on the global and biome scales. In Western Europe, North America, and East Asia, where there were abundant sites, we also compared the  $s_{\text{N}}$  of forests with N-limited or N-saturated status on a regional scale. The mean  $s_{\text{N}}$  was significantly different on global and regional scales ( $p < 0.001$ ; Fig. S9), proving that  $s_{\text{N}}$  can indicate N limitation or saturation status in forests.

Then we calculated an optimal threshold for  $s_{\text{N}}$  using data from 154 sites with field-observed N saturation status and  $s_{\text{N}}$  information. The bootstrap method accounted for the different sample sizes of N-limited and N-saturated sites (Davison & Hinkley, 1997). Specifically, from the 154 sites, we randomly sampled 10 N-limited and 10 N-saturated sites and selected a cutoff value for their  $s_{\text{N}}$  at a precision of  $0.0001 \text{ kgN}_2\text{O-N kgN}^{-1}$ . Sites in which  $s_{\text{N}}$  were above the cutoff value were classified as “N-saturated,” and the rest were classified as “N-limited.” The classified N saturation status of the sites was compared with field observations to determine the accuracy of the classification, which was calculated as the proportion of sites accurately classified into the same category as that observed. All possible cutoff values were tested, and the one with

the highest classification accuracy was the “optimal” cutoff value. Random sampling and detection of optimal cutoff values were repeated 5000 times, during which some optimal cutoff values were detected more frequently than others. The optimal threshold for  $s_N$  in all samples was the most frequently detected optimal cutoff value, which was  $0.0143 \text{ kgN}_2\text{O-N kgN}^{-1}$ .

The N saturation status of global forests was determined based on the optimal threshold. Forests with  $s_N$  above the threshold were classified as N-saturated, and the rest were classified as N-limited. The accuracy of the classification was higher than 70% on global and regional scales (Fig. S6). Based on the classification, we produced a rasterized map of N-limited and N-saturated forests ( $0.5^\circ \times 0.5^\circ$  resolution) in ArcGIS (ESRI, 2011).

## Text S2. Inferring the variation of methane production and oxidation rates from the variation of observed methane fluxes

Soil CH<sub>4</sub> flux observed on the soil-air interface is codetermined by methane production (methanogenesis) and oxidation rates (Eq. S3). However, it has been difficult to disentangle the responses of methane production and oxidation to N input, because of the limited ability to separately observe methanogenesis and methane oxidation processes in the field. Here, we inferred the variation of methane production and oxidation rates from the variation of observed methane fluxes. This could further support the “three stage” hypothesis we proposed.

$$R_{CH_4} = R_{CH_4_{prod}} - R_{CH_4_{oxid}} \quad (\text{Eq. S3})$$

where  $R_{CH_4}$  is the observed soil CH<sub>4</sub> flux (kg ha<sup>-1</sup> yr<sup>-1</sup>), positive  $R_{CH_4}$  value means methane emission, whereas negative  $R_{CH_4}$  value means methane uptake;  $R_{CH_4_{prod}}$  is methane production rate (kg ha<sup>-1</sup> yr<sup>-1</sup>);  $R_{CH_4_{oxid}}$  is methane oxidation rate (kg ha<sup>-1</sup> yr<sup>-1</sup>).

The change in methane production and oxidation rates could hardly be inversely calculated from the  $R_{CH_4}$  values. Here, we inferred the change in  $R_{CH_4_{prod}}$  and  $R_{CH_4_{oxid}}$  by analyzing the mean values and standard deviations of  $R_{CH_4}$ .

Firstly, the standard deviation of  $R_{CH_4}$  could be calculated from that of  $R_{CH_4_{prod}}$  and  $R_{CH_4_{oxid}}$  (not considering the interaction between  $R_{CH_4_{prod}}$  and  $R_{CH_4_{oxid}}$ ; Eq. S4).

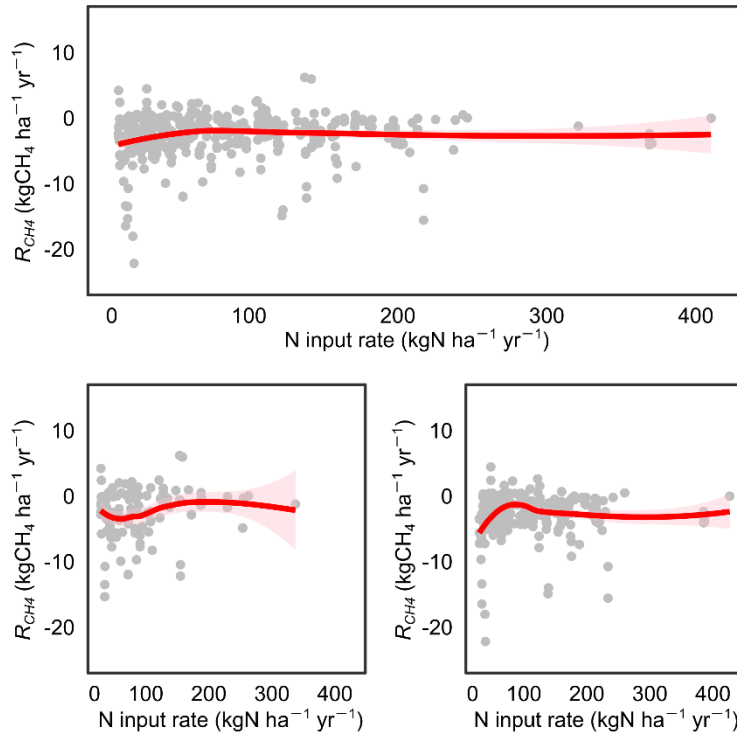
$$SD(R_{CH_4}) = \sqrt{SD(R_{CH_4_{prod}})^2 + SD(R_{CH_4_{oxid}})^2} \quad (\text{Eq. S4})$$

Usually, when the expected value of a variable becomes higher, its observations will be more dispersed. This is because the random errors in the observations are often proportional to their values. That is to say, statistical dispersion of  $R_{CH_4_{prod}}$  and  $R_{CH_4_{oxid}}$  (as indicated by their standard deviations) should be positively related to their mean values.

Therefore, the decrease in the standard deviation of  $R_{CH_4}$  under high N input (Fig. S2) may result from: (1)  $R_{CH_4_{prod}}$  decreased under high N input, and  $R_{CH_4_{oxid}}$  didn't change or slightly increased; (2)  $R_{CH_4_{oxid}}$  decreased under high N input, and  $R_{CH_4_{prod}}$  didn't change or slightly increased; (3) both  $R_{CH_4_{prod}}$  and  $R_{CH_4_{oxid}}$  decreased under high N input.

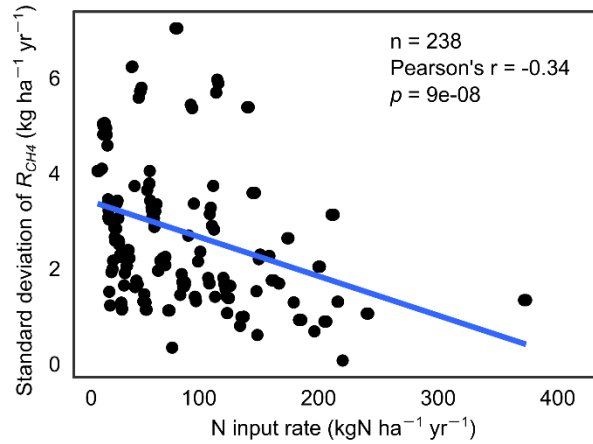
Meanwhile, we observed that the mean values of  $R_{CH_4}$  remained nearly unchanged under high N input (Fig. S1a), which may result from: (i) both  $R_{CH_4_{prod}}$  and  $R_{CH_4_{oxid}}$  increased under high N input; (ii) both  $R_{CH_4_{prod}}$  and  $R_{CH_4_{oxid}}$  decreased under high N input; (iii) both  $R_{CH_4_{prod}}$  and  $R_{CH_4_{oxid}}$  remained constant under high N input.

Combining the two evidences (standard deviation and mean values of  $R_{CH_4}$  under high N input), it can be inferred that only hypotheses (3) and (ii) can be true at the same time. That is, both  $R_{CH_4_{prod}}$  and  $R_{CH_4_{oxid}}$  decreased under high N input.

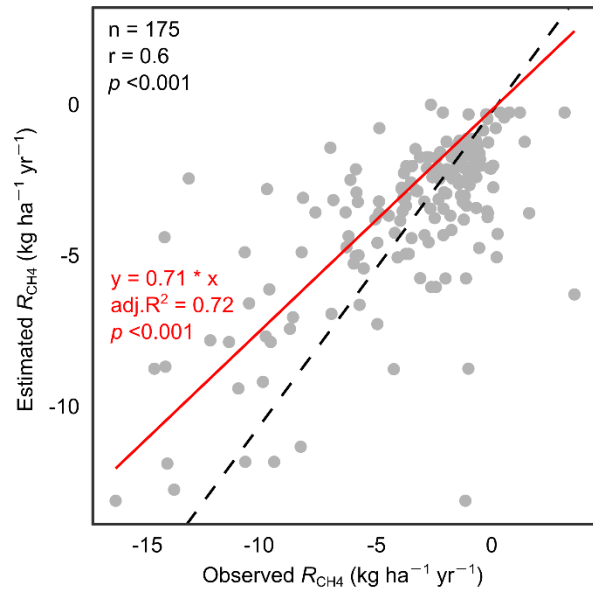


**Fig. S1.** Locally weighed regression ("LOWESS") model on soil CH<sub>4</sub> emission rate and N input rate. (a) Using all observations compiled from global N addition experiments, the N input rates of which were no greater than 400 kgN ha<sup>-1</sup> yr<sup>-1</sup> (n = 448). The few but variable observations on soil CH<sub>4</sub> fluxes at sites where N input rates were above 400 kgN ha<sup>-1</sup> yr<sup>-1</sup> (n = 17) were not used in further analysis. (b) LOWESS model constructed using data from N-limited sites and also where N addition experiments have been conducted for no more than 3 years when CH<sub>4</sub> emissions were observed (n = 131); (c) LOWESS model constructed using data from N-saturated forests, or data from sites where N addition experiments have been conducted for more than 3 years before observing the CH<sub>4</sub> fluxes (n = 317). Pink shadings represent the standard errors of the fitted models.



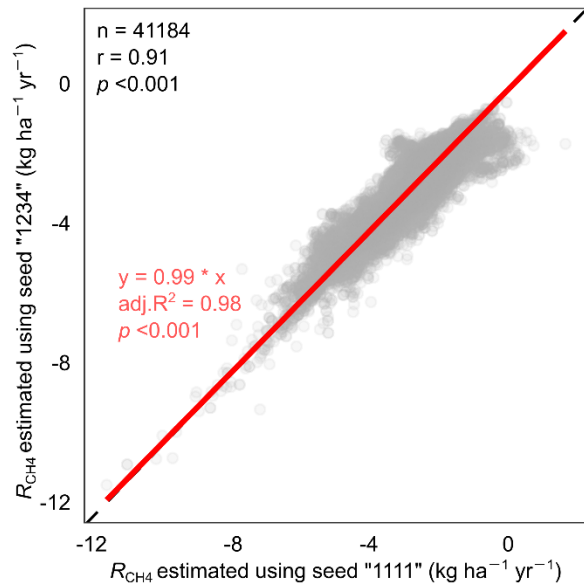


**Fig. S2.** Standard deviation of soil methane flux ( $R_{CH_4}$ ) was negatively correlated to N input rate. Data corresponding to N input levels above 400 kgN ha<sup>-1</sup> yr<sup>-1</sup> were not included in this analysis, because the very limited observations may not sufficiently reveal the statistical distribution of  $R_{CH_4}$ . There were 238 unique N input rates that was no greater than 400 kgN ha<sup>-1</sup> yr<sup>-1</sup>. In practice, standard deviation was calculated for  $R_{CH_4}$  corresponding to each N input rate, and N input rates less than 2 kgN ha<sup>-1</sup> yr<sup>-1</sup> in difference (e.g., standard deviation of  $R_{CH_4}$  corresponding to 5 kgN ha<sup>-1</sup> yr<sup>-1</sup> was calculated using observations whose N input rates were within the range of 3 to 7 kgN ha<sup>-1</sup> yr<sup>-1</sup>). That was to make sure that there were sufficient observations for each N input level.



**Fig. S3.** Comparing observed soil  $\text{CH}_4$  flux ( $R_{CH_4}$ ) in testing dataset with that estimated using averaged outputs from 1,000 random forest regression models. The red line and font indicate the fitted linear model on estimated and observed  $R_{CH_4}$  values.

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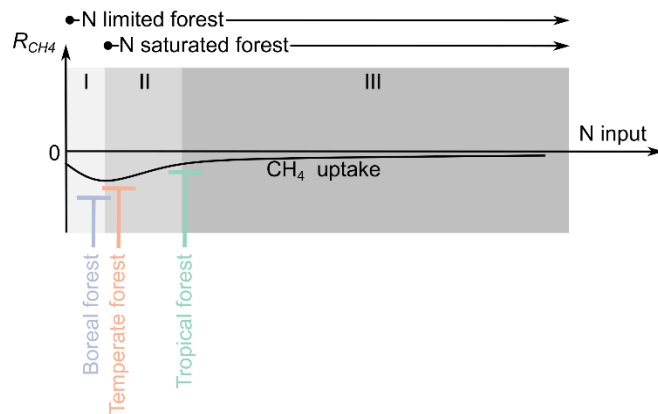
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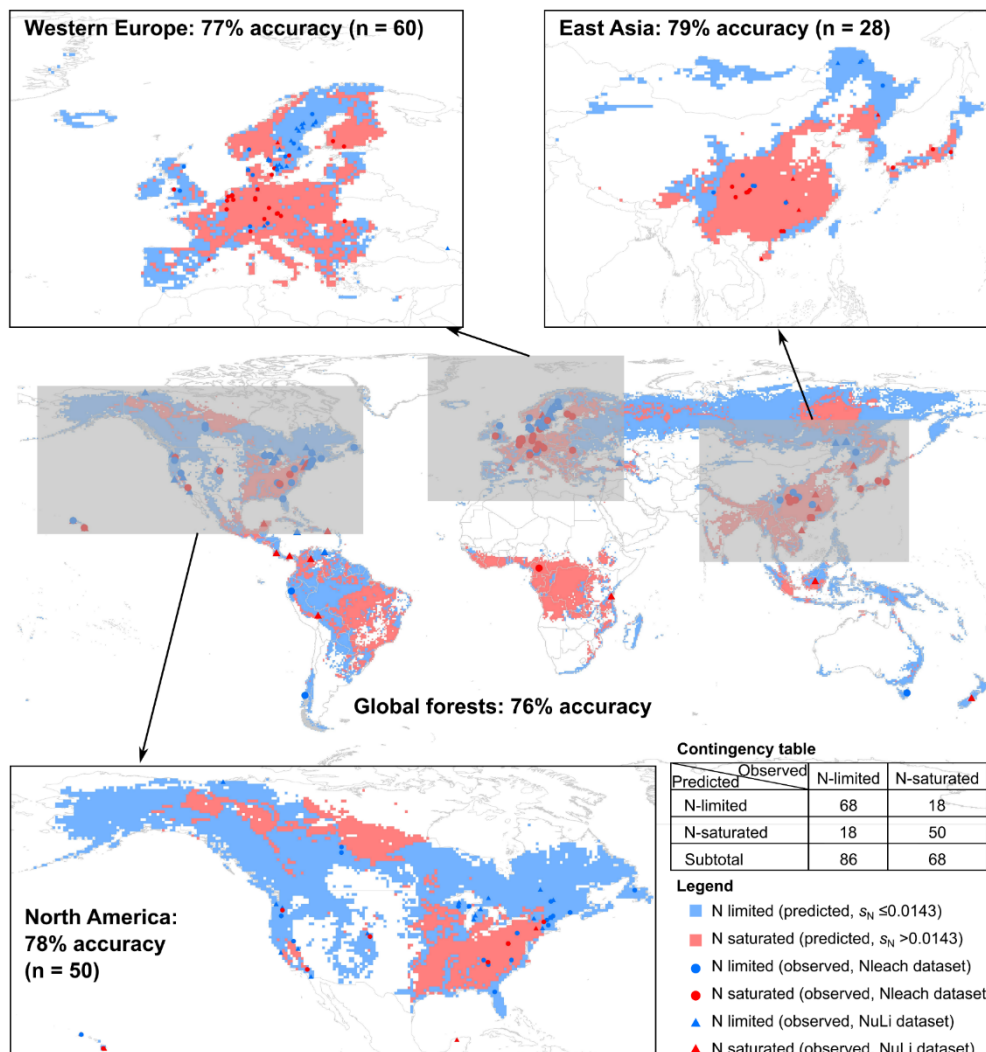
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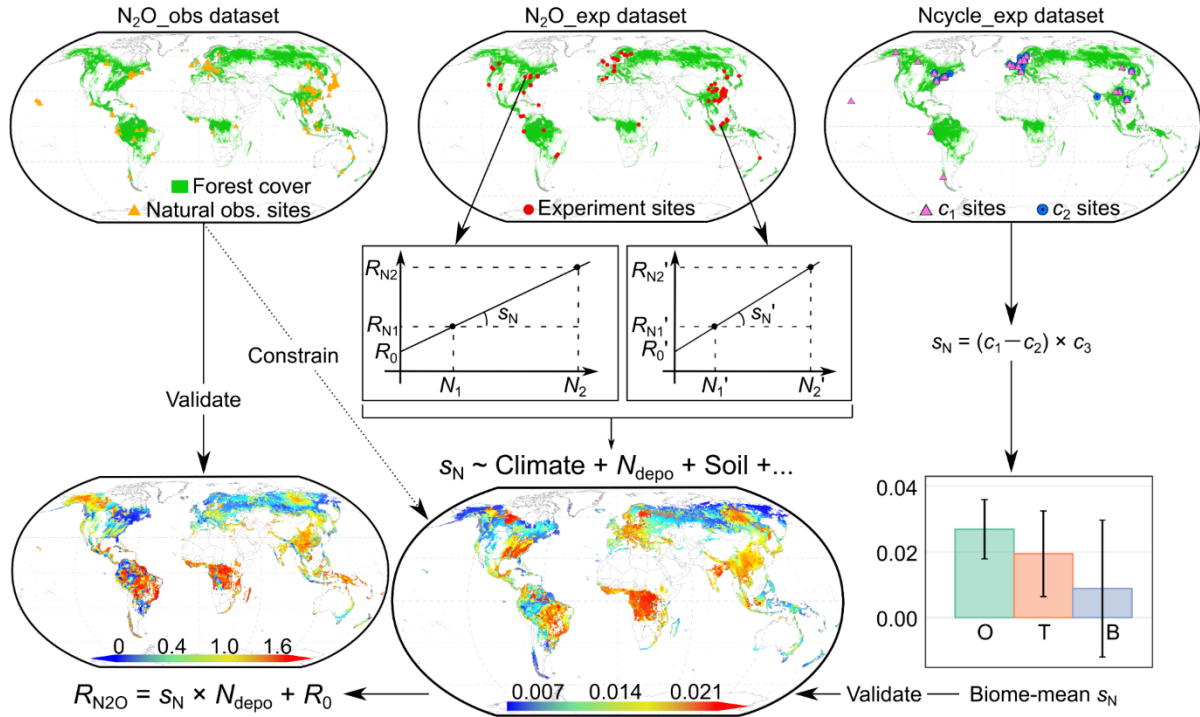
**Fig. S4.** Comparing soil CH<sub>4</sub> flux ( $R_{CH_4}$ ) estimated from different models built out of different training datasets. The sampling of observations to form a training (or testing) dataset was randomized by using different "seeds". Each seed corresponds to a determined set of samples, and different seeds lead to different samples. In this study, we randomly used seeds "1111" and "1234" for sampling. This analysis was to ensure that the estimated  $R_{CH_4}$  values were not dependent on which data were used for training and testing the models, so that the derived spatial pattern of  $R_{CH_4}$  was robust on grid level.



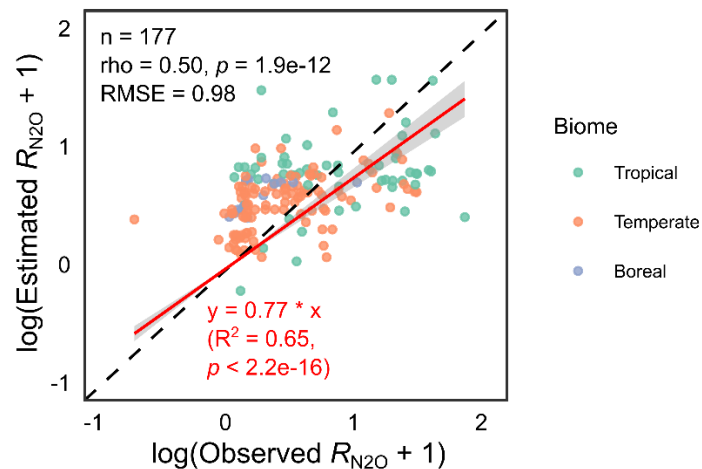
**Fig. S5.** Various forests are at different “stages” (in the stimulating-suppressing-weakening “three stages” framework), in accordance with the overall effects of N deposition on soil  $\text{CH}_4$  fluxes in the forests.



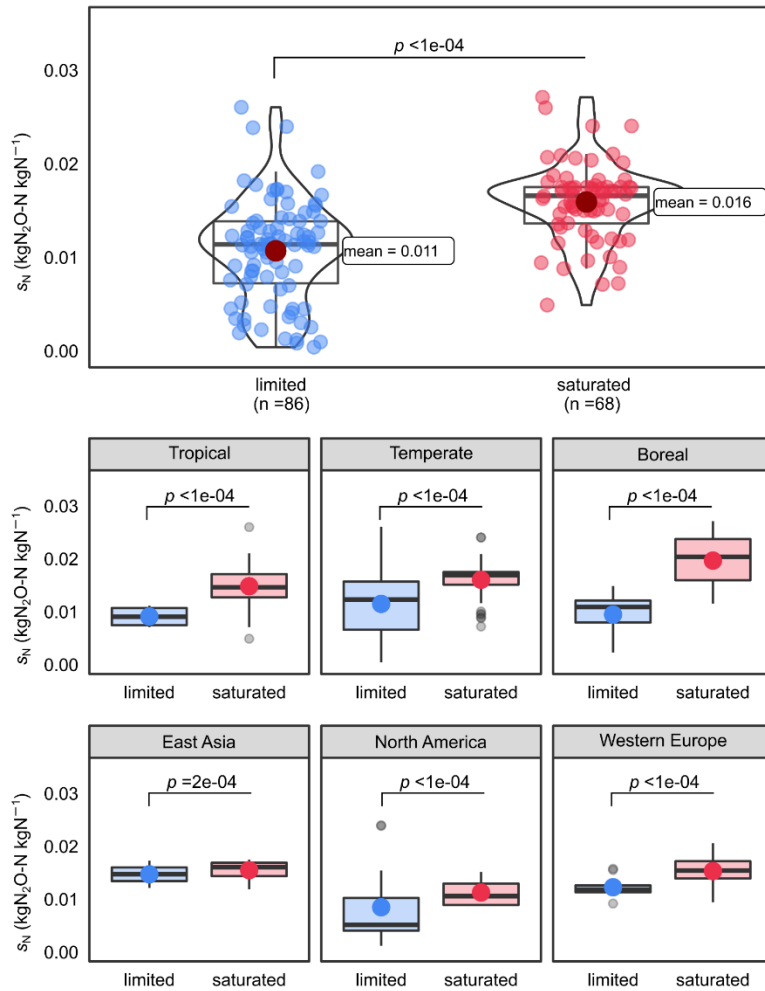
**Fig. S6.** Classified N-limited and N-saturated forests based on the sensitivity of soil  $N_2O$  emission to N deposition ( $s_N$ ) compared with field-observed N limitation or saturation status, with extra details in regions where field-observations were more abundant.



**Fig. S7.** Workflow illustrating the quantification and validation of the sensitivity of soil N<sub>2</sub>O emission to N deposition ( $s_N$ ) of global forests.  $N_1$  and  $N_2$  are different rates of low N input, and  $R_{N1}$  and  $R_{N2}$  are the corresponding soil N<sub>2</sub>O emission rates.  $N_{\text{depo}}$ : N deposition rate ( $\text{kgN ha}^{-1} \text{yr}^{-1}$ );  $R_0$ : background soil N<sub>2</sub>O emission rate ( $\text{kgN}_2\text{O-N ha}^{-1} \text{yr}^{-1}$ );  $c_1$ : sensitivity of N loss to N deposition ( $\text{kgN kgN}^{-1}$ );  $c_2$ : sensitivity of N leaching to N deposition ( $\text{kgN kgN}^{-1}$ );  $c_3$ : ratio of N<sub>2</sub>O to other gaseous end-products from nitrification and denitrification processes ( $\text{kgN}_2\text{O-N kgN}^{-1}$ ). O: Tropical; T: Temperate; B: Boreal.



**Fig. S8.** Comparing estimated and observed soil  $N_2O$  emission rates ( $R_{N_2O}$ ). Observations were aggregated to  $0.5^\circ \times 0.5^\circ$  grids to match with the spatial resolution of the environmental factors. Each point represents a grid-year. Points of different colors represent grid-years in different biomes. Dashed black line is the 1:1 line. The red line and fonts show a linear regression model on estimated and observed  $R_{N_2O}$ .



**Fig. S9.** Comparing the sensitivity of soil  $\text{N}_2\text{O}$  emission to N deposition ( $S_N$ ) of N-limited and N-saturated forests on global and regional scales.



**Table S1.** Parameters of segmented linear regression models on soil CH<sub>4</sub> flux ( $R_{CH4}$ ) and N input rate.

No.	Model ( $R_{CH4} \sim$ N input rate)	Parameters
①	$y = -0.037 \cdot x - 2.45$	$n = 53, R^2 = 0.01, p = 0.44$
②	$y = 0.045 \cdot x - 5.75$	$n = 49, R^2 = 0.06, p = 0.09$
③	$y = -0.004 \cdot x - 0.73$	$n = 29, R^2 = 0.00, p = 0.80$
④	$y = 0.096 \cdot x - 5.28$	$n = 121, R^2 = 0.10, p = 0.0003$
⑤	$y = -0.006 \cdot x - 1.53$	$n = 196, R^2 = 0.03, p = 0.02$

**Table S2.** Parameters of the constructed random forest regression models.

Model	$R_{CH4} \sim MAT + MAT.cv + MAP + MAP.cv + N_{depo} + N_{depo.cv} + Sand + Clay + s_N$	
	mtry	3
Parameters	ntree	1000
	Number of runs	1000

$R_{CH4}$ : soil CH<sub>4</sub> emission rate; MAT: mean annual temperature; MAP: mean annual precipitation;  $N_{depo}$ : mean annual N deposition; Sand: soil sand content; Clay: soil Clay content; MAT.cv, MAP.cv and  $N_{depo.cv}$  are the corresponding coefficients of temporal variation;  $s_N$ : sensitivity of soil N<sub>2</sub>O emission to N deposition, which indicates soil N limitation or saturation status. The predictors were selected based on mechanistic relevance and data availability.

**Table S3.** Linear models on soil N<sub>2</sub>O emission rate ( $R_{N_2O}$ ) and N input rate (model:  $R_{N_2O} \sim$  N input rate) built with low N input data (N addition rate  $\leq 150$  kgN ha<sup>-1</sup> yr<sup>-1</sup>) from global forest experiment sites, and the derived sensitivity ( $s_N$ ) of soil N<sub>2</sub>O emission to N deposition and background N<sub>2</sub>O emission rate ( $R_0$ ).

N o.	Longitude range	Latitude range	Biome	$s_N$	$R_0$	n	adj.R <sup>2</sup>	p value	References
1	(19,19.5)	(64,64.5)	Boreal	0.002	0.045	2	NA	NA	(Rutting et al., 2021)
2	(30.5,31)	(62.5,63)	Boreal	0.025	5.132	4	0.14	0.347	(Regina et al., 1998)
3	(22.5,23)	(62,62.5)	Boreal	0.013	0.538	2	NA	NA	(Ojanen et al., 2019)
4	(8,8.5)	(58.5,59)	Boreal	0.026	0.343	6	0.57	0.052	(Sitaula et al., 1995a, 1995b)
5	(-3.5,-3)	(55.5,56)	Temperate	0.02	0.258	6	0.18	0.224	(U. M. Skiba et al., 1998)
6	(-3,-2.5)	(55.5,56)	Temperate	0.006*	-0.009	6	0.73	0.019	(U. Skiba et al., 1999; U. M. Skiba et al., 1998)
7	(1.5,2)	(52.5,53)	Temperate	0.004	0.233	2	NA	NA	(U. M. Skiba et al., 1998)
8	(9.5,10)	(51.5,52)	Temperate	0.042*	0.51	10	0.48	0.015	(Borken et al., 2002; Brumme & Beese, 1992; Marife D Corre et al., 2003)
9	(128.5,129)	(47,47.5)	Boreal	0.015	0.777	11	0.02	0.300	(He, 2015; L. Song et al., 2017; Tian et al., 2018)
10	(8.5,9)	(47,47.5)	Temperate	0.003	-0.062	4	0.63	0.134	(Krause et al., 2013)
11	(-80.5,-80)	(43.5,44)	Temperate	0.009	1.374*	4	0.79	0.073	(Lutes et al., 2016)
12	(-72.5,-72)	(43,43.5)	Temperate	0.012	-0.216	2	NA	NA	(M. S. Castro et al., 1992)
13	(141,141.5)	(43,43.5)	Temperate	0.025	1.647	2	NA	NA	(Y. S. Kim et al., 2012)
14	(-72.5,-72)	(42.5,43)	Temperate	0.001	0.074	6	0.05	0.323	(Richard D. Bowden et al., 1991)
15	(128,128.5)	(42,42.5)	Temperate	0.01	0.67	2	NA	NA	(Geng et al., 2017)
16	(127.5,128)	(41.5,42)	Temperate	0.029	2.287	13	0.11	0.141	(Bai et al., 2014; Cheng et al., 2016; B. Peng et al., 2021)
17	(-80.5,-80)	(41.5,42)	Temperate	0.003	0.217	2	NA	NA	(R. D. Bowden et al., 2000)
18	(-4,-3.5)	(40,40.5)	Temperate	0.001*	0.026*	4	0.95	0.017	(Lafuente et al., 2020)
19	(112,112.5)	(36.5,37)	Temperate	0.056	2.754	3	0.98	0.068	(H. Yu, 2019)
20	(111,111.5)	(31.5,32)	Temperate	0.013**	0.483	27	0.28	0.003	(Zhaolan Lin, 2013; Zhaolan Lin et al., 2012; R. Wang, 2012; Xu et al., 2017)
21	(110,110.5)	(31.5,32)	Temperate	0.023	-0.31	4	0.54	0.166	(Pan, 2013)
22	(120.5,121)	(30.5,31)	Temperate	0.017	1.135	4	0.51	0.181	(Tu & Zhang, 2018)
23	(119.5,120)	(30,30.5)	Temperate	0.003	1.238***	16	0.01	0.308	(X. Chen, 2014; X. Chen et al., 2014; Ziwen Lin, 2019; X. Z. Song et al., 2020; Z. Wang, 2014)
24	(120,120.5)	(30,30.5)	Temperate	0.012**	0.834*	12	0.64	0.001	(J. Zhang, 2013; J. Zhang et al., 2013)
25	(106.5,107)	(29.5,30)	Temperate	0.025*	0.875*	3	1	0.018	(Xie et al., 2018)

26	(115.5,116)	(29.5,30)	Temperate	0.012	2.025	6	0.14	0.248	(C. Li et al., 2019)
27	(116.5,117)	(28,28.5)	Temperate	0.013	0.16	2	NA	NA	(Fan et al., 2020)
28	(118,118.5)	(27,27.5)	Tropical	0.015	1.948	9	0.12	0.190	(S. Chen, 2012)
29	(115,115.5)	(26.5,27)	Tropical	0.026***	-0.092	54	0.47	<0.001	(Dang, 2015; X. Li, 2017; X. Y. Li et al., 2015; Sun & Zhang, 2015; J. Wang, 2016; L. Wang, 2015; L. Wang et al., 2016; Y. Wang, 2015; Y. S. Wang et al., 2016; L. Zhang, 2013)
30	(117,117.5)	(26,26.5)	Tropical	0.007	0.5	3	0.55	0.313	(Wu, 2018)
31	(118,118.5)	(25.5,26)	Tropical	0.012	0.601	4	0.33	0.257	(Yuan, 2016)
32	(113,113.5)	(23.5,24)	Tropical	0.014	-0.226	3	0.77	0.220	(Cai, 2013)
33	(112.5,113)	(23,23.5)	Tropical	0.027*	0.19	22	0.15	0.041	(H. Chen et al., 2016; Gao et al., 2017; Mo et al., 2006)
34	(112.5,113)	(22.5,23)	Tropical	0.004	1.919***	14	0.11	0.129	(W. Zhang et al., 2014)
35	(106.5,107)	(22,22.5)	Tropical	0.012***	-0.038	8	0.84	0.001	(Hong, 2015)
36	(107,107.5)	(22,22.5)	Tropical	0.043*	-0.089	10	0.51	0.013	(R. Li et al., 2014, 2015; Yang, 2015; Kai Zhang et al., 2015)
37	(107.5,108)	(22,22.5)	Tropical	0.007**	0.589**	4	0.98	0.007	(K. Zhang et al., 2017)
38	(101,101.5)	(21.5,22)	Tropical	0.037	2.101*	9	0.18	0.144	(Yan, 2006; Zhou et al., 2016)
39	(110.5,111)	(21,21.5)	Tropical	0.018	3.195	3	0.69	0.256	(F. M. Wang et al., 2014)
40	(-80,-79.5)	(9,9.5)	Tropical	0.021**	0.674	8	0.71	0.005	(M. D. Corre et al., 2014; Koehler et al., 2009)
41	(-82.5,-82)	(8.5,9)	Tropical	0.019	1.063	8	0.32	0.083	(M. D. Corre et al., 2014; Koehler et al., 2009)
42	(116.5,117)	(6,6.5)	Tropical	0.007**	0.517**	10	0.61	0.005	(Hall et al., 2004)
43	(31.5,32)	(1.5,2)	Tropical	0.018***	1.756***	4	1	0.001	(Tamale et al., 2021)
44	(102,102.5)	(-1.5,-1)	Tropical	0.022**	0.919*	7	0.84	0.002	(Aini et al., 2015)
45	(-79.5,-79)	(-4,-3.5)	Tropical	0.005	0.135	3	0.44	0.356	(Muller et al., 2015)
46	(-79,-78.5)	(-4.5,-4)	Tropical	0.006	0.471	3	0.5	0.333	(Muller et al., 2015)
47	(-79.5,-79)	(-4.5,-4)	Tropical	0.006	-0.11	3	0.95	0.106	(Muller et al., 2015)

\*  $p < 0.05$ ; \*\*  $p < 0.01$ ; \*\*\*  $p < 0.001$ ; NA, not applicable

**Table S4.** Generalized linear models on environmental factors and the sensitivity ( $s_N$ ) of soil  $N_2O$  emission to N deposition and the background  $N_2O$  emission rate ( $R_0$ ).

	Estimate	SE	t	p
<b>Refined model on <math>s_N^{\dagger}</math></b> (Deviance explained = 91.1%, n=46)				
Clay	4.77E-03	1.83E-03	2.605	0.013*
Sand	3.15E-03	9.20E-04	3.419	0.001**
$\log(N_{\text{depo}})$	2.01E-02	1.14E-02	1.769	0.085
Clay $\times$ $\log(N_{\text{depo.cv}})$	2.13E-03	9.35E-04	2.282	0.028*
Sand $\times$ $\log(N_{\text{depo.cv}})$	1.17E-03	3.82E-04	3.056	0.004**
Clay $\times$ Sand	-1.90E-04	6.94E-05	-2.735	0.009**
Clay $\times$ Sand $\times$ $\log(N_{\text{depo.cv}})$	-1.14E-04	3.66E-05	-3.112	0.003**
<b>Refined model on <math>R_0^{\ddagger}</math></b> (Deviance explained = 43.2%, n = 45)				
$\log(N_{\text{depo.cv}})$	1.99E-01	9.56E-02	2.084	0.043*
MAT $\times$ Sand $\times$ Clay	3.04E-06	5.99E-07	5.072	0.000***
MAP $\times$ MAP.cv $\times$ $\log(N_{\text{depo}})$	-8.31E-04	2.91E-04	-2.854	0.007**

MAT: mean annual temperature; MAP: mean annual precipitation;  $N_{\text{depo}}$ : mean annual N deposition; Sand: soil sand content; Clay: soil clay content.

$s_N^{\dagger} \sim (\text{Clay} + \text{Sand} + \log(N_{\text{depo}}) + \text{Clay} \times \log(N_{\text{depo.cv}}) + \text{Sand} \times \log(N_{\text{depo.cv}}) + \text{Clay} \times \text{Sand} + \text{Clay} \times \text{Sand} \times \log(N_{\text{depo.cv}}))^2$

$R_0^{\ddagger} \sim \text{EXP}(\log(N_{\text{depo.cv}}) + \text{MAT} \times \text{Sand} \times \text{Clay} + \text{MAP} \times \text{MAP.cv} \times \log(N_{\text{depo}})) - 0.5$

\*  $p < 0.05$ ; \*\*  $p < 0.01$ ; \*\*\*  $p < 0.001$

**Data Set S1. (separate file)**

Compiled dataset on soil CH<sub>4</sub> flux from N addition experiments in global forests (CH<sub>4</sub>\_exp dataset in main text).

**Data Set S2. (separate file)**

Compiled data on soil CH<sub>4</sub> flux under natural conditions in global forests (CH<sub>4</sub>\_obs dataset in main text).

**Data Set S3. (separate file)**

Compiled dataset on soil N<sub>2</sub>O emission rate from N addition experiments in global forests (N<sub>2</sub>O\_exp dataset in Text S1).

**Data Set S4. (separate file)**

Compiled data on soil N<sub>2</sub>O emission rate under natural conditions in global forests (N<sub>2</sub>O\_obs dataset in Text S1).

**Data Set S5. (separate file)**

Compiled dataset on N loss rate, N leaching rate and change rate of soil N pool from N addition experiments in global forests (Ncycle\_exp dataset in Text S1).

**Data Set S6. (separate file)**

Compiled dataset on global forest N saturation status (limited or saturated) indicated by N leaching rate (Nleach dataset in Text S1).

**Data Set S7. (separate file)**

An existing dataset from Du et al. (2020) on global forest N saturation status (limited or saturated) indicated by plant growth response to N input (NuLi dataset in Text S1).

**Data Set S8. (separate file)**

Data on environmental factors (MAT, MAP, N deposition rate, etc.) in global forests, extracted from spatial datasets mentioned in Methods section.

**Data Set S9. (separate file)**

Global forest soil methane budgets estimated in previous studies.

**Code S1. (separate file)**

R code script used to carry out the data analysis processes, and produce the figures.